

DISSERTATION

RELATIONSHIPS AMONG ENVIRONMENTAL FACTORS AND
STREAM WATER ION YIELDS OF WATERSHEDS IN THE UNITED STATES

Submitted by:

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In partial fulfillment of the requirements

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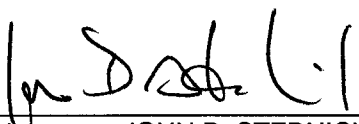
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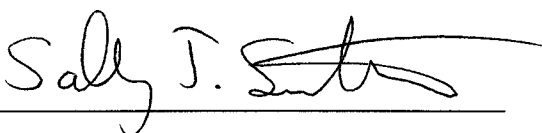

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Abstract of Dissertation

RELATIONSHIPS AMONG ENVIRONMENTAL FACTORS AND STREAM WATER ION YIELDS OF WATERSHEDS IN THE UNITED STATES

Understanding relationships among environmental factors and water quality constituents is very important for ion yield prediction, water quality management, and pollution prevention strategies. In this study, 13 stream water constituents – total dissolved solids, dissolved oxygen, pH, total nitrogen, orthophosphate, phosphorus, calcium, magnesium, sodium, potassium, sulfate, chloride, and bicarbonate from 47 watersheds in the United States were analyzed in relation to the environmental factors, including bedrock types (limestone, sandstone, and crystalline), land cover types, average annual precipitation, population density, watershed slope, and watershed area. Two regression analyses were applied, multiple linear regression using the “stepwise” method of model selection and principal component regression (PCR) followed by a prediction and validation process. Geographic information system (GIS) techniques were used to delineate watersheds, handle data, and examine the spatial distribution of the predicted ion yields.

The results of this study indicated that there was a strong exponential relationship among environmental factors and selected stream water ion yields. The most important factor affecting stream water ion yields was average annual precipitation, which either appeared uniquely in the multiple regression models or appeared in company with other factors, accounting for 12 to 94 percent of the variance of most annual ion yields. Bedrock type was the second importance for stream water ion yields. Some average annual ion yields of limestone watersheds are 2 to 5 times larger than those of sandstone

or crystalline watersheds, including total dissolved solids, bicarbonate, magnesium, and calcium. In addition, ion yields in limestone and sandstone watersheds are more closely related to the environmental factors than those in crystalline watersheds. Population density was the third most important factor, which accounted for 26 to 61 percent (in company with other factors) of the variance of ion yields of total dissolved solids (sandstone), sodium (limestone), magnesium (sandstone), potassium (limestone, sandstone), chloride (limestone), and bicarbonate (sandstone).

Land cover was less important. The effects of land cover were significant in company with other factors and accounted for 15 to 45 percent of the variances in the predictions of ion yields in limestone and sandstone watersheds, including orthophosphate, phosphorus, magnesium, and sulfate (limestone watersheds), total dissolved solids, sodium, and bicarbonate (sandstone watersheds). Watershed slope was an ineffective predictor, significant only in limestone watersheds and accounted (in company with other factors) for 12 to 42 percent of the variance of ion yields of sodium, potassium, magnesium, and dissolved oxygen. Watershed area was not significant in any ion yield regression analysis.

In this study, multiple linear regression (MLR) provided simpler models but larger R square, smaller standard error of estimate, and could detect more significant equations than principal component regression (PCR).

The results of this study can be helpful to understand effects and importance of environmental factors on water quality, to predict average annual stream water ion yields, especially for watersheds where water chemistry data are not available, and to evaluate the explanation of the predicted results contributed by each environmental factor. The

methods of this study may also be useful for those who want to develop their own models to predict water ion yields.

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I. INTRODUCTION

“It is important to understand the difficulties in identifying causes and, in particular, sources of pollution in impaired waters. For many waters, states and other jurisdictions classify the causes and sources of impairment as unknown. U. S. Environmental Protection Agency (EPA) and states are working to develop methodologies for both determining the causes and sources of impairment and describing the level of confidence in the classification.” (U. S. Environmental Protection Agency, 2000). The causes and/or the sources of pollution in impaired waters are not easy to be identified partly due to the fact that factors affecting water quality vary greatly across landscapes (Vose, 2006, as cited in USDA, 2006). In other words, the processes controlling solute concentrations in surface water are affected by many environmental factors, including bedrock formation, climate, topography, biota and soils, atmospheric deposition, hydrology, land use, time, and anthropogenic activities (Peters, 1984; Hem et al., 1990; Brown, 2000). It would be expected that properties and magnitude of the impacts of environmental factors on water quality are complicated and may change profoundly through time and space. While affecting water quality, environmental factors simultaneously affect each other. Water quality, therefore, is typically conceptualized as a reflection of the environmental factors occurring within a watershed (Stednick, 1991; Brown 2000).

For a long time, watershed scientists have attempted to find the "causes" of stream water ion yield variation. For example:

Why do water quality constituents differ considerably from this watershed to other

watersheds? Or what factors account for the fact that some watersheds have a higher/lower rate of constituent concentrations than others through the time? And

What can water quality managers do in terms of Best Management Practices to reduce non-point source water pollution in their watersheds?

To answer those questions, it is important to define and assess the effects of environmental factors on water quality. This work is necessary to set up inexpensive, effective but simple models for getting master plans for water pollution (especially for nonpoint source pollution) prevention strategies. Based on these research findings, water resource managers and environmental agencies can have a suitable time allocation, technologies (e.g. Best Management Practices), and enforcements to protect water resources properly. For example, a well known development strategy named “Low Impact Site Design” has been developed and implemented to minimize effects of environmental factors on water quality from stormwater based on a “set of “micromanagement”, nonstructural approaches for a more attractive, multifunctional landscape when compared to the conventional strategies that just rely on a “pipe-and-pond” approach.” (Massachusetts Low Impact Development Toolkit, 2006). In addition, to deal with nonpoint source pollution, the “Nonpoint Source Management Programs”, created by the Water Quality Act of 1987 (PL 100-4) through section 319 did require State governments to submit State Assessment Reports identifying significant nonpoint sources of pollution to the state’s navigable waters (U.S. G.P.O., 1988).

There have been numerous studies examining the relationship among specific aspects of water quality and environmental characteristics. However, “relatively few have

included the "common" ions (e.g., calcium and magnesium) and/or the entire range of environmental characteristics occurring within the United States" (Brown, 2000).

This study will examine the relationship between environmental factors (including geology (bedrock type), climatology (precipitation), anthropogenic (population), and land cover type) and 13 constituent ion yields, including dissolved oxygen, orthophosphate, phosphorus, pH, total nitrogen, sodium, potassium, magnesium, calcium, chloride, sulfate, bicarbonate, and total dissolved solids. This study also comprehensively takes advantage of the current techniques of multiple regression, principal component regression, and GIS-based models to investigate alternatives for prediction and delineation of annual ion loads on a spatially distributed basis.

II. LITERATURE REVIEW

Annual dissolved ion yields in streamwater are mainly derived from few sources, including bedrock and soil types of the drainage basin, atmospheric deposition, and substances derived from human activities (Hem, 1970, as cited in Peters, 1984). It is difficult to determine the influence of a specific source on a stream system due to the fact that the relationships between dissolved ion yields (distributed by specific sources) and the yields actually transported in streams are complex and variable (Peters, 1984).

In the United States, research projects aimed at assessing and defining the effects of environmental factors on water quality spatially and temporally have been conducted since the early 1970's (Brown, 2000). The most common examples of those projects are the National Water Summary Report like *Hydrologic Events and Stream Water Quality* (U.S. Geological Survey, 1993) and the National Water Quality Inventory (U.S. Environmental Protection Agency, 2000). The National Water Summary is a series of publications by the U.S. Geological Survey focused on increasing public understanding of the nature, quality, use, and trends of the nation's water resources. Each volume of the National Water Summary series includes general papers on the specific topic and state summaries. Having been prepared since 1975 under Section 305(b) of the Federal Clean Water Act, the *National Water Quality Inventory Report to Congress* is a biennial report to Congress and the public about the quality characteristics of our nation's water bodies. This becomes the primary vehicle for identifying and informing about water quality problems in rivers and streams, sources of pollutants and stressors impacting water quality in the United States (U.S. Environmental Protection Agency, 2000).

Of the most common sources of environmental effects, the bedrock type underlying a drainage basin is considered a major source of influences on the chemical composition of the stream water in that basin (Douglas, 1968, as cited in Peters 1984). Stream draining granite basins were reported to have dissolved solids concentration about 2 to 3 times greater than that in streams draining quartzite basins (Miller, 1961, as cited in Peters 1984). Plutonic, granite, or sandstone bedrock basins were reported to contain lower concentrations of dissolved solids than those of basins containing sedimentary, limestone or dolomite bedrocks (e.g., Johnson and Reynolds, 1977, as cited in Peters 1984). According to Hack (1960, as cited in Peters 1984), the lowest concentrations of solutes were found in streams draining quartzite basins and the highest were found in streams underlain by carbonate basins. In general, streams of noncalcareous and nonevaporitic basins drained the lowest dissolved solids concentration, while the highest dissolved solids concentrations were found in the sedimentary basins (Peters, 1984). Streams draining limestone bedrock typically produced ion yields that were approximately twice those of streams draining sandstone or crystalline bedrock (Brown, 2000).

The atmosphere is also a major source of ions in stream water (Peters, 1984). In basins underlain by bedrocks containing granites or quartzites, the ions found in stream water can be attributed mostly to atmospheric precipitation (Fisher, 1968, as cited in Peters 1984). In the Southwestern United States, high concentrations of sulfate in precipitation were reported due to the oxidation of H_2S produced by decaying organic matter (Junge and Werby, 1958, as cited in Peters 1984) while in the Northeastern United States, the most important source of sulfates found in streams was attributed to

the localized industrial areas (Fisher et al., 1968; Likens et al., 1977, as cited in Peters 1984).

The in-stream environmental factors also affect ion yield. For example, as water temperature increases, mineral solubility usually increases although the relationship is complicated by other effects. At higher temperatures, water in equilibrium contains less CO₂ than at lower temperatures. Cold water, therefore, may have more carbonic acid available for weathering than warm water (Peters, 1984). According to Sweeting (1966, as cited in Peters 1984), pH can decrease when stream temperature decreases, but carbonate mineral dissolution can increase. Streamflow also affects chemical concentrations. For example, most major stream solute concentrations were flow dependent, indicating a decreased concentration with increasing streamflow (Peters, 2004).

Forest land can play an important role and is among one of the most effective factors on water quality. Although the effects of forest land on runoff, stream flow, and erosion are both site- and species-specific (Calder, 2002), hundreds of research projects confirming the role of forest on improving water quality have been documented. For example, references and abstracts of more than 286 scientific articles particularly relevant to forest water quality were compiled in “A Forest Water Quality Literature Review” (Austin, 2006). Most of those articles were concerned with the effects of forest management and silviculture activities on water quality. In general, forest buffer strips are widely recommended for trapping nonpoint source pollutants and protecting surface water quality (Endreny, 2002). The presence of a forest or grass vegetative buffer predictably filtered between 40 and 95 percent of the runoff load where runoff carries

sediment and/or particulate phosphorus (Haycock et al., 1997; Reed-Anderson et al., 2000, as cited in Endreny, 2002). “Even a Mile of Forest Makes a Difference in Water Quality”, was stated by researchers after conducting research for the USDA Forest Service Southern Research Station on the effects of forest on water quality. The research findings “showed a definite "cleaning" affect on the stream from passing through a mile of National Forest, with evidence of significant reductions in concentrations of chemicals such as nitrates, ammonium, and phosphorous” (Vose, J. and Clinton, B., 2006, as cited in USDA).

Human activities can profoundly affect stream water composition by contributing constituents directly through both point and non-point sources; indirectly by disturbing the soils and underlying rocks, and by releasing gases and smoke into the air (Peters, 1984). These elements then go back to the stream water system via atmospheric deposition and precipitation. The effects of human activities on the composition and loads of dissolved constituents in streams are the subject of many studies. Among those activities, population density and land use data have been shown to affect stream water quality in many cases (Peters, 1984). Population density was reported to be correlated well with all constituents transported in the river except phosphorus and silica (Ceasar et al., 1976, as cited in Peters 1984). Road salting activities have been considered a principal source of sodium and chloride in streams (Peters and Turk, 1981; Mattson and Godfrey, 1994). The level of potassium was typically increased by runoff from fertilized land (Berner and Berner, 1987, as cited in Brown, 2000). Total dissolved solids were reported to be affected by a number of human activities such as irrigation return flow (Engberg, 1983, as cited in Brown, 2000) and acid mine drainage (Herlihy et al., 1990,

as cited in Brown, 2000). In their research (Day et al., 2002), the authors indicated that the degradation of water quality was more likely related to the increases in local populations and residential by-products rather than the increases in the associated infrastructure such as roads. The authors also suggested using the insights gained from their study to extrapolate the probable future impacts of growth on water quality for other watersheds that do not have the detailed data available.

The average yearly runoff of land covered by corn/soybean is about 4 times, and of commercial land in Central Indiana is about 24 times larger than that in forest (Frankenberger, undated). Therefore, the land use can affect water quality considerably. Land use was one of the most influential factors affecting nitrogen and phosphorous species in the Susquehanna River (Lystrom et al., 1978, as cited in Peters 1984). Disturbed land surface can cause considerably higher concentrations and loads of chemical constituents than undisturbed basins (Musser and Whetstone, 1964). According to Choi and Deal (2005), total nitrogen, total suspended particle, and total phosphorus in the year 2000 in five counties (Clinton, Jersey, Madison, Monroe, and St. Clair) in Illinois and five counties (St. Louis City, St. Louis, St. Charles, Jefferson, and Franklin) in Missouri were about 0.2 to 4.0 % in forest land, while those in agriculture land ranged from 55 to approximately 70%. Concentrations of sulphate increased with the loss of forest canopy by logging, and concentrations of nitrate increased due to increased nitrification of the forest floor caused by landscape excavation for drainage, reforestation, and liming (Havel, et al., 1999). Pollution from urban and agricultural land transported by precipitation and runoff was reported as the leading source of stream water impairment (U.S. Environmental Protection Agency, 2000). The increases of aquatic nutrient

concentration have been related to the use of phosphorus detergents, urbanization, disposal of untreated human and animal wastes, and the leaching due to applications of nitrogen and phosphorus fertilizers (Peters and Donohue, 2001).

Today, one of the most influential factors affecting hydrology and water quality is attributed to human activity. Anthropogenic activities such as urbanization, transportation, farming, deforestation, and afforestation affect both water quality and quantity by altering not only water balance but also processes controlling water quality (Peters, 2000). In addition to human influences such as land use, soil type also affects water quality. Effects of soil on water quality have been documented in many studies. For example, total N and total P in tributaries of clay soils were lower than those in tributaries of siliceous sandy soils (Donohue, et al., 2001).

Topographic condition is also an important factor that indirectly affects water quality and ion yield via effect of it on runoff and soil erosion. As described in the “rational” erosion-estimating equation (Smith and Whitt, 1948, as cited in Renard et al., 1996) and in the empirical Universal Soil Loss Equation (USLE) (Renard et al., 1996), not only slope, but also the length of slope is taken into account. Size of watersheds can affect water quality and ion yields as well. The effect of watershed size is indirectly reflected via travel time; in turn, this affects the decay and decomposition processes of the contributing substances draining to the stream water system. The effects of watershed size on water quality can also be reflected indirectly via sediment yield and transport. However, conflicting results have been reported in previous studies about effects of subwatershed sizes on sediment yield (Arabi et al., 2004). For example, Bingner et al. (1997) and Jha et al. (2004) (as cited in Arabi et al., 2004) indicated that sediment yield at

the outlet is very sensitive to the number and size of subwatersheds. However, FitzHugh and Mackay (2000) (as cited in Arabi et al., 2004) reported that observed sediment yield at the outlet was almost unaffected by the number and size of subwatersheds. Although slope and watershed size have not been extensively studied, they warrant attention because many ion concentrations are related to sediment and runoff transport.

One of the most comprehensive, national scale studies describing relationships among concentration or transport of a dissolved constituent in streams and various basin characteristics throughout the United States was conducted by Peters (1984). In this study, the relationships among seven major dissolved ions in streams – sodium, potassium, magnesium, calcium, chloride, sulfate, bicarbonate, and dissolved solids and four environmental factors including bedrock type, annual precipitation, population density, and stream temperature were examined by using linear and multiple linear regression analysis. Water ion yields were modeled as a linear function of the 4 environmental characteristics using ordinary least squares (OLS) regression. The general form of the linear regression equation was:

$$y_i = a + b_1x_1 + b_2x_2 + \dots + b_nx_n$$

where:

y_i = the annual yield of constituent i ,

x_1, x_2, \dots, x_n = the bedrock type (x_1) and other basin characteristics.

In his study, Peters treated a parallel slope coefficient for a certain basin characteristic in different bedrocks when taking the class variable (dummy variable) into

the multiple linear regression model to compare the effects of bedrock on ion yield. That assumption, however, may not be true. The author also combined the data from three bedrock types to do the regression analysis. Obviously, this approach is not as appropriate as doing regression analysis for each bedrock type separately.

Peters (1984) noted that, of the 4 factors analyzed, annual precipitation and the type of bedrock underlying a basin are the most important in determining the yields of dissolved major ions in streams. Population density was found to be effective in predicting only the yields of sodium, potassium, chloride, and sulfate in company with other variables. Stream water temperature had a negative effect on the yields of potassium, calcium, sulfate, bicarbonate, and dissolved solids. Peters suggested that the prediction of a constituent yield may be improved by determining the amount contributed from atmospheric deposition, and that stream water ion yield prediction may also be improved by taking more specific contributions from human activities, such as land use instead of population density in the multiple linear regression analysis. In addition, the prediction of ion yield can also be improved by developing multiple linear regression models for basins in each bedrock type separately (Peters, 1984).

Summary of the previous methods used to predict water quality concentrations and ion yields:

The various techniques used to identify the effects of environmental factors on water quality in large watersheds can be classified as one of two general approaches (Haith, 1976, as cited in Brown, 2000). The first approach is based on a particular set of environmental characteristics used to calculate the specified ion yields. In this approach,

the study area is first partitioned into discrete areas of relatively homogenous conditions of interest (e.g., land use/cover). Water samples are then collected to define ion concentrations for each area of concern. Total ion yields are calculated by multiplying the mean concentration by the runoff volume. This approach is inexpensive and offers a relatively rapid means of prioritizing watersheds for water quality remediation plans (Adamus and Bergman, 1995; Brown, 2000). Nevertheless, there are some limitations involved in this approach. For example, this approach is considered unable to predict ion yields for combinations of the environmental characteristics which are not included in the sampled population (Brown, 2000).

The second approach is to model water quality loading based on a range of environmental characteristics. In this approach, the locations that integrate combinations of the environmental conditions are considered for the site selection (not relatively homogenous areas of interest) (Brown, 2000). Water samples collected at each site are then used to define ion concentrations. The total ion yield is calculated as average concentration multiplied by runoff volume. A function of the environmental characteristics occurring within a watershed is then built to model the prediction of ion yields in that watershed. This approach can bring back an advantage overwhelming the first approach, that is, we would be able to predict ion yields for combinations of environmental characteristics which are not included in the sampled population (Brown, 2000).

Traditionally, the ordinary least squares (OLS) regression was used to analyze the relationship between dependent and independent variables. However, in many situations, data collected do not satisfy OLS regression assumptions. To deal with those cases,

alternative regression methods to OLS have been developed and they can provide greater predictive efficiency and comprehension of structure when data do not satisfy OLS regression assumptions (Efron and Tibshirani, 1991; Brown, 2000). Multiple linear regression analyses have been used in many studies to predict water concentration yield (Lystrom et al., 1978; Peters, 1984). Brown (2000) used a regression tree model to analyze the relationship between environmental factors and ion yields. The author noted that analogous to the OLS approach, regression tree models revealed the linear nature, non-constant variance, and the importance of variables observed in relationships between ion yields and environmental factors.

The uses of linear and multiple linear regression analysis are among aggregation approaches. These approaches can calculate predicted discrete single values of the water concentration of interest. However, using these approaches, we do not know the spatial variation of the project concentrations and ion yields.

Recent advances in geographic information systems (GIS) and remote sensing technologies have enabled us to build more realistic relationships among environmental factors and water concentration. These techniques include disaggregating approaches that can bring in both temporal and spatial results, rather than just predicted discrete single values. This is important, especially when we want to predict water ion yields for large watersheds.

Some of the most comprehensive models developed for predicting spatial distribution of water concentrations are useful due to applications of GIS techniques (Saunders, 1996; Dartiguenave, 1997; Wong et al., 1997; Naranjo, 1997; Quenzer, 1998;

Melancon, 1999; Osborne, 2000; Khalil, 2002). Although these GIS models were developed from different data and places, they were done in a similar way, that is, the loading of a given project constituent is a function of runoff and estimated or expected mean concentration (EMC) of that constituent. The expected mean concentrations for each predefined land use/land cover and corresponding runoff coefficient are derived from previous studies. Runoff coefficients can be also calculated using the Curve Number (CN) method (e.g. Khalil, 2002). Then the model uses a grid laid over the landscape, accounting for the constituent loadings and runoff derived from each cell. By tracing the flow of water from cell-to-cell basis, the movement of a predefined constituent over the landscape and into the river system is simulated (e.g. Naranjo, 1997).

In general, as reported in those studies (e.g., in Saunders, 1996), most of them have either some or all of the following limitations. *First*, the transport of constituents is pre-assumed to be conservative with no material loss or decay/decomposition. In fact, that assumption might be contrary with reality. *Second*, only average annual assessments are considered, runoff flow and constituent/pollutant loads are pre-assumed to be steady state parameters from year to year and within any year. *Third*, constituent concentrations on the land surface are considered to be directly related to general land use categories and do not vary with more specific types of land uses (e.g., from event to event). *Fourth*, some constituent loads are estimated due to data limitation. *Fifth*, most of those studies used Expected Mean Concentration or Estimated Mean Concentration (EMC) derived from each pre-defined land use type to estimate total loadings for that type of land use. Since the available research findings of EMC are still very limited, the application of these models is somewhat limited. In addition, using runoff and EMC to estimate total loadings

of ion yields is not easy to explain the contribution of environmental factors on the estimated results. Nevertheless, in spite of these assumptions and limitations, those GIS models still be useful to make the results easier to analyze in a visual format and can be used to represent both the observed and predicted concentration/ion yield data spatially in many cases.

HYPOTHESIS

Water quality can be predicted from watershed characteristics including land cover, population, precipitation, topology, and geology in the United States.

OBJECTIVES

(1) To develop and compare regression equations for annual ion yields measured at different watersheds located throughout the United States using water quality and environmental data available via the internet.

(2) To determine the importance (weight) of environmental factors on water quality constituent yields.

(3) To generate an alternative Principal Component Regression technique for a subset of watersheds to estimate annual water quality ion yields based on the environmental factors and compare the results from this technique with the results from the traditional multiple linear regression analysis approaches.

(4) To predict water quality and use the GIS techniques to display different areas affecting water quality in a subset of watersheds.

III. METHODOLOGY

The overall methodology used in this study (Figure 1) is described in detail in the following sections:

1. Bedrock types, station number and station geo-coordinates

Following an earlier study (Peters, 1984), 56 gaged stream stations in the conterminous United States and Hawaii having drainage basins underlain by limestone, sandstone, or crystalline bedrocks were selected for this study. However, due to the lack of available data, only 47 study areas were actually resampled (Table 2, Appendix 1). For each selected basin, the following data sets were compiled:

(1) Annual yields of major dissolved ions, bicarbonate (HCO_3), calcium (Ca), chloride (Cl), dissolved oxygen (DO), total dissolved solids (TDS), magnesium (Mg), phosphate (P), potassium (K), orthophosphate (PO_4), sodium (Na), sulfate (SO_4), total nitrogen, and pH.

(2) Average annual precipitation; (3) Mean population density; (4) Average slope; (5) Total basin area.

(6) National Land Cover Dataset.

Watersheds were selected throughout the United States to provide a wide range in stream water constituents and environmental factors. The three rock types, limestone, sandstone, and crystalline, were chosen due to the reasons as the following (Peters, 1984): limestone is relatively abundant (constituting 20 percent by volume of the world's

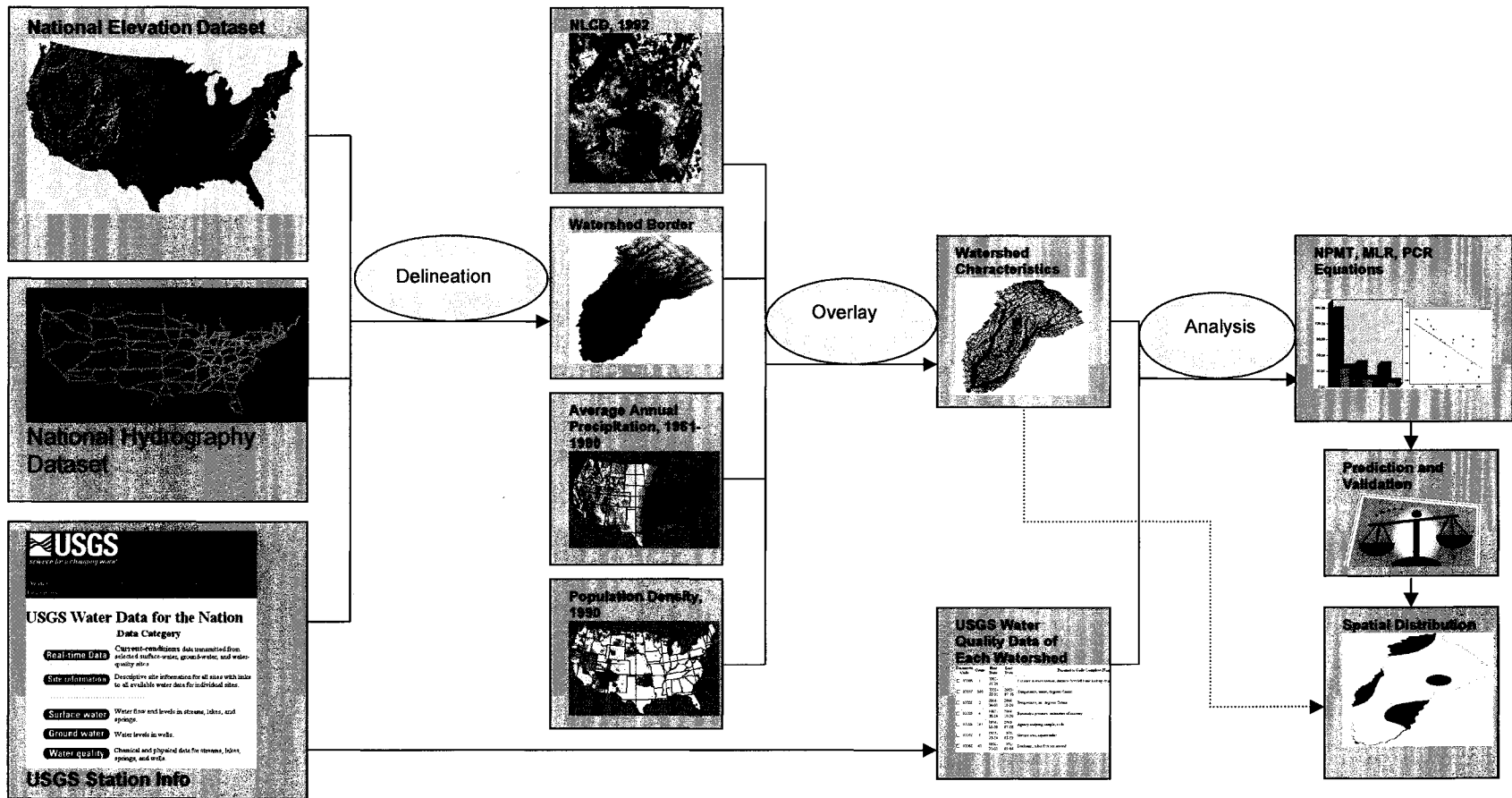


Figure 1. Methodology used for the study.

(NLCD: National Land Cover Dataset; NPMT: Nonparametric test; MLR: Multiple linear regression; PCR: Principal component regression)

sedimentary rocks) and it is composed primarily of easily weathered minerals of calcium and magnesium; sandstone consists of silicate minerals, which are more resistant to weathering than carbonates; crystalline rock constitutes 75 percent by volume of the Earth's crust and occupies the widest range in mineral texture and composition (crystalline was chosen for contrast with the other two types). Of the 47 basins selected, 19 were limestone, 11 were sandstone, and 17 were crystalline. Locations of USGS selected gage stations and rock types of the drainage basin are presented in Figure 2.

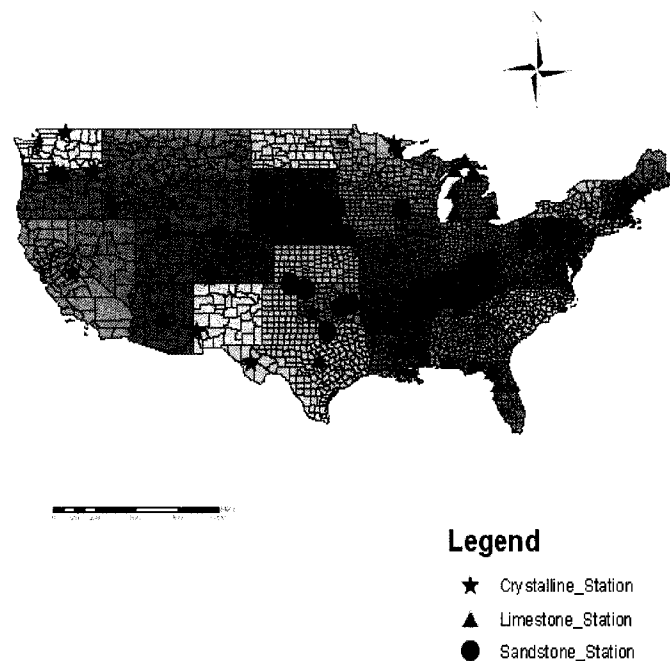


Figure 2. Locations of the selected USGS gaged stream stations for drainage watersheds underlain by single bedrock types in the United States.

Since there were many large basins containing a mixture of rock types, the conceptual determination of each bedrock type was based on Peters (1984) rationales (1) if a basin included more than 50 percent of the bedrock containing carbonate, then it was

designated as limestone basin; (2) if the carbonate minerals were virtually absent, the basin was designated as sandstone; (3) a basin would be designated as crystalline if it contained more than 80 percent metamorphic or igneous rock.

The selected stations and corresponding compiled data representing their basin characteristics (latitude, longitude, HUC, and drainage area in square miles) are summarized in Table 2 and in Appendix 1.

2. Sources and periods of selected data

Based on available data accessible via the internet, and to make the study consistent and compatible among the data types, the data were designed to compile as close to the same source and the same period as possible (Table 1). For metadata information, please go to the corresponding website (Table 1 and in the list of Literature Cited). These data were originally collected for multiple purposes, and this study uses them as secondary data based on EPA's Information Quality Guidelines (IQG) (EPA, 2003).

Table 1. Data sources for the study.

	Basin stations And Bedrock Types	Basin and sub-basin border	National Elevation Dataset	Land Cover	Average Annual Precipitation (inch)	Population Density (People/square mile)	Water Quality Constituents
Data Source	Peters, N., 1984 U.S. Geological Survey	National Hydrography Dataset	U.S. Geological Survey (1 Arc Second)	U.S. Geological Survey	U.S. Geological Survey	U.S. Geological Survey National Atlas of the United States	U.S. Geological Survey
Time/ Periods	1992	1961-1990	1990	1980-1990*
Internet Address	http://waterdata.usgs.gov/nwis	http://nhd.usgs.gov/	http://seamless.usgs.gov/website/seamless/viewer.Php http://www.mrlc.gov	http://seamless.usgs.gov/website/seamless/viewer.php http://www.mrlc.gov	http://datagateway.nrcs.usda.gov/GatewayHome.html	http://www.nationalatlas.gov	http://water.usgs.gov/watch/ http://waterdata.usgs.gov/nwis

(*: See Appendix 2 for more details).

Table 2. General information of the selected USGS gaged stream stations in the United States.

Site No.	Station No.	Station name	Lat (NAD27)	Long (NAD27)	HUC	Drainage Area (SQMI)
Limestone Bedrock						
1	02313000	WITHLACOCOCHEE RIVER NR HOLDER, FLA.	28°59'19"	82°20'59"	03100208	1,820.00
2	02320500	SUWANNEE RIVER AT BRANFORD, FLA.	29°57'20"	82°55'40"	03110205	7,880.00
3	02329000	OCHLOCKONEE RIVER NR HAVANA, FLA.	30°33'14"	84°23'03"	03120003	1,140.00
4	02368000	YELLOW RIVER AT MILLIGAN, FLA.	30°45'10"	86°37'45"	03140103	624
5	03085000	Monongahela River at Braddock, PA	40°23'28"	79°51'30"	05020005	7,337.00
6	03254000	LICKING RIVER AT BUTLER, KY	38°47'22"	84°21'05"	05100101	3,385.00
7	03290500	KENTUCKY RIVER AT LOCK 2 AT LOCKPORT, KY	38°26'20"	84°57'48"	05100205	6,180.00
8	03425000	CUMBERLAND RIVER AT CARTHAGE, TN	36°14'53"	85°57'19"	05130201	10,690.00
9	03543005	TENNESSEE RIVER AT WATTS BAR DAM (TAILWATER), TN	35°37'13"	84°47'00"	06010201	17,310.00
10	03571850	TENNESSEE RIVER AT SOUTH PITTSBURG, TN	35°00'41"	85°41'51"	06030001	22,640.00
11	03593005	TENNESSEE RIVER AT PICKWICK LANDING DAM (LL), TN	35°03'54"	88°15'08"	06040001	32,820.00
12	04045500	TAHQAMENON RIVER NEAR PARADISE, MI	46°34'30"	85°16'10"	04020202	790
13	04057004	MANISTIQUE RIVER ABOVE MANISTIQUE, MI	45°58'18"	86°14'35"	04060106	1,445.00
14	04108690	KALAMAZOO RIVER AT SAUGATUCK, MI	42°38'50"	86°11'53"	04050003	2,020.00
15	04126520	MANISTEE R AT MANISTEE, MI	44°15'02"	86°19'09"	04060103	1,928.00
16	04132052	CHEBOYGAN R (POND) AT LINCOLN AVE AT CHEBOYGAN, MI	45°38'02"	84°28'52"	04070004	1,500.00
17	04165500	CLINTON RIVER AT MORAVIAN DRIVE AT MT. CLEMENS, MI	42°35'45"	82°54'32"	04090003	734
18	08103900	S Fk Rocky Ck nr Briggs, TX	30°54'41"	98°02'12"	12070203	33.3
19	13018300	CACHE CREEK NEAR JACKSON, WY	43°27'08"	110°42'12"	17040103	10.6
Sandstone Bedrock						
1	01545600	Young Womans Creek near Renovo, PA	41°23'22"	77°41'28"	02050203	46.2
2	03049625	Allegheny River at New Kensington, PA	40°33'52"	79°46'22"	05010009	11,546.00
3	03237280	Upper Twin Creek at McGaw OH	38°38'37"	83°12'57"	05090201	12.2
4	08452000	WHITE R NEAR OACOMA, SD	43°44'54"	99°33'22"	10140204	10,200.00
5	06478500	JAMES R NEAR SCOTLAND, SD	43°11'09"	97°38'07"	10160011	20,653
6	07026000	OBION RIVER AT OBION, TN	36°15'04"	89°11'33"	08010202	1,852.00
7	07231500	Canadian River at Calvin, OK	34°58'40"	96°14'36"	11090202	27,952
8	07234000	Beaver River at Beaver, OK	36°49'20"	100°31'08"	11100102	7,955
9	07237500	North Canadian River at Woodward, OK	36°26'12"	99°16'41"	11100301	11,589
10	07245000	Canadian River near Whitefield, OK	35°15'50"	95°14'21"	11090204	47,576
11	06438000	BELLE FOURCHE RIVER NEAR ELM SPRINGS, SD	44°22'11"	102°33'56"	10120202	7,210
Crystalline Bedrock						
1	01059400	ANDROSCOGGIN RIVER AT BRUNSWICK, ME	43°55'03"	69°58'25"	01040002	3,434.00
2	01066000	Saco River at Cornish, Maine	43°48'28.94"	70°46'53.54"	01060002	1,293.00
3	04014500	BAPTISM RIVER NEAR BEAVER BAY, MN	47°20'15"	91°12'02"	04010101	140
4	05124480	KAWISHIWI RIVER NEAR ELY, MN	47°55'22"	91°32'06"	09030001	254
5	08623800	ENCAMPMENT RIVER AB HOG PARK CR, NR ENCAMPMENT, WY	41°01'25"	106°49'27"	10180002	72.7
6	07311200	Blue Beaver Creek near Cache, OK	34°37'24"	98°33'48"	11130203	24.6
7	08431700	Limpia Ck abv Ft Davis, TX	30°36'48"	104°00'04"	13070005	52.4
8	09352900	VALLECITO CREEK NEAR BAYFIELD, CO.	37°28'39"	107°32'35"	14080101	72.5
9	09430600	MOGOLLON CREEK NEAR CLIFF, NM	33°10'00"	108°38'57"	15040001	69
10	09508300	WET BOTTOM CREEK NEAR CHILDS, AZ.	34°09'39"	111°41'32"	15060203	36.4
11	11264500	MERCED R A HAPPY ISLES BRIDGE NR YOSEMITE CA	37°43'54"	119°33'28"	18040008	181
12	12447390	ANDREWS CREEK NEAR MAZAMA, WA	48°49'23"	120°08'41"	17020008	22.1
13	13169500	BIG JACKS CREEK NR BRUNEAU ID	42°47'06"	115°59'00"	17050102	253
14	13331500	MINAM RIVER NEAR MINAM, OR	45°37'12"	117°43'32"	17060105	240
15	14048000	JOHN DAY RIVER AT MCDONALD FERRY, OR	45°35'16"	120°24'30"	17070204	7,580.00
16	14103000	DESCHUTES RIVER AT MOODY, NEAR BIGGS, OR	45°37'20"	120°54'16"	17070306	10,500.00
17	14113000	KLICKITAT RIVER NEAR PITT, WA	45°45'24"	121°12'32"	17070106	1,297.00

(Source: <http://waterdata.usgs.gov>)

3. Watershed/Basin delineation

The watershed delineation process was based on Desktop GIS software (9.1) and the latest Hydrology Modeling Module/Extension made by ESRI (ESRI, 2005) in combination with the reference from basin and sub-basin borders of the National Hydrography Dataset (<http://nhd.usgs.gov>).

Watershed delineation was the most time-consuming process of the data preparation. The process is summarized as below:

(1) First get the geo-coordinates of the gaged-stream stations, corresponding hydrologic unit code (HUC), corresponding basin drainage area by entering the station number into the “Site information” link shown on the USGS website (<http://waterdata.usgs.gov/nwis>). Then, summarize this information in a Table (e.g. Table 2).

(2) Get a “draft” basin area by entering the HUC into the National Hydrography Dataset (<http://nhd.usgs.gov>). For large and very large basins (e.g. Canadian River near Whitefield, OK, station number: 07245000), it is necessary to get the basin area of a lower digit number of the HUC (e.g. 2 or 4 digit HUC) to make sure that the “draft” area is big enough to contain the actual area of the basin wanted.

(3) Get the coordinates of the “draft” border gathered in step 2, and enter them into the USGS website at <http://seamless.usgs.gov/website/seamless/viewer.php> or <http://www.mrlc.gov> and download the corresponding National Elevation Dataset (NED)

(4) Use the geo-coordinates of each station to create a corresponding point shapefile in ArcMap.

(5) Use Desktop GIS software and Hydrology Modeling extension to delineate the watershed based on the National Elevation Data gathered in step 3 and the point shapefile feature created in step 4.

The delineated watershed area will be compared with the corresponding area from USGS (in Table 2) by calculating the area variation in percent. The formula for this is expressed as:

$$\text{Area Difference (\%)} = (\text{Delineated area} - \text{USGS area}) * 100 / (\text{USGS area}) \quad (1)$$

The watersheds delineated in the process are very important input to get other GIS data such as population, precipitation, and land cover.

4. Data calculations for the stream water constituents

The raw stream water ion data were downloaded from a USGS website at <http://waterdata.usgs.gov>. The data were then calculated for average values per record and compiled in a table. For pH, instead of the mean, the median was calculated because of the fact that a change of one pH unit represents a tenfold change in hydrogen ion concentration (Project SEARCH, undated). This means that the pH distribution is right skewed, therefore, a median pH can be more appropriate to be expressed rather than a mean pH.

The formula for converting a measure of a constituent from mg/L to kg/ha/year is

expressed as:

$$\text{Ion yield of a constituent (kg/ha/yr)} = 3.45 \times \text{Average number of mg/L of the constituent} \times \text{Average number of instantaneous discharge (cfs)/ Drainage area of the station} \quad (2)$$

(3.45 = 365 x 60 x 60 x 24 x 0.3048 x 10/27878400; 1 square mile = 27878400 square feet; 1 foot = 0.3048 meter); (constant of 10 = 10000/1000)

Due to lack of available data, dissolved solids were calculated indirectly from the specific conductance by multiplying specific conductance by a constant of 0.6 (Stednick, 1991). Also, bicarbonate was calculated by multiplying alkalinity as CaCO₃ by a constant of 1.22 (Stednick, 1991).

5. GIS data preparation and calculations for the environmental factors

GIS data layers of land cover, precipitation, population are delineated based on the corresponding delineated watershed using Arctoolbox of Desktop GIS. After those layers are derived, they were exported to *.dbf files, to calculate the data as follows:

(1) For land cover data:

Originally, these data were divided into 21 classes/categories (Appendix 3a_4). However, as stated, all classes may not be represented in a specific state data set. The class numbers representing the digital value of the class in the data set are as the followings: (http://sain.nbio.org/downloads/regional_map/gisdata/nlcd_faq.shtml):

Water: 11 Open Water; 12 Perennial Ice/Snow

Developed: 21 Low Intensity Residential; 22 High Intensity Residential; 23 Commercial/Industrial/Transportation

*Barren: 31 Bare Rock/Sand/Clay; 32 Quarries/Strip Mines/Gravel Pits;
33 Transitional Vegetated;*

Natural Forested Upland:

41 Deciduous Forest; 42 Evergreen Forest; 43 Mixed Forest

Shrubland: 51 Shrubland

Non-natural Woody: 61 Orchards/Vineyards/Other

Herbaceous Upland: 71 Grasslands/Herbaceous

*Herbaceous Planted/Cultivated: 81 Pasture/Hay; 82 Row Crops;
83 Small Grains; 84 Fallow; 85 Urban/Recreational Grasses*

Wetlands: 91 Woody Wetlands; 92 Emergent Herbaceous Wetlands

The area of each land cover class (in square mile – SQMI) is calculated from the dbf file using the following equation:

$$\text{Area of a land cover class (SQMI)} = \frac{\text{Number of counts of that class} \times \text{Area of the watershed}}{\text{Total count of the watershed}} \quad (3)$$

One example of area calculation of each land cover class is presented in Table 3 as follows:

Table 3. Land cover classification and corresponding area of each class for station number 06478500 - JAMES R NEAR SCOTLAND, SD (sandstone).

Station number: 06478500 - JAMES R NEAR SCOTLAND,SD (Sandstone)				
ObjectID	Value of land cover type	Count	Area fraction (%)	Area (SQMI)
0	11	680019	0.8	176.2
1	21	105367	0.1	27.3
2	22	28343	0.0	7.3
3	23	169145	0.2	43.8
4	31	363	0.0	0.1
5	32	1462	0.0	0.4
6	33	24	0.0	0.0
7	41	685083	0.8	177.5
8	42	7194	0.0	1.9
9	43	819	0.0	0.2
10	51	41613	0.1	10.8
11	71	12931117	16.0	3350.3
12	81	21736287	26.9	5631.6
13	82	39021698	48.2	10110.1
14	83	2633738	3.3	682.4
15	84	43169	0.1	11.2
16	85	18262	0.0	4.7
17	91	22358	0.0	5.8
18	92	2825651	3.5	732.1
Total		80951712	100.0	20973.7

Since 21 classes of land cover are not practical for regression and model selection processes, the classes were recombined into 4 categories only, namely *Water*, *Developed*, *Forest*, and *Agriculture*, in which:

Water = land cover values of (11 + 12 + 91 + 92).

Developed = land cover values of (21 + 22 + 23 + 31 + 32 + 33).

Forest = land cover values of (41 + 42 + 43 + 51) (note: 51 is shrubland but it is more reasonable to add this class into Forest than into other categories).

Agriculture = land cover values of (61 + 71 + 81 + 82 + 83 + 84 + 85)

One example of the recombination of land cover classes into new categories is illustrated in Table 4 as below:

Table 4. Recombination of land cover classes into 4 categories for station number 06478500 - JAMES R NEAR SCOTLAND, SD (sandstone).

Station number: 06478500 (sandstone, square mile)				
Water	Developed	Forest	Agriculture	Total area (SQMI)
914.1	78.9	190.4	19790.3	20973.7

The unit of area is in square mile (SQMI).

(2) For precipitation data:

These average data were compiled based on the dbf data file exported from the precipitation layer of each watershed using *area weighted fractions* and the range of precipitation for that area.

$$\text{Annual Average Precipitation} = \sum \text{Area of range}(i) \times \text{Precipitation of range}(i) / \text{Total area}, \quad (4)$$

(i runs from 1 to n).

The unit of average annual precipitation is in inch (in).

One example of precipitation calculation is indicated in Table 5 as below:

Table 5. Average annual precipitation for basin station number 06478500 (sandstone).

Station number: 06478500 - JAMES R NEAR SCOTLAND,SD (sandstone)			
Range of Precipitation (In)	Area in square degree	Area fraction	Precipitation (in)
21	1.3	0.2	4.3
19	3.2	0.5	9.9
17	1.2	0.2	3.2
23	0.6	0.1	2.0
Total	6.2	1	19.4

(3) For population data:

These data were compiled using the spatial summary of the GIS population layer (or can be calculated in the same way as precipitation data were. The two results are almost the same).

The unit of population is an average of people/square mile (people/SQMI).

(4) For slope data:

The average slope of each watershed is calculated by using the slope function of

the Spatial Analysis in ArcGIS.

The unit of average slope is in degrees (°).

(5) For watershed area data:

These data are available from delineated watersheds using area function of ArcGIS. The unit of watershed area is in square mile (SQMI).

6. Building regression models for estimating constituent yield

Empirical relationships between ion yield and the selected environmental characteristics (independent variables) were examined. Various multivariate techniques are available to quantify the relation between the ion yields and those independent variables, the linear and multiple linear regressions were preferred for this study. These techniques were considered most suitable to provide quantitative information on the relationship between the environmental characteristics and stream ion yields (Lystrom et al., 1978; Peters, 1984).

The general form of a multiple linear regressions (after Peters, 1984) can be expressed as:

$$Y_i = a + b_1X_1 + b_2X_2 + \dots + b_nX_n; \quad (5)$$

Where

Y_i = the annual yield of constituent,

$X_1, X_2 \dots, X_n$ = the basin characteristics.

All regressions and other related procedures were performed by using SPSS version 13.0 (SPSS, 2004).

As tested in this study, most distributions of stream ion concentrations and the environmental factors are non-normal with a p-value less than 0.05. These results are consistent with the research findings of Davis (1973, as cited in Peters, 1984). Hydrologic data are positively skewed, meaning that values smaller than the mean are more common than values greater than the mean (Gray, 1970, as cited in Peters, 1984).

The distribution of log-transformed data is closer to normal than the distribution of original data, and the linear regression model works better with normal variables (SPSS, 2004). Therefore, these data were transformed into the common logarithm transformation (base 10) (Peters, 1984). In addition, for any new recombined categories of the land cover data with a value of zero, a value of 0.000347 SQMI is replaced to make the log transform possible (the original spatial resolution of the Landsat-5 TM images, which the land cover data were derived, was 30 meters or $30 \times 30 = 900$ square meters, or equal to 0.003474 SQMI). Therefore, any areas less than that value are not considered due to unclear spectrals influenced by other land cover types and objects. Bedrock types were treated as class/dummy variables in the multiple linear regression analysis (Peters, 1984).

The form of the log-transformed regression was:

$$\text{Log } Y_i = \log a + b_1 D_1 + b_2 D_2 + b_3 \log X_1 + b_4 \log X_2 + \dots + b_n \log X_n \quad (6)$$

Where:

Y = Average annual ion yield,

D₁ and D₂ = rock types (since there were 3 bedrock types, only two dummy variables need to be entered, the other is implied to be the control).

and X₁, etc. = remaining basin characteristics.

7. Multicollinearity diagnostics

Multicollinearity can occur when there are near-constant linear functions of two or more of the predictor variables. If multicollinearities exist, the variances of the estimated regression coefficients may become very large. That may lead to unstable estimates of the regression equation. Several approaches have been proposed to overcome the multicollinearity problem. For example, the equation does not contain multicollinearities if only a subset of the predictor variables are used (Jolliffe, 2002).

In this study, multicollinearity among the independent variables can be detected using the correlation matrix and the regression collinearity diagnostics based on the “Enter” method of the regression model selection (Appendices 4, 5, 6). In more detail, this can be based on the following indicators (SPSS, 2004):

(1) Based on the tolerance:

Tolerance equals to $1 - R^2_i$, “where R^2_i is the squared multiple correlation of the i^{th} variable with other independent variables” (Liu et al., 2002). In other words, “the tolerance is the percentage of the variance in a given predictor that cannot be explained by the other predictors”. If the tolerances are close to 0, there is high multicollinearity

and the standard error of the regression coefficients will be inflated.” (SPSS, 2004).

(2) Based on the Variance Inflation Factor (VIF):

It is usually considered problematic if variance inflation factor greater than 2 (SPSS, 2004). In fact, VIF is reciprocal of tolerance. The variables with low tolerance tend to have large VIF values. Variables with low tolerance and large VIF, therefore, suggest that they have a collinearity (Liu et al., 2002).

(3) Based on the initial total eigenvalues:

If the initial total eigenvalues in the collinearity diagnostics are close to 0, the predictors are highly intercorrelated and “small changes in the data values may lead to large changes in the estimates of the coefficients” (SPSS, 2004).

(4) Based on the Condition Indices:

“The condition indices are computed as the square roots of the ratios of the largest eigenvalue to each successive eigenvalue” (SPSS, 2004). If condition index value is greater than 15, there is a possible problem with collinearity. When condition index value is greater than 30, it is considered a serious problem with collinearity (SPSS, 2004).

There are two ways to minimize the multicollinearity problem used in this study. First, use the stepwise method of model selection to include only the most useful variables in the model (SPSS, 2004) and therefore, multicollinearity among independent variables can be reduced.

Second, apply principal component regression. In a principal component regression, the independent variables are, not separate variables per se, but the components of all original input independent variables. The principal component regression analysis does not only make collinearity diagnosis for each independent variable, but also solve the collinearity problem (Liu et al., 2002).

To do a principal component regression (PCR), it is necessary to extract components by implementing the principal component analysis (PCA). Since the measurement scale of the independent variables in this study is different, a correlation matrix (not covariance matrix) was used for the extraction analysis. In addition, for ease of interpretation, a rotated component matrix was performed. The rotated component matrix is useful to determine what the components represent (SPSS, 2004).

In a principal component analysis, initial total eigenvalues is the expression of the amount of variance in the original variables accounted for each component (SPSS, 2004). “For the initial solution, there are as many components as variables, and in a correlations analysis, the sum of the initial total eigenvalues equals to the number of components.” (SPSS, 2004). Since one of the aims of this study was to model the most effective components making the most considerable contributions on the variation of the dependent variables (ion yields), only initial eigenvalues over 1.0 were extracted and taken into account for the regression performance.

For compatibility with the multiple linear regressions, principal component regressions also used the log transformed data and the stepwise method of the model selection. The form of this regression model is as below:

$$\text{Log } Y_i = a + b_1 C_1 + b_2 C_2 + \dots + b_n C_n \quad (7)$$

$$C_i = \sum a_{ij} Z(X'_j) \quad (i = 1, \dots, n ; j = 1, \dots, n) \quad (8)$$

$$Z(X'_j) = (X_j - \bar{X}_j) / S_{x_j} \quad (j = 1, \dots, n) \quad (9)$$

Where,

Y = Average ion yield,

b_i = Slope coefficient of the i^{th} principal component

C_i = The i^{th} principal component,

n = Number of independent variables,

$Z(X'_j)$ = The j^{th} standardized independent variable (Z-score) of the j^{th} independent log-transformed variable (X_j),

\bar{X}_j = The mean of the j^{th} log-transformed independent variable,

S_{x_j} = The standard deviation of the j^{th} log-transformed independent variable,

a_{ij} = The coefficient of the principal component coefficient matrix.

Once the regression in equation (7) is created, it is possible to transform those C_i back to $Z(X'_j)$ and then to the original X_j

8. Determining the most effective factors on constituent yields

The factors with significant slope coefficients in the regression equation are more important than those not. The higher the standardized slope coefficient (beta), the more

the variation of the dependent variable is accounted for (SPSS, 2004). The adjusted R square, part, partial, and statistical significance (p) are also useful to determine the importance of the environmental factors. Since regression coefficient R square is affected by the number of independent variables and the corresponding sample size of the model, the adjusted R square (not R square) is usually used to compare the goodnesses of fit between different linear models (Liu et al., 2002).

9. Examining the effects of bedrock types on ion yields

One way to examine the effects of bedrock types on ion yields is to use comparison tests. If ion yields of a same constituent for any two bedrock types are significantly different, then the difference of the bedrocks can be accounted for the difference of that constituent.

Since the sample size of each bedrock type is not large (limestone -19, sandstone – 11, and crystalline – 17) and the distribution of the data is non-normal, a non parametric comparison using Mann Whitney U Test in SPSS (2004) was conducted to examine the differences between constituents of any two bedrock types.

Another way to examine the effects of bedrock types on ion yields is to use bedrock types as class (dummy) variables in the multiple linear regression analysis (Equation 6). However, equation 6 is compatible for the assumption that the slope coefficients of the same independent variables are the same in the two equations. This assumption is not appropriate in many circumstances. Therefore, that model is good for testing the significant differences of the intercepts only.

Since most of the differences among constituents of the three bedrock types were significant, equation (6) was used to fit for each type of bedrock separately, meaning that the class variable was not taken into account. However, if a same independent variable was detected in any two equations (of different bedrock types) for the same constituent, a comparison test was performed to test the significance of that coefficient slope and the intercept. This is another way to test the effects of bedrock types on ion yields. The general form for this regression can be expressed as:

$$\begin{aligned} \text{Log } Y_i = & \log (b_0) + b_1 D_1 + b_2 D_2 + b_{3(1)} \log X_1 + b_{4(1)} \log X_2 + \dots + b_{n(1)} \log X_n + b_{3(2)} D_1(\log X_1) + b_{4(2)} D_1 \\ & (\log X_2) + \dots + b_{n(2)} D_1(\log X_n) + b_{3(3)} D_2(\log X_1) + b_{4(3)} D_2(\log X_2) + \dots + b_{n(3)} D_2(\log X_n) \end{aligned} \quad (10)$$

Where

Y = Average ion yield,

D_1 and D_2 = rock types (since there were 3 bedrock types, only two dummy variables need to be entered, the other is implied to be the control one).

and

X_1 , etc. = remaining basin characteristics.

In equation (10), if the b_1 and or b_2 coefficients are significant, they are due to the effects of bedrock types on the ion yields. If any slope coefficient for terms with bedrock dummy variables (e.g. $b_i D_j$) is significant, the corresponding independent variable is interpreted to be different between/among the bedrock types. Otherwise, it is not.

10. Stream water ion yield prediction, model validation, and spatial distribution of the predicted ion yields.

The models of principal component regression and the traditional multiple linear regression were compared to each other. Then, the most significant regression equation models were selected for ion yield prediction. The ArcGIS-based models were used to display and classify the areas of different affects on stream water ion yields.

The predicted values are compared with actual data of the average control observed constituent concentrations to determine which model might better represent prediction for the dependent responses. The model verification and validation procedures are based on the following criteria:

- High regression correlation ($R^2 \geq 0.64$)
- No or low collinearity
- Low Average Percent Error (APE) for each model:

$$\text{APE (\%)} = (\text{actual value} - \text{predicted value}) / \text{actual value} \quad (11)$$

In general, the percentage in equation (11) is expected to be as small as possible (e.g. less than or equal to 5%). That was also known as the “acid test”. However, depending on each application and situation, higher APE may be also accepted.

Due to the lack of data, only 4 sandstone stations were available for the prediction and validation of the models.

IV. RESULTS AND DISCUSSION

1. Descriptions of the project watersheds

1. 1. Watershed delineation and corresponding areas:

Watershed delineations were implemented based on the National Elevation Dataset, ArcGIS and the Hydrology Modeling Extension. The delineations of 47 watersheds can be found in Appendix 8. Figure 3 shows an example of a sandstone watershed delineated from National Elevation Dataset.

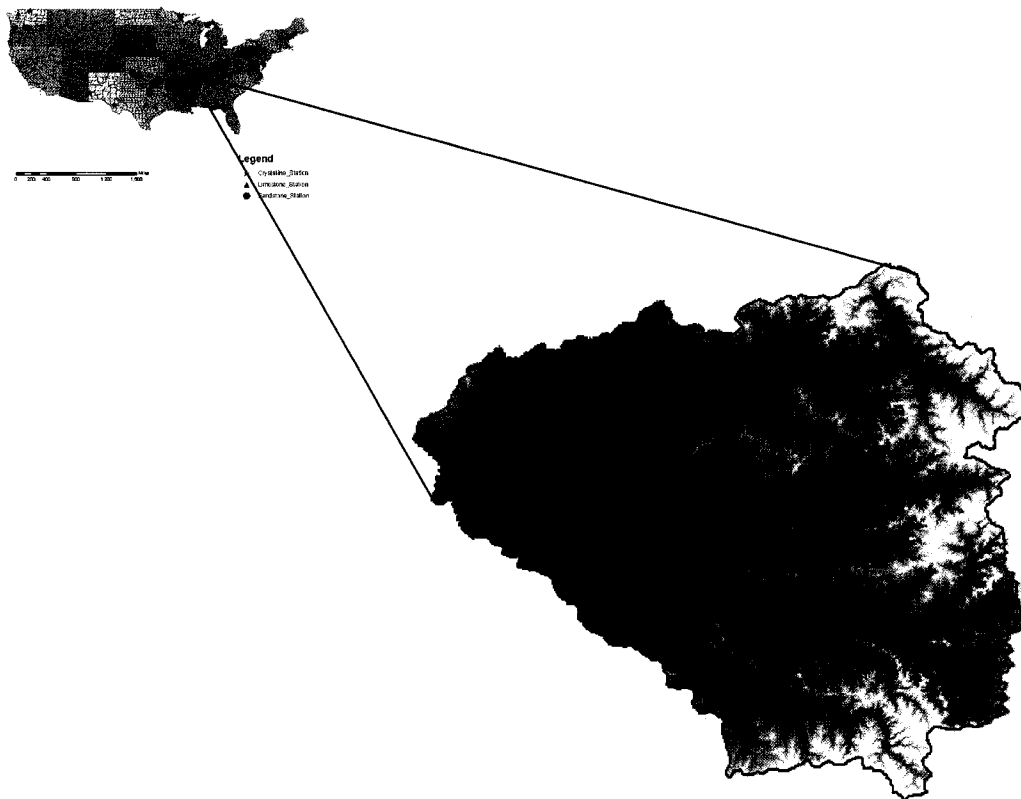


Figure 3. A sandstone watershed delineated from National Elevation Dataset.

It can be noted that the watershed shapes are very diverse and irregular (Appendix 8, Figure 3). The areas (SQMI) of those watersheds are also variable (Table 6 and Table 7a). They range from small (10.8 SQMI, limestone) to large (47223.1 SQMI, sandstone) (Table 7a). Furthermore, the area differences between delineated and USGS watersheds are very small although most delineated watersheds are slightly smaller than those of USGS data (indicated by negative sign, Table 7a). The mean area difference of sandstone bedrock is 1.4%, of limestone bedrock is 1.9%, and of crystalline is 3.5% (Table 6). Those numbers indicate that almost delineated watersheds were acceptable when compared to those of USGS (the differences < 5 %). It should also be noted that, however, there were some watersheds with a larger difference (e.g. 18.6% (USGS 08103900, limestone), 13.3% (USGS 04014500, crystalline)) (Table 6 and Appendix 3).

Table 6. Descriptive statistics of the delineated watershed areas and the area differences in Table 7a (in square mile, SQMI).

	n	Minimum	Maximum	Mean	Std. Deviation
Limestone_Delineated (SQMI)	19	10.8	32784.1	6314.9	8889.0
Limestone_Difference (%)	19	.0	18.6	1.9	4.1
Sandstone_Delineated (SQMI)	11	12.1	47223.1	13241.4	14085.9
Sandstone_Difference (%)	11	.2	3.0	1.4	.9
Crystalline_Delineated(SQMI)	17	22.2	10004.1	1473.6	2925.4
Crystalline_Difference (%)	17	.1	13.3	3.5	4.0

Table 7a. Differences between the delineated and USGS areas of the project watersheds.

Area (SQMI)			
	Delineated	USGS	Difference (%)
Basin Station	Limestone Bedrock		
USGS 02313000	1804.5	1820.0	-0.9
USGS 02320500	7698.4	7880.0	-2.3
USGS 02329000	1145.9	1140.0	0.5
USGS 02368000	628.8	624.0	0.8
USGS 03085000	7332.3	7337.0	-0.1
USGS 03254000	3371.3	3385.0	-0.4
USGS 03290500	6168.3	6180.0	-0.2
USGS 03425000	10690.2	10690.0	0.0
USGS 03543005	17266.2	17310.0	-0.3
USGS 03571850	22627.5	22640.0	-0.1
USGS 03593005	32784.1	32820.0	-0.1
USGS 04045500	771.8	790.0	-2.3
USGS 04057004	1461.5	1445.0	1.1
USGS 04108690	2023.0	2020.0	0.1
USGS 04126520	1902.6	1928.0	-1.3
USGS 04132052	1537.6	1500.0	2.5
USGS 04165500	718.5	734.0	-2.1
USGS 08103900	39.5	33.3	18.6
USGS 13018300	10.8	10.6	1.9
	Sandstone Bedrock		
USGS 01545600	44.8	46.2	-3.0
USGS 03049625	11445.1	11546.0	-0.9
USGS 03237280	12.1	12.2	-0.9
USGS 06452000	9913.5	10200.0	-2.8
USGS 06478500	20973.7	20653.0	1.6
USGS 07026000	1848.7	1852.0	-0.2
USGS 07231500	27609.2	27952.0	-1.2
USGS 07234000	7826.6	7955.0	-1.6
USGS 07237500	11696.9	11589.0	0.9
USGS 07245000	47223.1	47576.0	-0.7
USGS 06438000	7061.8	7210.0	-2.1
	Crystalline Bedrock		
USGS 01059400	3462.5	3434.0	0.8
USGS 01066000	1287.5	1293.0	-0.4
USGS 04014500	121.4	140.0	-13.3
USGS 05124480	237.2	254.0	-6.6
USGS 06623800	72.2	72.7	-0.7
USGS 07311200	22.2	24.6	-9.9
USGS 08431700	52.0	52.4	-0.8
USGS 09352900	78.2	72.5	7.8
USGS 09430600	73.4	69.0	6.4
USGS 09508300	36.5	36.4	0.1
USGS 11264500	179.9	181.0	-0.6
USGS 12447390	22.6	22.1	2.3
USGS 13169500	244.2	253.0	-3.5
USGS 13331500	237.1	240.0	-1.2
USGS 14048000	7629.7	7580.0	0.7
USGS 14103000	10004.1	10500.0	-4.7
USGS 14113000	1291.3	1297.0	-0.4

To assure that the delineated watershed area and corresponding USGS area are not different, a nonparametric test was conducted. The results (Table 7b) indicate that they are not statistically different, the P-values range from 0.847 (sandstone watersheds) to 0.919 (crystalline watersheds) and to 0.977 (limestone watersheds) (Table 7b).

Table 7b. Nonparametric test for the differences between delineated watershed areas and USGS watershed areas.

	Area_Limestone	Area_Sandstone	Area_Crystalline
Mann-Whitney U	179.0	57.0	141.0
Wilcoxon W	369.0	123.0	294.0
Z	-.044	-.230	-.121
Asymp. Sig. (2-tailed)	.965	.818	.904
Exact Sig. [2*(1-tailed Sig.)]	.977	.847	.919

1. 2. Basin characteristics (environmental factors):

Land cover, population, and precipitation were clipped for each watershed separately using the delineated results from the National Elevation Dataset. These results are described in the following sections:

1. 2.1. Land cover:

The 21 land cover classes of the watersheds in each bedrock type and the recombined 4 land cover categories were presented in Appendix 3a_4.

The descriptive statistics of the 4 new categories of land cover of each watershed are displayed in Table 8. The land cover types range considerably from zero SQMI (e.g. *Water* of crystalline bedrock) to approximately 36,000 SQMI (*Agriculture* of sandstone bedrock).

Table 8. Descriptive statistics of the 4 land cover categories of the project watersheds (in square mile, SQMI).

	n	Minimum	Maximum	Mean	Std. Deviation
Limestone Bedrock					
Water	19	0	1923.6	369.3	471.8
Developed	19	.2	1050.8	231.9	310.4
Forest	19	8.9	23514.5	4383.5	6782.3
Agriculture	19	1.8	7159.9	1330.2	1750.0
Sandstone Bedrock					
Water	11	0	914.1	213.8	276.9
Developed	11	0	563.9	147.8	173.8
Forest	11	12.0	10206.0	2668.9	3909.3
Agriculture	11	.1	35934.6	10210.8	11125.0
Crystalline Bedrock					
Water	17	0	271.8	35.2	70.0
Developed	17	0	194.9	29.5	51.8
Forest	17	5.5	8707.1	1271.6	2555.7
Agriculture	17	.5	989.5	137.3	297.8

In general, average *Water* and *Developed* lands of each bedrock type occupied only very small percentages (1.0% to 13.6%) while *Forest* and *Agriculture* lands are major categories occupying up to 79.6% (crystalline) and 62.3% (sandstone), respectively (Table 9).

Table 9. Percentage of the 4 land cover categories of the project watersheds (%).

	<i>Water</i>	<i>Developed</i>	<i>Forest</i>	<i>Agriculture</i>
Limestone	13.6	4.8	57.0	24.7
Sandstone	1.8	1.0	34.9	62.3
Crystalline	4.2	3.5	79.6	12.7
Average	6.5	3.1	57.2	33.2

1. 2. 2. Population:

The average population density (people/SQMI) of the watersheds of each bedrock type was analyzed to get the descriptive statistics (Table 10).

Table 10. Average population of each watershed (people/SQMI, 1990).			
Site No.	Limestone	Sandstone	Crystalline
1	146.7	39.0	51.9
2	41.3	169.8	42.7
3	74.5	122.4	5.0
4	62.1	4.1	4.5
5	225.7	8.1	1.3
6	49.4	55.2	103.0
7	104.2	18.3	1.0
8	55.8	5.9	14.1
9	112.4	5.8	2.7
10	113.2	27.0	9.2
11	105.3	5.7	22.1
12	11.1		6.0
13	8.0		1.0
14	186.1		7.2
15	37.2		3.1
16	31.2		11.9
17	1285.6		25.4
18	22.0		
19	3.0		
Descriptive Statistics			
Minimum	3.0	4.1	1.0
Maximum	1285.6	169.8	103.0
Mean	140.8	41.9	18.4
Std. Deviation	283.9	55.0	26.3

Average population density ranges from 1.0 people/SQMI (crystalline) to 1285.6 people/SQMI (limestone). The mean population density of each bedrock type also varied considerably from 18.4 (crystalline) to 140.8 (limestone). Limestone bedrock had the largest mean, however, the standard deviation was also highest (283.9) (Table 10).

1. 2. 3. Precipitation:

The lowest average annual precipitation (in/yr) was found in crystalline bedrock (11.5 in) while the highest amount belonged to limestone bedrock (60.3 in) (Table 11). Sandstone bedrock had the lowest mean annual precipitation (28.4 in) and limestone bedrock had the highest (43.7 in). Also, the standard (std.) deviation of each bedrock type was not very different from each other (from 10.2 to 13.7 in).

Site No.	Limestone	Sandstone	Crystalline
1	51.3	41.8	45.7
2	51.6	43.9	50.1
3	53.2	41.0	29.0
4	60.3	17.8	29.0
5	46.2	19.4	46.5
6	45.0	53.0	29.8
7	47.0	19.6	20.3
8	53.3	17.5	40.6
9	52.4	18.9	26.7
10	53.9	22.3	22.0
11	54.7	17.1	43.4
12	33.0		34.2
13	32.5		11.5
14	35.8		46.4
15	32.5		18.4
16	31.7		20.5
17	31.1		34.6
18	31.0		
19	34.6		
Descriptive Statistics			
Minimum	31.0	17.1	11.5
Maximum	60.3	53.0	50.1
Mean	43.7	28.4	32.3
Std. Deviation	10.2	13.5	11.7

1. 2. 4. Average slope:

The average slope (degree) of each watershed was low, ranging from 0.7 degrees (limestone) to 27.5 degrees (crystalline) (Table 12). The means of the average slope of

each bedrock type were also low and very close to each other, ranging from 5.2 (sandstone) to 6.4 (limestone) and to 13.2 (crystalline). However, the standard deviations of the mean slope of each bedrock type was large, ranging from 5.1 (sandstone) to 6.2 (limestone) and to 8.1 (crystalline) (Table 12). This suggests that slope may not affect the average annual stream ion yields.

Table 12. Average slope of each watershed (degree).

Site No.	Limestone	Sandstone	Crystalline
1	0.7	12.3	7.7
2	0.8	7.0	8.5
3	1.7	17.2	2.7
4	2.7	4.0	3.0
5	11.1	1.0	12.6
6	9.6	2.5	6.2
7	13.3	3.8	14.0
8	11.3	1.0	26.8
9	13.6	1.2	27.5
10	12.7	2.9	20.0
11	10.4	4.8	17.8
12	1.1		21.2
13	0.9		5.7
14	1.6		22.9
15	2.6		12.8
16	2.3		6.8
17	1.1		8.5
18	2.8		
19	21.7		
Descriptive Statistics			
Minimum	.7	1.0	2.7
Maximum	21.7	17.2	27.5
Mean	6.42	5.24	13.23
Std. Deviation	6.23	5.14	8.10

1. 3. Stream water ion yields (kg/ha/yr):

Stream water ion yields were calculated (Appendix 3). The average annual ion yields (kg/ha/yr) varied dramatically among watersheds and among bedrock types (Table 13a). In general, total dissolved solids (TDS) was the highest for the three bedrock types, with TDS of limestone (722.6), sandstone (301.3), crystalline (230.3). These differences were significantly different ($p < 0.05$, Table 16). The lowest values were for orthophosphate and phosphorus, being just less than 1.0 for all three bedrocks. The next lowest values were for total nitrogen, ranging from 2.5 (sandstone) to 4.3 (limestone). Potassium ranged from 4.6 (sandstone) to 7.0 (limestone); magnesium from 9 (crystalline) to 36 (limestone); sodium from 21.4 (crystalline) to 43.6 (limestone); dissolved oxygen from 28.3 (sandstone) to 50.4 (crystalline); calcium from 40.2 (crystalline) to 144.0 (limestone); chloride from 16.0 (crystalline) to 61.3 (limestone); sulfate from 30.8 (crystalline) to 123.0 (limestone); bicarbonate from 89.5 (sandstone) to 441.3 (limestone). These results are similar to the earlier study (Peters 1984).

The values of pH among different bedrock types were very close to each other, ranging from 7.5 crystalline to 7.7 (limestone, and sandstone).

In general, when compared to Peters 1984, most average constituent yields of the current research appear higher, except for potassium, magnesium, calcium, and sulfate in limestone watersheds (Table 13b and Appendix 7) although they are not statistically different ($p > 0.05$, Table 13c). The slight differences are partly due to the fact that the delineated watershed areas are a little smaller than those of USGS data as shown in Table 7a.

**Table 13a. Descriptive statistics of average annual stream water ion yields
of each bedrock type (kg/ha/yr).**

Water constituent	<i>n</i>	<i>Minimum</i>	<i>Maximum</i>	<i>Mean</i>	<i>Std. Deviation</i>
Limestone Bedrock					
TDS	19	199.6	1415.5	722.6	384.4
DO	19	8.9	64.0	40.6	15.6
pH	19	7.0	8.4	7.7	0.4
Total Nitrogen	16	0.6	11.7	4.3	3.2
Orthophosphate	19	0.1	2.6	0.7	0.7
Phosphorus	19	0.0	0.7	0.2	0.2
Ca	19	17.3	270.4	144.0	72.8
Mg	19	6.7	71.9	36.0	20.1
Na	19	7.1	178.0	43.6	43.4
K	19	0.9	17.7	7.0	4.9
Chloride	19	4.2	305.1	61.3	71.6
Sulfate	19	14.8	618.7	123.0	143.8
Bicarbonate	14	87.1	830.0	441.3	255.1
Sandstone Bedrock					
TDS	11	55.4	1352.6	301.3	355.8
DO	9	0.5	98.2	28.3	35.0
pH	11	6.7	8.3	7.7	0.6
Total Nitrogen	9	0.1	8.9	2.5	3.1
Orthophosphate	11	0.0	0.9	0.2	0.3
Phosphorus	9	0.0	0.4	0.1	0.1
Ca	11	6.7	218.5	42.0	59.5
Mg	11	2.0	58.8	15.0	16.0
Na	11	5.8	102.8	34.8	27.0
K	11	0.3	15.3	4.6	4.6
Chloride	11	4.1	132.3	37.1	42.2
Sulfate	11	16.8	563.9	119.8	160.0
Bicarbonate	10	11.7	326.8	89.5	92.6
Crystalline Bedrock					
TDS	17	28.4	1397.3	230.3	308.5
DO	17	3.6	171.4	50.4	42.4
pH	17	6.6	8.3	7.5	0.5
Total Nitrogen	17	0.2	19.7	3.3	4.5
Orthophosphate	16	0.0	13.5	1.0	3.3
Phosphorus	17	0.0	3.2	0.3	0.8
Ca	17	4.0	271.6	40.2	61.1
Mg	17	1.0	39.9	9.0	8.9
Na	17	3.4	105.5	21.4	24.9
K	17	0.9	57.3	6.6	13.2
Chloride	17	1.8	92.4	16.0	25.1
Sulfate	17	0.3	197.4	30.8	45.8
Bicarbonate	14	37.8	247.3	128.5	62.8

The area differences are not statistically different (Table 7b). Nevertheless, these results suggest that there may be a trend of increasing water constituent yields from the period of 1970-1980 to 1980-1990 in the project watersheds in the United States.

Table 13b. Average annual ion yields (kg/ha/yr) in watersheds of each bedrock type from Peters 1984 and the current study (compiled for the corresponding watersheds only).								
	Na	K	Mg	Ca	Chloride	Sulfate	Bicarbonate	Dissolved Solids
Limestone								
This study	43.6	6.99	36.0	144.0	61.3	123.0	441.3	722.6
Peters 1984	41.1	7.01	36.1	145.6	59.1	123.5	439.1	894.5
Changes (%)	6.0	-0.3	-0.1	-1.1	3.7	-0.4	0.5	-19.2
Sandstone								
This study	34.7	4.9	14.4	42.1	40.2	109.3	96.0	306.3
Peters 1984	19.7	4.3	12.9	30.1	23.6	76.9	61.0	227.6
Changes (%)	76.4	14.6	11.9	40.2	70.7	42.2	57.3	34.6
Crystalline								
This study	21.4	6.6	9.0	40.2	16.0	30.8	128.5	230.3
Peters 1984	15.9	3.2	6.8	26.1	11.8	21.2	122.7	198.4
Changes (%)	34.4	102.8	32.6	54.2	35.6	45.4	4.8	16.1

Table 13c. Nonparametric test for the differences between stream water ion yields of Peters 1984 and the current study.

	Limestone							
	Na	K	Mg	Ca	Chloride	Sulfate	Bicarbonate	Dissolved Solids
Mann-Whitney U	158.50	175.0	169.0	178.5	160.5	167.0	94.0	148.0
Wilcoxon W	348.50	365.0	359.0	368.5	350.5	357.0	214.0	338.0
Z	-.642	-.161	-.336	-.058	-.584	-.394	-.480	-.949
Asymp. Sig. (2-tailed)	.521	.872	.737	.953	.559	.693	.631	.343
Exact Sig. [2*(1-tailed Sig.)]	.525	.885	.751	.954	.563	.708	.652	.354
Sandstone								
Mann-Whitney U	28.5	39.0	37.5	32.0	39.0	26.0	29.0	34.0
Wilcoxon W	83.5	94.0	92.5	87.0	94.0	81.0	74.0	89.0
Z	-1.626	-.832	-.945	-1.361	-.832	-1.814	-1.015	-1.209
Asymp. Sig. (2-tailed)	.104	.406	.345	.173	.406	.070	.310	.226
Exact Sig. [2*(1-tailed Sig.)]	.105	.436	.353	.190	.436	.075	.340	.247
Crystalline								
Mann-Whitney U	119.0	125.0	128.0	131.50	137.0	130.0	91.5	118.5
Wilcoxon W	272.0	278.0	281.0	284.5	290.0	283.0	196.5	271.5
Z	-.878	-.672	-.568	-.448	-.259	-.499	-.299	-.896
Asymp. Sig. (2-tailed)	.380	.501	.570	.654	.796	.617	.765	.370
Exact Sig. [2*(1-tailed Sig.)]	.394	.518	.586	.658	.812	.634	.769	.375

The average annual yield of sulfate in sandstone watersheds and of potassium in crystalline watersheds are good examples of demonstrating the increasing trend of some stream water ion yields in the project watersheds (Figure 4 and Figure 5). These increases might be due to the effects of human and industrial activities such as waste, fertilizers, smoke, and water treatment plants.

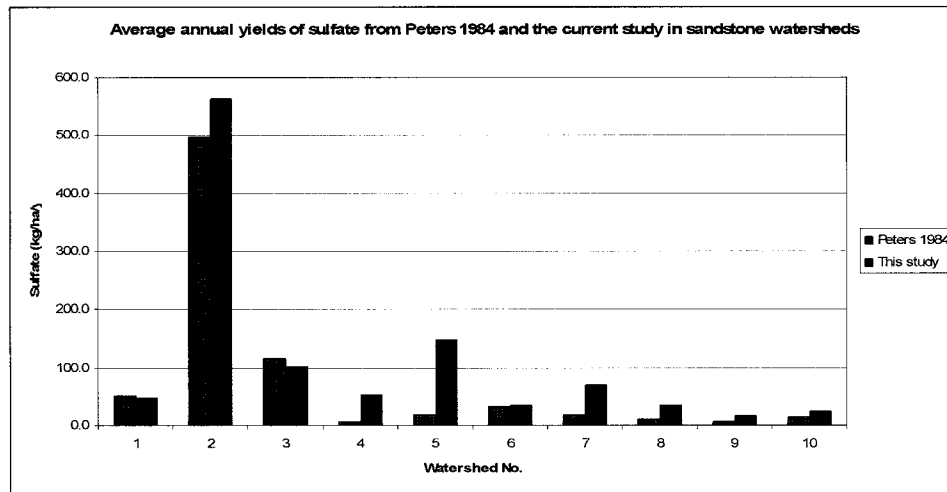


Figure 4. Comparison of average annual sulfate yields (kg/ha/yr) between Peters 1984 and the current research in limestone watersheds.

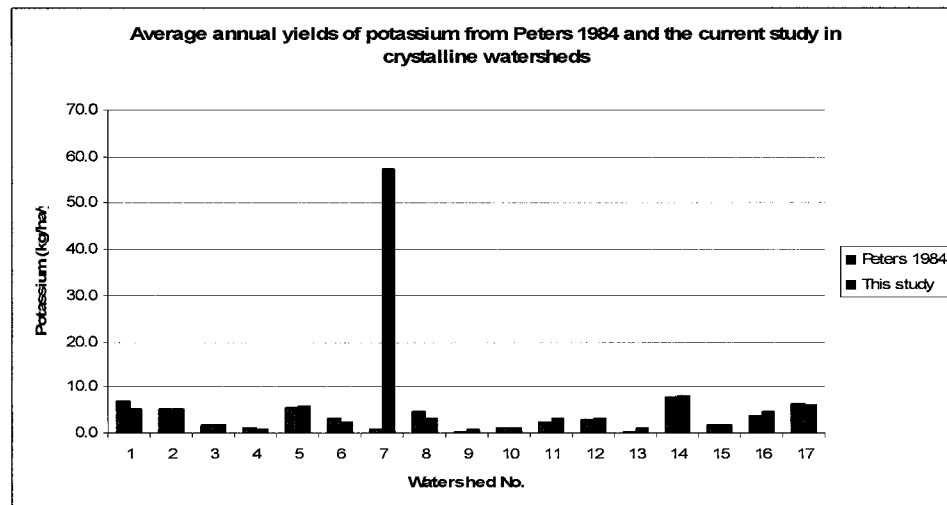


Figure 5. Comparison of average annual potassium yields (kg/ha/yr) between Peters 1984 and the current research in crystalline watersheds.

1. 4. Kolomogorov-Smirnov Tests:

The One-Sample Kolmogorov-Smirnov (KS) test is used to test the null hypothesis that a sample comes from a particular distribution (e.g. normal, uniform, Poisson, or exponential). In this study, the null hypothesis was that a sample comes from a normal distribution. Since the sample size (the number of watersheds) of each bedrock type was small, the KS test was conducted for the combined data set of 47 watersheds.

For the stream water ion yield, the results (Table 14) indicated that almost all stream water ion yields were non-normal, except for pH ($p = 0.758$) and DO ($p = 0.509$). Magnesium, sodium, and bicarbonate were almost close to normality (p approximately equals to 0.05). However, since the data size of 47 was still small, the normality test might be biased.

Table 14. Kolomogorov-Smirnov Test for the stream ion yields combined from the three bedrocks (limestone, sandstone, and crystalline).

		TDS	DO	pH	Total_ Nitrogen	Ortho- phosphate	Phos- phorus	Ca	Mg	Na	K	Cl	SO ₄	Bicar- bonate
N		47	45	47	42	46	45	47	47	47	47	47	47	38
Normal Parameters (a,b)	Mean	446.0	41.8	7.7	3.5	.7	.2	82.6	21.3	33.5	6.3	39.2	88.9	233.4
	Std. Deviation	415.0	32.3	.5	3.7	2.0	.5	82.2	19.9	34.8	8.7	54.9	127.9	229.0
Most Extreme Differences	Absolute	.25	.122	.098	.213	.368	.362	.257	.187	.193	.247	.248	.259	.216
	Positive	.25	.122	.069	.213	.347	.362	.257	.187	.185	.234	.224	.259	.216
	Negative	-.16	-.100	-.098	-.181	-.368	-.343	-.169	-.154	-.193	-.247	-.248	-.244	-.166
Kolmogorov-Smirnov Z		1.72	.82	.67	1.38	2.50	2.43	1.76	1.28	1.33	1.69	1.70	1.78	1.33
Asymp. Sig. (2-tailed)		.006	.509	.758	.044	.000	.000	.004	.074	.060	.007	.006	.004	.058

a Test distribution is Normal.
b Calculated from data.

For the watershed characteristics normality test (Table 15), only *Precipitation* and *Slope* were normal ($p = 0.651$ and $p = 0.212$, respectively). All other factors were significant statistically, meaning that they are not normal.

Table 15. Kolomogorov-Smirnov Test for the environmental factors combined from the three bedrocks (limestone, sandstone, and crystalline).

		Water	Developed	Forest	Agriculture	Precipitation	Population	Slope	Area	Bedrock
N		47	47	47	47	47	47	47	47	47
Normal Parameters (a,b)	Mean	212.1	139.0	2856.7	2977.2	36.0	73.4	8.6	6185.0	2.0
	Std. Deviation	356.7	230.7	5049.1	6689.4	13.1	188.8	7.5	9856.3	.9
Most Extreme Differences	Absolute	.276	.273	.301	.328	.107	.351	.155	.266	.265
	Positive	.233	.258	.301	.321	.107	.291	.155	.259	.265
	Negative	-.276	-.273	-.286	-.328	-.098	-.351	-.146	-.266	-.243
Kolmogorov-Smirnov Z		1.893	1.874	2.065	2.250	.736	2.405	1.059	1.820	1.817
Asymp. Sig. (2-tailed)		.002	.002	.000	.000	.651	.000	.212	.003	.003

a Test distribution is Normal.

2. Effects of bedrock types on stream ion yields

The descriptive statistics of average annual stream water ion yields of each bedrock type (kg/ha/yr) are displayed (Table 13a and Appendix 3). Stream water ion yields of limestone bedrock are quite different from those of the other two (Table 16), including total dissolved solids, orthophosphate, calcium, magnesium, and bicarbonate ($p < 0.05$), and consistent with Peters 1984. However, sandstone and crystalline ion yields are not much different, except for chloride ($p = 0.033$) and sulfate ($p = 0.002$). These results in combination with the results in Table 13a, suggest that yield of total dissolved solids of limestone watersheds is about 2.4 times greater than sandstone watersheds and 3.1 times greater than crystalline watersheds.

The ratio of ion yields in limestone watersheds to those in crystalline watersheds and in sandstone watersheds is about: bicarbonate, 5:1.4:1; orthophosphate, 3.5:5:1; calcium, 3.4:1:1, and magnesium, 2.4:0.6:1 (Table 13a). These ratio results are consistent with Peters 1984.

Chloride and sulfate concentrations from limestone and sandstone watersheds are significantly different from those of crystalline watersheds, and the yield of chloride from limestone bedrock is about 4 times greater than that from crystalline bedrock, and sandstone is about 1.4 times greater than crystalline bedrock. The yields of sulfate from limestone and sandstone bedrocks are 4 times greater than that for crystalline bedrock (Table 13a). This may be related to both climate and farming practices (Peters, 1984). When compared to the crystalline watersheds, the limestone watersheds generally have a higher population density and higher annual precipitation (Table 10 and Table 11), which consequently bring in larger quantities of sulfate either by atmospheric deposition or fertilizer application and other farming activities (Peters, 1984). However, the sulfate yields from limestone watersheds are not statistical different from those of sandstone watersheds like Peters 1984. That might be due to other factors affecting sulfate, such as, developed land, forest, and agriculture.

The difference test between bicarbonate concentrations from sandstone and crystalline watersheds resulted in a p-value of 0.064 (Table 16). Dissolved oxygen is not different between limestone and crystalline watersheds ($p = 0.684$) while there is just only slight difference between dissolved oxygen yields of limestone and sandstone ($p = 0.105$) or sandstone and crystalline watersheds ($p = 0.085$). Potassium is almost different

between limestone and sandstone ($p = 0.103$) or limestone and crystalline watersheds ($p = 0.061$). Sodium yields of limestone and crystalline or of sandstone and crystalline bedrocks are also almost different ($p = 0.057$ and 0.100 , respectively). The statistical significance of total nitrogen yields of limestone and sandstone watersheds is slightly larger than 0.05 ($p = 0.074$).

Table 16. Nonparametric Test Statistics for the ion yield differences between two bedrock types of the three bedrock studied.

	TDS	DO	pH	Total Nitrogen	PO ₄	Phosphorus	Ca	Mg	Na	K	Cl	Sulfate	Bicarbonate
Limestone vs. Sandstone													
Mann-Whitney U	28.0	52.0	88.5	40.0	39.0	45.0	23.0	32.0	97.0	66.0	73.0	91.0	9.0
Wilcoxon W	94.0	97.0	278.5	85.0	105.0	90.0	89.0	98.0	163.0	132.0	139.0	157.0	64.0
Z	-3.29	-1.648	-.691	-1.812	-2.819	-1.99	-3.507	-3.1	-.323	-1.657	-1.356	-.581	-3.572
Asymp. Sig. (2-tailed)	.001	.099	.490	.070	.005	.046	.000	.002	.747	.098	.175	.561	.000
Exact Sig. [2*(1-tailed Sig.)]	.001	.105	.497	.074	.004	.048	.000	.001	.767	.103	.185	.582	.000
Limestone vs. Crystalline													
Mann-Whitney U	26.0	148.0	128.0	102.0	67.0	106.0	34.0	33.0	101.0	102.0	52.0	45.0	20.0
Wilcoxon W	179.0	338.0	281.0	255.0	203.0	259.0	187.0	186.0	254.0	255.0	205.0	198.0	125.0
Z	-4.294	-.428	-1.065	-1.225	-2.815	-1.759	-4.04	-4.1	-1.917	-1.885	-3.470	-3.692	-3.584
Asymp. Sig. (2-tailed)	.000	.669	.287	.221	.005	.079	.000	.000	.055	.059	.001	.000	.000
Exact Sig. [2*(1-tailed Sig.)]	.000	.684	.300	.231	.004	.081	.000	.000	.057	.061	.000	.000	.000
Sandstone vs. Crystalline													
Mann-Whitney U	66.0	44.0	69.5	60.0	72.0	67.0	91.0	65.0	58.0	90.0	48.0	30.0	38.0
Wilcoxon W	219.0	89.0	222.5	105.0	138.0	112.0	157.0	218.0	211.0	156.0	201.0	183.0	93.0
Z	-1.294	-1.752	-1.132	-.889	-.790	-.512	-.118	-1.3	-1.670	-.165	-2.140	-2.987	-1.874
Asymp. Sig. (2-tailed)	.196	.080	.257	.374	.430	.609	.906	.180	.095	.869	.032	.003	.061
Exact Sig. [2*(1-tailed Sig.)]	.208	.085	.264	.396	.451	.634	.926	.191	.100	.890	.033	.002	.064

Phosphorus yield of sandstone is the same as that of crystalline, but is close to the difference level between limestone vs. crystalline ($p = 0.081$) and is significantly different between limestone vs. sandstone watersheds ($p = 0.048$). This indicates that the yield of phosphorus of limestone watersheds is 2 times greater than that of sandstone watersheds.

The pH values of the three project bedrocks showed no difference between bedrock type.

Ion yields for limestone and sandstone bedrocks are more closely related to the environmental factors than those in crystalline bedrock (Table 17). Most of the ion yields in crystalline watersheds have no relation with the environmental factors, except for DO, pH, and total nitrogen. In the three bedrock types, the present of significant environmental factors of the same ion yield (and the corresponding adjusted R square) is almost different from this bedrock compared to the other two. For example, in limestone bedrock, bicarbonate is related to precipitation, in sandstone, bicarbonate is related to population and water, and in crystalline, bicarbonate is not related to any factor (Table 17). Precipitation is significant in the relation with pH in all three bedrocks. It is also significant in the relation with dissolved oxygen and total nitrogen in sandstone and crystalline bedrocks (Table 17). In addition, the slope of a regression of an ion yield on any specific environmental factors (if significant) is the same of the three bedrocks. pH, for example, has a negative relation with precipitation for all three bedrocks (Table 17).

Table 17. Summary of the multiple linear regression results of the relationship between Log of ion yield and Log of the environmental factors in the three bedrocks.

Ion yield	Limestone	Sandstone	Crystalline
TDS	ND	Population(+), Water(+); (.530)	ND
DO	Slope(+); (.385)*	Precipitation(+); (.821)	Precipitation(+); (.447)
pH	Precipitation(-); (.611)	Precipitation(-); (.932)	Precipitation(-); (.379)
Total Nitrogen	ND	Precipitation(+); (.761)	Precipitation(+); (.290)
Orthophosphate	Developed(+), Water(-); (.388)	ND	ND
Phosphorus	Developed(+); (.261)	Precipitation(+); (.447)	ND
Ca	ND	ND	ND
Mg	Precipitation(-), Slope(+), Agriculture(+); (.834)	Population(+); (.307)	ND
Na	Population(+), Slope(+), Precipitation(-); (.740)	Water(+); (.392)	ND
K	Slope(+), Population(+); (.530)	Population(+); (.540)	ND
Chloride	Population(+); (.585)	ND	ND
Sulfate	Forest(+), Precipitation(-); (.354)	ND	ND
Bicarbonate	Precipitation(-); (.311)	Population(+), Water(+); (.636)	ND

Notations in Table 17:

ND: No regression coefficient detected to be significant

Precipitation(-): Significant negative relation (between the corresponding Log of ion yield and Log of precipitation)

Population(+): Significant positive relation (between the corresponding Log of ion yield and Log of population)

Significant level (Alpha) = 0.05

(.385): Adjusted R square of the regression

To know how bedrock types affect the relation between the environmental factors and the selected ion yields, it is also necessary to distinguish the slope coefficient of the same predictor (environmental factor) for the same ion yield regression among bedrock types. As mentioned earlier, precipitation proved to be significant in relation to pH in all three bedrocks, and with dissolved oxygen and total nitrogen in sandstone and crystalline bedrocks.

To test for effects of bedrocks on significant coefficient slope of precipitation on pH in the three bedrock types, the data were recompiled (Table 18) and regression analysis prepared (Table 19).

Table 18. Data for the test of effects of bedrocks on significant coefficient slope of Precipitation on pH of the three bedrock types (except for Sandstone, all other data are in log-transformed)

Site No.	Station No.	pH	Precipitation	Limestone	Sandstone	Limestone x Precipitation	Sandstone x Precipitation
1	02313000	0.88	1.71	1	0	1.71	0
2	02320500	0.89	1.71	1	0	1.71	0
3	02329000	0.85	1.73	1	0	1.73	0
4	02368000	0.85	1.78	1	0	1.78	0
5	03085000	0.88	1.66	1	0	1.66	0
6	03254000	0.89	1.65	1	0	1.65	0
7	03290500	0.89	1.67	1	0	1.67	0
8	03425000	0.88	1.73	1	0	1.73	0
9	03543005	0.88	1.72	1	0	1.72	0
10	03571850	0.88	1.73	1	0	1.73	0
11	03593005	0.87	1.74	1	0	1.74	0
12	04045500	0.89	1.52	1	0	1.52	0
13	04057004	0.89	1.51	1	0	1.51	0
14	04108690	0.91	1.55	1	0	1.55	0
15	04126520	0.9	1.51	1	0	1.51	0
16	04132052	0.91	1.5	1	0	1.5	0
17	04165500	0.91	1.49	1	0	1.49	0
18	08103900	0.9	1.49	1	0	1.49	0
19	13018300	0.92	1.54	1	0	1.54	0
20	01545600	0.84	1.62	0	1	0	1.62
21	03049625	0.87	1.64	0	1	0	1.64
22	03237280	0.85	1.61	0	1	0	1.61
23	06452000	0.91	1.25	0	1	0	1.25
24	06478500	0.91	1.29	0	1	0	1.29
25	07026000	0.83	1.72	0	1	0	1.72
26	07231500	0.91	1.29	0	1	0	1.29
27	07234000	0.92	1.24	0	1	0	1.24
28	07237500	0.91	1.28	0	1	0	1.28
29	07245000	0.89	1.35	0	1	0	1.35
30	06438000	0.92	1.23	0	1	0	1.23
31	01059400	0.85	1.66	0	0	0	0
32	01066000	0.82	1.7	0	0	0	0
33	04014500	0.88	1.46	0	0	0	0
34	05124480	0.84	1.46	0	0	0	0
35	06623800	0.89	1.67	0	0	0	0
36	07311200	0.86	1.47	0	0	0	0
37	08431700	0.86	1.31	0	0	0	0
38	09352900	0.89	1.61	0	0	0	0
39	09430600	0.89	1.43	0	0	0	0
40	09508300	0.9	1.34	0	0	0	0
41	11264500	0.83	1.64	0	0	0	0
42	12447390	0.88	1.53	0	0	0	0
43	13169500	0.92	1.06	0	0	0	0
44	13331500	0.88	1.67	0	0	0	0
45	14048000	0.92	1.27	0	0	0	0
46	14103000	0.91	1.31	0	0	0	0
47	14113000	0.89	1.54	0	0	0	0

Table 19. Model summary and the analysis of coefficients of the relationship between pH and precipitation in the combined data of the three bedrock types.

Model Summary						
Model	R	R Square		Adjusted R Square	Std. Error of the Estimate	
1	.631(a)	.398		.385	.022	
2	.799(b)	.638		.622	.017	
a Predictors: (Constant), Precipitation						
b Predictors: (Constant), Precipitation, Limestone						
Coefficients(a)						
Model		Unstandardized Coefficients		Standardized Coefficients	t	Sig.
		B	Std. Error	Beta		
1	(Constant)	1.033	.028		37.395	.000
	Precipitation	-.098	.018	-.631	-5.454	.000
2	(Constant)	1.086	.024		45.623	.000
	Precipitation	-.142	.016	-.909	-8.715	.000
	Limestone	.031	.006	.564	5.401	.000
a Dependent Variable: pH						

Excluded Variables(c)								
Model		Beta In	t	Sig.	Partial Correlation	Collinearity Statistics		
						Tolerance	VIF	Minimum Tolerance
1	Limestone	.564(a)	5.401	.000	.631	.756	1.323	.756
	Sandstone	-.149(a)	-1.210	.233	-.179	.877	1.140	.877
	Limestone_Precipitation	.569(a)	5.314	.000	.625	.727	1.375	.727
	Sandstone_Precipitation	-.180(a)	-1.516	.137	-.223	.926	1.080	.926
2	Sandstone	.037(b)	.356	.724	.054	.772	1.296	.665
	Limestone_Precipitation	-.473(b)	-.377	.708	-.057	.005	187.435	.005
	Sandstone_Precipitation	.015(b)	.143	.887	.022	.794	1.260	.648
a Dependent Variable: pH								
b Predictors in the Model: (Constant), Precipitation, Limestone								
c Dependent Variable: pH								

As shown in Table 19, *constant*, and *coefficients of Precipitation and Limestone* are significant while *Sandstone, Limestone x Precipitation, and Sandstone x Precipitation* are not significant ($p > 0.05$). Thus, they are not presented in the output of the “stepwise” model selection). This means that only pH in limestone is different from crystalline.

However, the coefficient of limestone, which is small (0.031), can be considered negligible. Since the slopes of the product of *Limestone x Precipitation* and of *Sandstone x Precipitation* (as shown in Table 18) are not significant, no effect of bedrocks on the predictor (precipitation) for pH is found.

Second, to test for effects of bedrocks on significant coefficient slope of precipitation on dissolved solids, pH, and on total nitrogen in between sandstone and crystalline bedrock types, the data were recompiled as shown in Table 20. The results of this regression analysis are expressed in Table 21, 23, and 24, for dissolved solids, pH, and nitrogen, respectively.

Table 20. Data for the test of effects of bedrocks on significant coefficient slopes of precipitation on DO, pH, and on total nitrogen in sandstone and crystalline bedrock types (except for Sandstone, all other data are in log-transformed).

Site No.	Station No.	pH	DO	total nitrogen	Precipitation	Sandstone	Sandstone x Precipitation
1	01545600	0.84	1.81	0.51	1.62	1.00	1.62
2	03049625	0.87	1.99	0.95	1.64	1.00	1.64
3	03237280	0.85	1.54	0.31	1.61	1.00	1.61
4	06452000	0.91	0.6	-0.07	1.25	1.00	1.25
5	06478500	0.91	0.4	-0.51	1.29	1.00	1.29
6	07026000	0.83	1.66	0.78	1.72	1.00	1.72
7	07231500	0.91	0.6	-0.24	1.29	1.00	1.29
8	07234000	0.92	Nodata	Nodata	1.24	1.00	1.24
9	07237500	0.91	-0.29	-0.88	1.28	1.00	1.28
10	07245000	0.89	Nodata	Nodata	1.35	1.00	1.35
11	06438000	0.92	0.27	-0.31	1.23	1.00	1.23
12	01059400	0.85	1.68	0.56	1.66	.00	.00
13	01066000	0.82	1.98	0.46	1.7	.00	.00
14	04014500	0.88	1.7	0.42	1.46	.00	.00
15	05124480	0.84	1.49	0.12	1.46	.00	.00
16	06623800	0.89	1.8	0.55	1.67	.00	.00
17	07311200	0.86	1.14	0	1.47	.00	.00
18	08431700	0.86	2.23	1.29	1.31	.00	.00
19	09352900	0.89	1.68	0.69	1.61	.00	.00
20	09430600	0.89	1	-0.36	1.43	.00	.00
21	09508300	0.9	0.93	-0.37	1.34	.00	.00
22	11264500	0.83	1.9	0.52	1.64	.00	.00
23	12447390	0.88	1.82	0.65	1.53	.00	.00
24	13169500	0.92	0.56	-0.63	1.06	.00	.00
25	13331500	0.88	1.93	0.53	1.67	.00	.00
26	14048000	0.92	1.01	-0.26	1.27	.00	.00
27	14103000	0.91	1.36	0.04	1.31	.00	.00
28	14113000	0.89	1.69	0.52	1.54	.00	.00

Table 21. Model summary and the analysis of coefficients of the relationship between DO and precipitation in the combined data of sandstone and crystalline bedrock types.

Model Summary						
Model	R	R Square		Adjusted R Square	Std. Error of the Estimate	
1	.754(a)	.568		.550	.433	
2	.832(b)	.693		.666	.373	
3	.871(c)	.759		.726	.338	
a Predictors: (Constant), Precipitation						
b Predictors: (Constant), Precipitation, Sandstone						
c Predictors: (Constant), Precipitation, Sandstone, Sandstone_Precipitation						
Coefficients(a)						
Model		Unstandardized Coefficients		Standardized Coefficients	t	Sig.
		B	Std. Error	Beta		
1	(Constant)	-2.558	.696		-3.673	.001
	Precipitation	2.652	.472	.754	5.619	.000
2	(Constant)	-2.201	.611		-3.601	.002
	Precipitation	2.521	.409	.716	6.163	.000
	Sandstone	-.473	.155	-.355	-3.055	.006
3	(Constant)	-1.105	.712		-1.551	.135
	Precipitation	1.779	.479	.506	3.715	.001
	Sandstone	-3.166	1.109	-.2378	-2.855	.009
	Sandstone_Precipitation	1.852	.757	2.028	2.448	.023
a Dependent Variable: DO						

In Table 21 (model 3), *constant*, *Precipitation*, *Sandstone*, and the product of *Sandstone x Precipitation* are significant. This indicates that both the log-transformed value of DO in sandstone and the coefficient slope of precipitation of sandstone are different from those in crystalline bedrock. The differences between DO in sandstone and crystalline bedrocks are not significant (Table 16). That is contrary with the result in Table 21. The reason for that may be due to the collinearity of *sandstone* and *precipitation* with the product of *sandstone x precipitation*.

Table 22. Model summary and the analysis of coefficients of the relationship between pH and precipitation in the combined data of sandstone and crystalline bedrock types.

Model Summary						
Model	R	R Square	Adjusted R Square	Std. Error of the Estimate		
1	.797(a)	.635	.621	.019		
a Predictors: (Constant), Precipitation						
Coefficients(a)						
Model		Unstandardized Coefficients		Standardized Coefficients	t	Sig.
		B	Std. Error	Beta		
1	(Constant)	1.082	.030		35.849	.000
	Precipitation	-.139	.021	-.797	-6.723	.000
a Dependent Variable: pH						

The results (Table 22 and Table 23) indicate that there are no differences of pH and total nitrogen and the significant slope coefficients between sandstone and crystalline bedrocks because the slopes of *Sandstone* and *Sandstone x Precipitation* are not significant ($p > 0.05$) (therefore *Sandstone* and *Sandstone x Precipitation* are not presented in Table 22 and Table 23).

Table 23. Model summary and the analysis of coefficients of the relationship between total nitrogen and precipitation in the combined data of sandstone and crystalline bedrock types.

Model Summary						
Model	R	R Square		Adjusted R Square	Std. Error of the Estimate	
1	.712(a)	.506		.486	.383	
a Predictors: (Constant), Precipitation						
Coefficients(a)						
Model		Unstandardized Coefficients		Standardized Coefficients	t	Sig.
		B	Std. Error	Beta		
1	(Constant)	-2.826	.615		-4.592	.000
	Precipitation	2.069	.417	.712	4.961	.000
a Dependent Variable: Total_Nitrogen						

The effects of bedrock types on stream water ion yields and their relationship with other environmental factors are discussed in later sections.

3. Effects of the other environmental factors on stream water ion yields

3.1. Effects of land cover:

The 21 class National Land Cover Dataset is grouped into 4 new categories, including *Water*, *Developed*, *Forest*, and *Agriculture*. Therefore, the effects of land cover are analyzed in effects of the 4 categories separately.

3.1.1. Effects of Water:

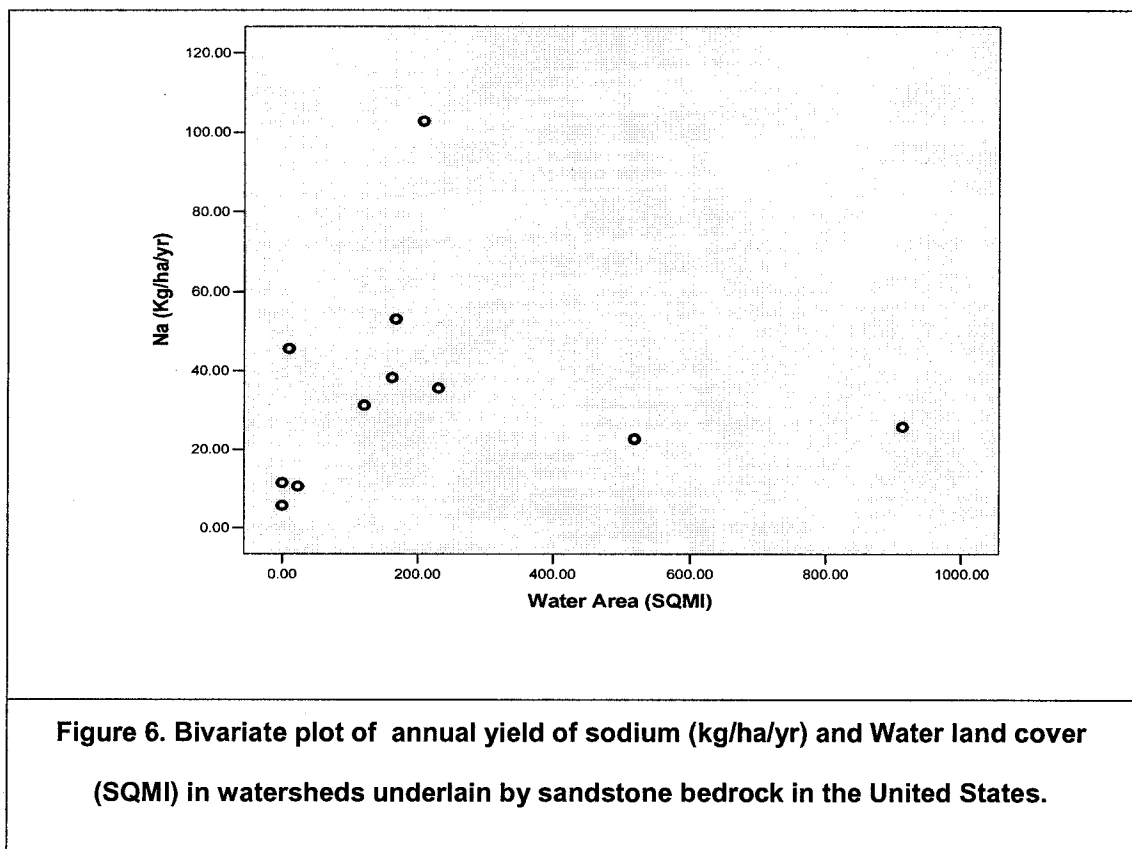
As listed in Table 9, *Water* of each bedrock type occupied only very small percentages (1.8% in sandstone to 13.6% in limestone), averaging at 6.5 %. However, there are still some slope coefficients of this land cover type significant in regressions of orthophosphate and magnesium in limestone, in regression of total dissolved solids, sodium, and bicarbonate in sandstone bedrock. In more details, the equations for these relations are found in Table 24:

Table 24. The presence of significant slope coefficients of *Water* in the regression equations in limestone and sandstone bedrocks ($p < 0.05$).

	R^2	Adjusted R^2
Limestone		
Orthophosphate = $10^{(-0.696)} \times (\text{Developed})^{(0.417)} \times (\text{Water})^{(-0.20)}$ (12)	0.456	0.388
Sandstone		
TDS = $10^{(1.59)} \times (\text{Population})^{(0.489)} \times (\text{Water})^{(0.091)}$ (13)	0.624	0.53
Na = $10^{(1.272)} \times (\text{Water})^{(0.118)}$ (14)	0.453	0.392
Bicarbonate = $10^{(0.777)} \times (\text{Population})^{(0.655)} \times (\text{Water})^{(0.129)}$ (15)	0.717	0.636

As shown in Table 24 and in Appendix 4 and 5, the slope coefficient of *Water* and its standardized coefficient are always smaller than those for other factors in a same equation. This indicates that *Water* is a less important predictor when compared to the others, yet more important than those factors found not significant in those equations.

In equation (12), when *Water* increases, orthophosphate decreases. This might be due to the effects of sorption by wetlands and/or the effects of dilution water on orthophosphate ion yields. In equations (13), (14), and (15) *Water* is directly proportional to TDS, Na (Figure 6), and bicarbonate, respectively. Those relations may be attributed to the contributions of ice/snow from high elevations and organic materials in the wetlands.



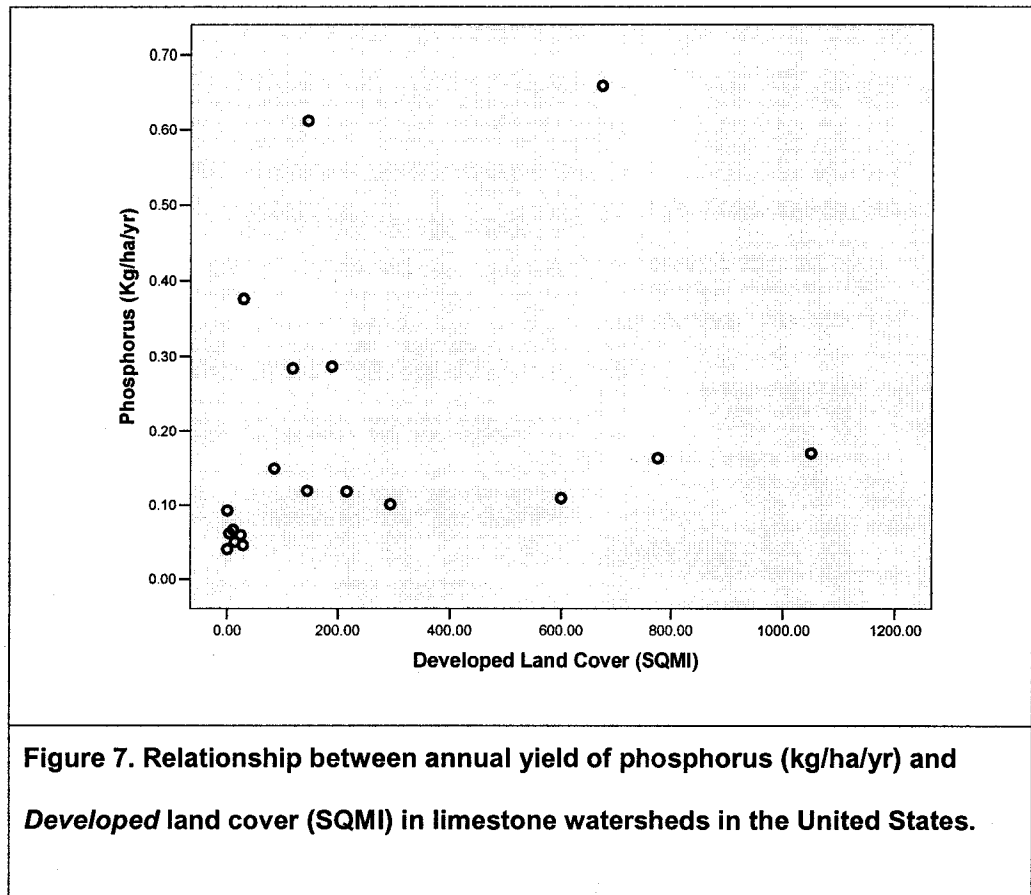
3.1.2. Effects of Developed land cover:

As shown in Table 9, *Developed* land cover occupies even smaller percentages than *Water*, ranging from 1.0% in sandstone to 4.8% in limestone, and averaged at 3.1%. The slope coefficients of this land cover type are found to be significant in limestone bedrock only (Table 25).

Table 25. The presence of significant slope coefficients of *Developed* in the regression equations in limestone bedrock (p < 0.05).

Limestone	R ²	Adj. R ²
Orthophosphate = $10^{(-0.696)} \times (\text{Developed})^{(0.417)} \times (\text{Water})^{(-0.20)}$ (16)	0.456	0.388
Phosphorus = $10^{(-1.202)} \times (\text{Developed})^{(0.186)}$ (17)	0.302	0.261

In equations (16) and (17), *Developed* is in direct relationship to orthophosphate and phosphorus (Figure 7). In equation (16), *Developed* has the highest standardized coefficient (Beta = 1.07) and partial or part correlations (0.666, or 0.658 respectively) (Appendix 4), meaning that *Developed* is the most important predictor for the prediction of orthophosphate in limestone bedrock. In equation (17) only *Developed* is a significant positive predictor for phosphorus. In both equations (16 and 17) when *Developed* increases, the corresponding dependent variable also increases, possibly due to human sewage surface disturbances causing increased rates of erosion and accumulated downstream sediment which carries phosphate.



3.1.3. Effects of Forest:

Forest cover is a highest percentage of all land cover types, ranging from about 35% in sandstone to approximately 80% in crystalline and averaged at 57.2% in all bedrock types (Table 9). However, this land cover is found to be just significant for a regression of sulfate in limestone bedrock (Table 26).

Table 26. The presence of significant slope coefficients of *Forest* in the regression equations in limestone bedrock ($p < 0.05$).

Limestone	R ²	Adj. R ²
Sulfate = $10^{(4.251)} \times (\text{Forest})^{(0.372)} \times (\text{Precipitation})^{(-2.136)}$ (18)	0.426	0.354

Forest's standardized coefficients and the corresponding partial or part correlations in equations (18) are also higher than those of *Precipitation* in a same equation (Appendix 4), indicating that *Forest* is the most important predictor for prediction of sulfate in limestone watersheds.

The role of *Forest* area in equation (18) in increasing yields of sulfate may be due to the increase of organic matter containing sulfate from the forest. Also shown in Appendix 4, *Forest* is highly correlated positively with *Agriculture* ($R = 0.90$). *Forest* and *Agriculture* plants can contribute a significant amount of sulfate into the stream system (because the organic matters come from forest and agriculture products can contain sulfate). Also, since *Forest* is correlated with *Precipitation* ($R = 0.61$, Appendix 4), the effect of acid rain (containing SO_4) is another possible reason for the increase of sulfate in the stream water.

3.1.4. Effects of Agriculture:

Agriculture is the second largest size of the 4 land cover categories, varying from 24.7% in limestone to 62.3% in crystal, and averaged at 33.2% (Table 9). However, *Agriculture* is not very significant in ion yield predictions, except for magnesium in limestone bedrock (Table 27). In equation (19), the beta value of *Agriculture* is 0.50, which is smaller than that of *Slope* (0.63) and *Precipitation* (-1.13) (Appendix 4). This means that *Agriculture* is the least important predictor for the prediction of Mg in limestone bedrock.

Table 27. The presence of significant slope coefficients of *Agriculture* in the regression equations in limestone bedrock (p < 0.05).

Limestone	R²	Adj. R²
$\text{Mg} = 10^{(6.403)} \times (\text{Precipitation})^{(-3.437)} \times (\text{Slope})^{(0.397)} \times (\text{Agriculture})^{(0.170)} \quad (19)$	0.862	0.834

In summary, land cover types may have some profound effects on ion yields, especially for orthophosphate, phosphorus, and magnesium in limestone bedrock. These effects are somehow reduced in sandstone and are not significant in crystalline bedrocks. The reasons for those phenomena are not clear and need more study.

3.2. Effects of Population:

The effects of *Population* are defined and quantified by the influence of population density (people/SQMI). In crystalline watersheds, *Population* is found of no significant contribution, but in sandstone it is significant for estimation of TDS, magnesium, potassium, and bicarbonate. In limestone watersheds, it is significant for estimation of sodium, potassium, and chloride (Table 28).

Population is always in direct proportion to the dependent variable corresponding to its contribution (Figure 8 and Figure 9). These might be due to human activities of road salt releases, chemical uses or fertilizer application, and farming activities.

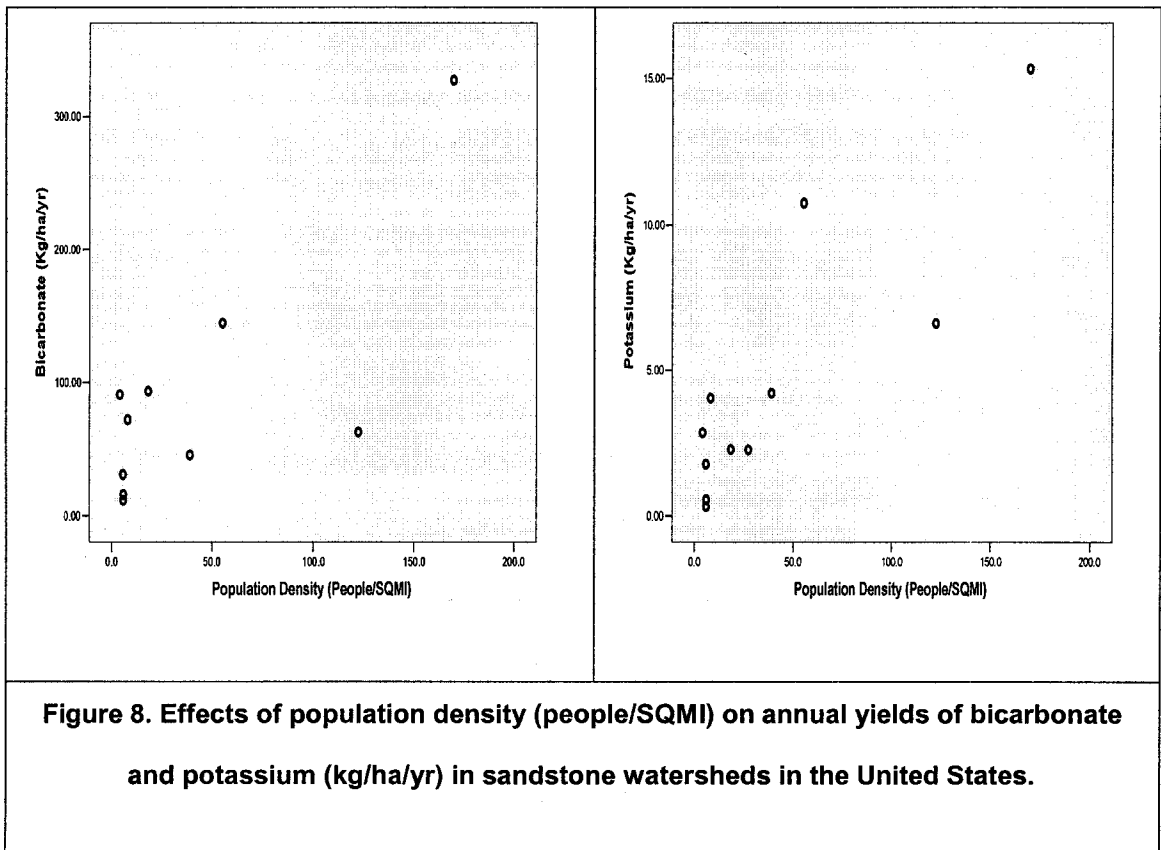
Table 28. The presence of significant slope coefficients of *Population* in the regression equations in limestone and sandstone bedrocks (p < 0.05).

Limestone		R²	Adj. R²
$Na = 10^{(2.645)} \times (Population)^{(0.619)} \times (Slope)^{(0.388)} \times (Precipitation)^{(-1.533)}$	(20)	0.783	0.740
$K = 10^{(-0.021)} \times (Slope)^{(0.399)} \times (Population)^{(0.299)}$	(21)	0.582	0.530
$Chloride = 10^{(0.435)} \times (Population)^{(0.634)}$	(22)	0.608	0.585
Sandstone			
$TDS = 10^{(1.59)} \times (Population)^{(0.489)} \times (Water)^{(0.091)}$	(23)	0.624	0.530
$Mg = 10^{(0.386)} \times (Population)^{(0.468)}$	(24)	0.376	0.307
$K = 10^{(-0.416)} \times (Population)^{(0.672)}$	(25)	0.586	0.540
$Bicarbonate = 10^{(0.777)} \times (Population)^{(0.665)} \times (Water)^{(0.129)}$	(26)	0.717	0.636

In equations (20), the beta value of *Population* is 0.877 which is almost double from that of *Slope* (0.476) and *Precipitation* (-0.390) (Appendix 4). Likewise, the values of partial and part correlations of *Population* are also the highest compared to those of *Slope* and *Precipitation* (Appendix 4). In equation (23), the values of beta, partial, and part correlations of *Population* (0.84, 0.78, and 0.763, respectively) are higher than those of *Water* (Appendix 5). Those comparisons are also likely to occur in equation (26). These results suggest that *Population* is the most important predictor in those regressions.

The positive effects of population density on sodium and chloride (equation 20 and equation 22) in this study are compatible with Peters 1984. However, the regression

correlations in this study, which are 0.74 (between population and sodium) and 0.78 (between population and chloride) (Appendix 4) are much higher than the corresponding regression correlations of Peters 1984, which were only 0.36 and 0.25, respectively.



The effects of population density on ion yields in this study are not consistent with what was found in Peters 1984. Population density was an ineffective indicator of all regressions of constituents except for sodium and chloride (Peters, 1984).

In this study, population density is significant in the regressions not only for sodium and chloride (in limestone watersheds), but also for total dissolved solids, magnesium, potassium, and bicarbonate (in sandstone watersheds) (Table 28). Furthermore, the regression correlations in these equations are much higher than those of

Peters 1984, ranging from 0.6 to 0.77 (Table 28 and Appendix 4, Appendix 5). Again, these results indicate that population density is important for estimations of these ion yields.

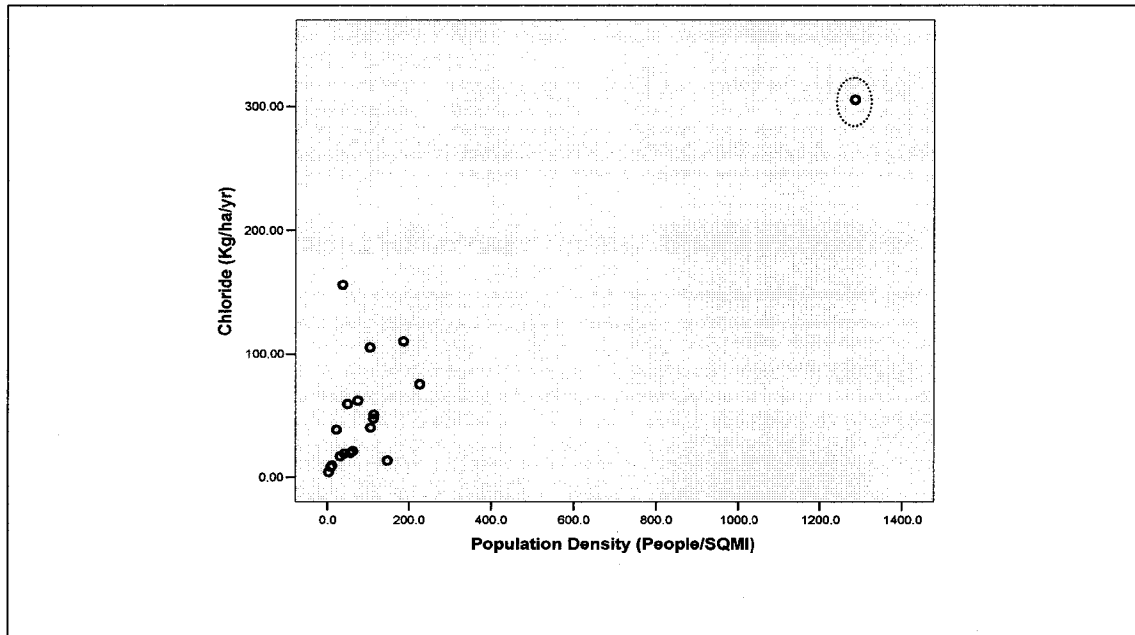


Figure 9. Effects of population density (people/SQMI) on annual yields of chloride (kg/ha/yr) in limestone watersheds in the United States.

Regression equation with the farthest point resulted in a slight larger R^2 (0.61) compared to not taking that point into account (0.50). Since that point does not affect regression very much, it is kept in the regression model (equation 22, Table 28).

3.3. Effects of Precipitation:

Unlike land cover types and *Population*, *Precipitation* is statistically significant in all three bedrocks, especially for the estimations of DO, pH, and total nitrogen (Table 17). The significant contributions of *Precipitation* to the regressions of dependent variables are shown in Table 29.

Table 29. The presence of significant slope coefficients of *Precipitation* in the regression equations in limestone, sandstone, and crystalline bedrocks ($p < 0.05$).

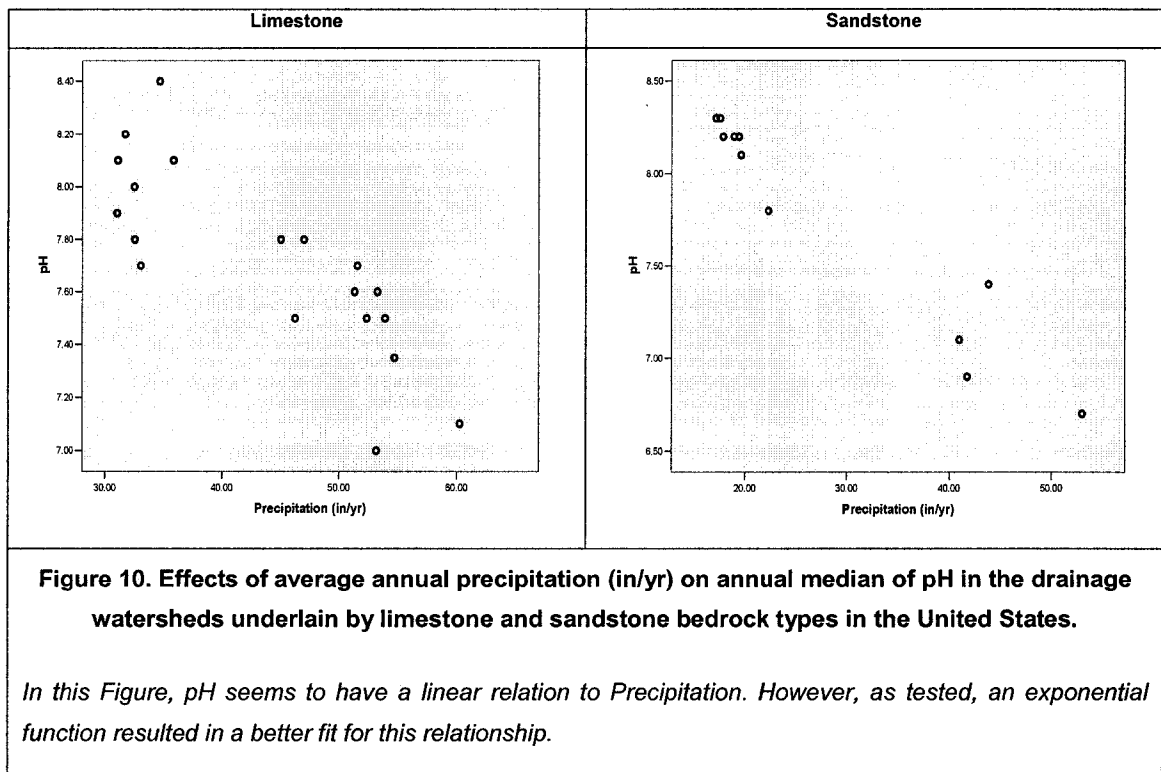
Limestone	R²	Adj. R²
$\text{pH} = 10^{(1.139)} \times (\text{Precipitation})^{(-0.155)}$ (27)	0.633	0.611
$\text{Mg} = 10^{(6.403)} \times (\text{Precipitation})^{(-3.437)} \times (\text{Slope})^{(0.397)} \times (\text{Agriculture})^{(-0.170)}$ (28)	0.862	0.834
$\text{Na} = 10^{(2.645)} \times (\text{Population})^{(0.619)} \times (\text{Slope})^{(0.388)} \times (\text{Precipitation})^{(-1.533)}$ (29)	0.783	0.740
$\text{Sulfate} = 10^{(4.251)} \times (\text{Forest})^{(0.372)} \times (\text{Precipitation})^{(-2.136)}$ (30)	0.426	0.354
$\text{Bicarbonate} = 10^{(5.407)} \times (\text{Precipitation})^{(-1.737)}$ (31)	0.364	0.311
Sandstone		
$\text{DO} = 10^{(-4.271)} \times (\text{Precipitation})^{(3.631)}$ (32)	0.843	0.821
$\text{pH} = 10^{(1.137)} \times (\text{Precipitation})^{(-0.176)}$ (33)	0.939	0.932
$\text{Total nitrogen} = 10^{(-3.790)} \times (\text{Precipitation})^{(2.678)}$ (34)	0.791	0.761
$\text{Phosphorus} = 10^{(-4.22)} \times (\text{Precipitation})^{(2.033)}$ (35)	0.516	0.447
Crystalline		
$\text{DO} = 10^{(-1.105)} \times (\text{Precipitation})^{(1.779)}$ (36)	0.482	0.477
$\text{pH} = 10^{(1.036)} \times (\text{Precipitation})^{(-0.180)}$ (37)	0.418	0.379
$\text{Total nitrogen} = 10^{(-2.083)} \times (\text{Precipitation})^{(1.599)}$ (38)	0.334	0.290

In equation (28), when compared to *Slope* and *Agriculture*, *Precipitation* has the highest values of beta (-1.13), partial and part correlations (-0.914 and -0.837) (Appendix 4). In equation (29), *Precipitation* has values of beta, partial, and part coefficients less than the corresponding values of *Population* and *Slope* (Appendix 4). In equation (30), the coefficients of *Precipitation* are also lower than those of *Forest* (Appendix 4).

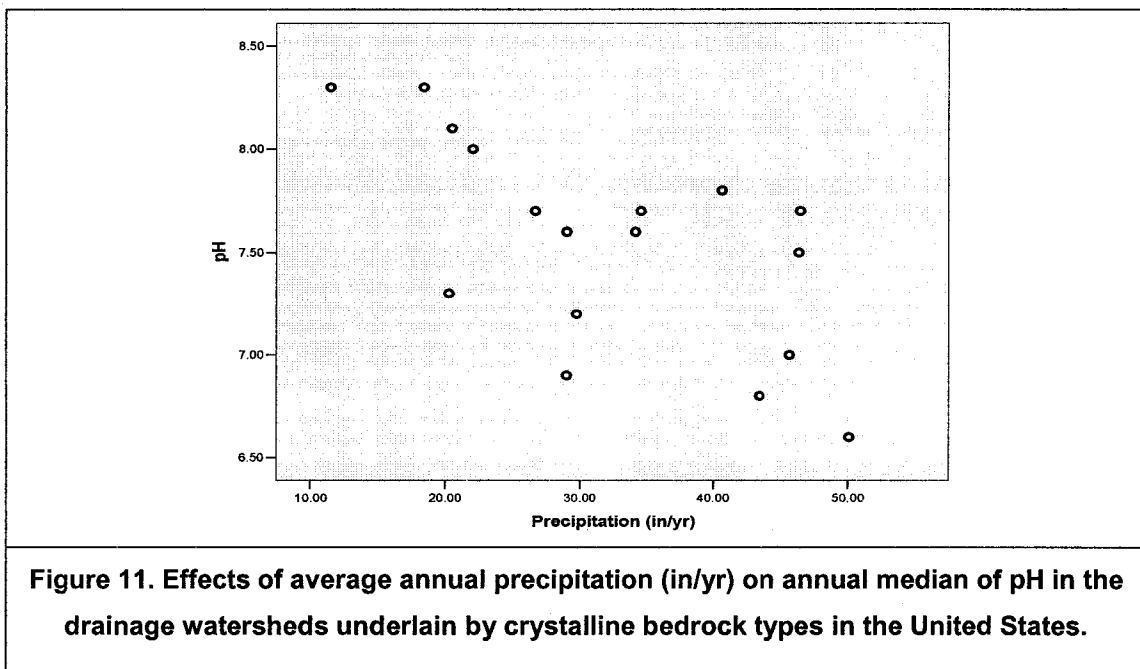
The reason for the negative relation between pH and precipitation on all three

bedrocks (Figure 10 and Figure 11) might be attributed to acid rain. Another reason for that is due to the increase of precipitation increasing the decay and decomposition of the organic materials in the soil and on the land surface. In turn, these processes release more carbon dioxide into the flow network system to the stream. In addition, the effective carbonic acid weathering in humid regions may be more closely related to the increase of dissolved solutes with increasing mean annual precipitation than in dry regions (Peters, 1984).

The atmospheric composition might be one of the main sources explaining the contribution of *Precipitation* to total nitrogen yields, phosphorus, magnesium, sodium, sulfate, and bicarbonate (Table 29). That conclusion is comparable with what was stated earlier (Peters 1984). *Precipitation* is also directly related to dissolved oxygen (Table 29) because of rainfall increasing streamflow and resulting in more mixing in the streams.



Precipitation is not significant in the regressions of total dissolved solids, calcium, potassium, and chloride like the results of Peters 1984. However, in this study, precipitation is statistically significant in many regressions for other ion yields, such as bicarbonate, pH, phosphorus, and total nitrogen (Table 29).



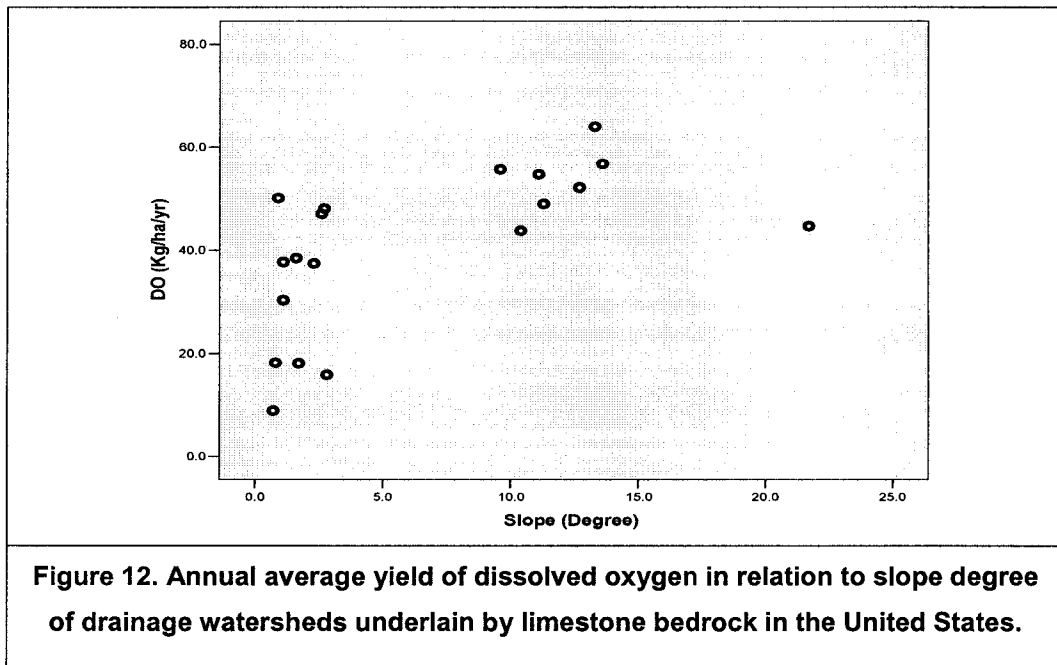
3.4. Effects of Slope:

Slope is found to be significant only in limestone watersheds. However, the effects of *Slope* are different in the regressions of different constituent yields. In the regressions with DO, *Slope* is the unique predictor (Table 30) while in the regressions of magnesium and sodium (equations 40 and 41, Table 30) *Slope* is significant in the company of other watershed characteristics (e.g. *Precipitation*, *Population*, and *Agriculture*). In the regression of potassium (equation 42), *Slope* has higher values of beta, partial, and part correlations compared to *Population* in the same equation (Appendix 4).

Table 30. The presence of significant slope coefficients of *Slope* in the regression equations in limestone and crystalline bedrocks ($p < 0.05$).

Limestone	R ²	Adj. R ²
$DO = 10^{(1.397)} \times (Slope)^{(0.296)}$ (39)	0.419	0.385
$Mg = 10^{(6.403)} \times (Precipitation)^{(-3.437)} \times (Slope)^{(0.397)} \times (Agriculture)^{(-0.170)}$ (40)	0.862	0.834
$Na = 10^{(2.645)} \times (Population)^{(0.619)} \times (Slope)^{(0.388)} \times (Precipitation)^{(-1.533)}$ (41)	0.783	0.740
$K = 10^{(-0.021)} \times (Slope)^{(0.399)} \times (Population)^{(0.299)}$ (42)	0.582	0.530

In limestone bedrock, in equation 39 (Table 30), *Slope* is directly proportional to dissolved oxygen (Figure 12). This may be due to the fact that when slope increases, the velocity of the surface flow increases resulting in more mixing in the streams. *Slope* is also in direct proportion to magnesium (equation 40), sodium (equation 41), and potassium (equation 42) (Table 30). This might be attributed to the effects of slope in increasing erosion and sediment transport.



3.5. Effects of Area:

The *Area* variable is not significant in any regression with the stream water ion yields when the significant level of alpha equals to 0.05. This indicates that *Area* is not an important variable in estimating stream water ion yields for this study. This can be explained with two reasons. The watershed sizes taken into account in this study might be too large (over the effective threshold) so that their area effects appear not discriminated from one to the others on a given water ion yield.

4. Effects of principal components of the environmental variables on stream water ion yields

As mentioned earlier in methods, principal component regression (PCR) is conducted based on principal component analysis using principal components extraction of the log transforms of the 8 environmental variables, which is rotated for ease of interpretation, including *Water, Developed, Forest, Agriculture, Precipitation, Population, Slope, and Area* of each watershed. Only components with eigenvalues greater than 1 are used for the principal component regression procedures.

4.1. Limestone watersheds:

4.1.1. Principal component analysis:

“The extraction communalities are estimates of the variance of each variable accounted for by the components in a principal component analysis” (SPSS, 2004). In this study, almost all of the extraction communalities of the principal components of the log values of the 8 environmental variables in this bedrock are high (except *Population*

and *Precipitation*), ranging from 0.87 to 0.95 (Table 31).

Table 31. Extraction communalities of the principal components of the log transforms of the 8 environmental variables in limestone watersheds.

Communalities		
	Initial	Extraction
Water	1.000	.867
Developed	1.000	.945
Forest	1.000	.904
Agriculture	1.000	.949
Precipitation	1.000	.614
Population	1.000	.473
Slope	1.000	.931
Area	1.000	.942
Extraction Method: Principal Component Analysis.		

The Rotation Sums of Squared Loadings maintains the cumulative percentage of variation explained by the extracted components (SPSS, 2004) (82.8 %, Table 32).

However, that variation is a little bit spread more over the components (after SPSS, 2004) (from 65.3 and 17.48 % to 63.5 and 19.3 %) (Table 32).

Table 32. Total variance explained for principal components in limestone watersheds.

Total Variance Explained									
Component	Initial Eigenvalues			Extraction Sums of Squared Loadings			Rotation Sums of Squared Loadings		
	Total	% of Variance	Cumulative %	Total	% of Variance	Cumulative %	Total	% of Variance	Cumulative %
1	5.2	65.3	65.3	5.2	65.3	65.3	5.1	63.5	63.5
2	1.4	17.5	82.8	1.4	17.5	82.8	1.5	19.3	82.8
3	.721	9.009	91.828						
4	.508	6.356	98.184						
5	.068	.856	99.039						
6	.050	.620	99.659						
7	.026	.322	99.982						
8	.001	.018	100.000						
Extraction Method: Principal Component Analysis.									

Together, the first two extracted components explain nearly 83% of the variability in the log transforms of the original 8 variables. Thus, we can reduce the complexity of the data set by using these two components (SPSS, 2004).

The rotated component matrix indicates what the components represent (SPSS, 2004) (Table 33). The first component is most highly correlated with *Developed* (component coefficient = 0.970), *Agriculture* (0.948), and *Area* (0.946). It is possible to estimate each variable based on the Rotated Component Matrix (Table 33). For example, $Developed = 0.970 \times Component\ 1 + 0.07 \times Component\ 2$. However, *Developed* is the best representative for component 1, because it is less correlated with component 2 (0.970 compared to 0.070). The second component is most highly correlated with *Slope* (component coefficient = 0.964) and *Precipitation* (0.505). Nevertheless, *Slope* is a better representative for the second component because it is also much less correlated with the first component (Table 33).

It should note that using the two saved components (as shown in Table 32 and Table 33) is more preferable than using each individual variable such as *Developed*, *Agriculture*, and *Slope* because the two components are representative of all log transforms of the 8 original variables, and they are not linearly correlated with each other (after SPSS, 2004).

Table 33. The Rotated Component Matrix of the log transforms of the 8 variables in limestone watersheds.

Rotated Component Matrix(a)		
	Component	
	1	2
Water	.875	-.320
Developed	.970	.070
Forest	.868	.387
Agriculture	.948	.221
Precipitation	.599	.505
Population	.682	-.089
Slope	-.048	.964
Area	.946	.217
Extraction Method: Principal Component Analysis. Rotation Method: Varimax with Kaiser Normalization.		
a Rotation converged in 3 iterations.		

The Regression Factor Scores (REGR factor score, Table 35) is computed by multiplying the case's standardized variable values by the component's score coefficients (SPSS, 2004) (Table 34).

Table 34. Component Score Coefficient Matrix of the log transforms of the 8 variables in limestone watersheds.

Component Score Coefficient Matrix		
	Component	
	1	2
Water	.217	-.310
Developed	.198	-.049
Forest	.145	.182
Agriculture	.178	.058
Precipitation	.076	.291
Population	.153	-.130
Slope	-.107	.675
Area	.178	.055
Extraction Method: Principal Component Analysis. Rotation Method: Varimax with Kaiser Normalization. Component Scores.		

Table 35. Regression Factor Scores for the first 2 principal components of the log transforms of the 8 variables in limestone watersheds.

Station No.	REGR factor score 1	REGR factor score 2
USGS 02313000	0.520	-1.090
USGS 02320500	0.905	-0.805
USGS 02329000	0.183	-0.413
USGS 02368000	-0.201	0.163
USGS 03085000	0.564	0.788
USGS 03254000	-0.016	0.857
USGS 03290500	0.320	1.055
USGS 03425000	0.523	1.086
USGS 03543005	0.831	1.109
USGS 03571850	0.952	1.106
USGS 03593005	1.156	0.989
USGS 04045500	-0.657	-1.110
USGS 04057004	-0.315	-1.273
USGS 04108690	0.296	-0.917
USGS 04126520	-0.160	-0.542
USGS 04132052	-0.191	-0.719
USGS 04165500	0.214	-1.567
USGS 08103900	-2.026	-0.103
USGS 13018300	-2.895	1.385

*(REGR factor score 1 = 0.217 * ZWater + 0.198 * ZDeveloped+ 0.145 * ZForest + 0.178 * ZAgriculture + 0.076 * ZPrecipitation + 0.153 * ZPopulation + - 0.107 * ZSlope + 0.178 * ZArea;*

*REGR factor score 2 = -0.310 * ZWater + -0.049 * ZDeveloped + 0.182 * ZForest + 0.058 * ZAgriculture + 0.291 * ZPrecipitation + -0.130 * ZPopulation + 0.675 * ZSlope + 0.055 * ZArea)*

(ZWater = (log transformed value of Water (i) – average log transformed value of all log transformed Water values)/std. deviation of all log transformed Water values). This formula is applied similarly for Zscores of the other variables).

4.1.2. Principal component regression:

There are no significant regressions between the first two selected principal components with total dissolved solids, total nitrogen, orthophosphate, calcium, magnesium, potassium, chloride, sulfate, and bicarbonate (Appendix 4). The PCRs of dissolved oxygen, pH, and phosphorus are shown in Table 36. However, the squared correlations are not very high, ranging from 0.217 to 0.284. It is noted that component 1 is significant in the regressions of pH because as mentioned earlier in equation (27) (Table 29), pH is negatively related to *Precipitation* and the component coefficient of *Precipitation* in component 1 is higher than that in component 2 (0.60 compared to 0.51, Table 33). Thus, the reason that pH is also negatively related to component 2 because of the effect of *Precipitation* on component 2 and *Precipitation*'s relation to pH. Similarly, the reason that phosphorus is positively related to component 1 because it is positively related to *Developed* as shown in equation (17) (Table 25) while *Developed* also has the highest correlation with component 1 as shown in Table 33 (the component correlation coefficient equals to 0.970). Likewise, dissolved oxygen is positively related to component 2 because it is positively related to *Slope* as shown in equation (39, Table 30) while *Slope* also has the highest correlation with component 2 (the component correlation coefficient equals to 0.964) (Table 33).

Table 36. Significant principal component regressions in limestone watersheds (p < 0.05).

Limestone	R ²	Adj. R ²
DO = 10 [1.562 + 0.123 x (REGR factor score 2)] (43)	.284	.242
pH = 10 [0.887 - 0.01 x (REGR factor score 1)] (44)	.256	.212
Phosphorus = 10 [-0.882 + 0.17 x (REGR factor score 1)] (45)	.217	.171

It is necessary to understand that although component 1 has correlations with pH and phosphorus lower than that of component 2 (Table 36), as shown in Table 32, component 1 is much more important because it is accounted for 63.5 % of the total variation of the 8 variables while component 2 is accounted for 19.3 % of the total variation only. In addition, the squared correlations of the principal component regressions in Table 36 are lower than those of the corresponding multiple linear regressions in equation (39) (Table 30), equation (27) (Table 29), and in equation (17) (Table 25).

4.2. Sandstone watersheds:

4.2.1. Principal component analysis:

The extractions of the principal components of the environmental variables for sandstone watersheds are higher than those of limestone watersheds, ranging from 0.71 to 0.99 (Table 37). Therefore, the extracted components represent the variables well (SPSS, 2004).

Table 37. Extraction communities of the principal components of the log transforms of the 8 environmental variables in sandstone watersheds.

Communalities		
	Initial	Extraction
Water	1.000	.944
Developed	1.000	.974
Forest	1.000	.906
Agriculture	1.000	.982
Precipitation	1.000	.841
Population	1.000	.936
Slope	1.000	.707
Area	1.000	.989
Extraction Method: Principal Component Analysis.		

The cumulative percentage of variation explained by the extracted components is maintained of the rotation loadings method (approximately 91%) (Table 38). However, that variation is spread more over the components (53.9 and 37.1 instead of 71.6 and 19.4) (after SPSS, 2004) (Table 38).

Table 38. Total variance explained for principal components in sandstone watersheds.

Total Variance Explained									
Component	Initial Eigenvalues			Extraction Sums of Squared Loadings			Rotation Sums of Squared Loadings		
	Total	% of Variance	Cumulative %	Total	% of Variance	Cumulative %	Total	% of Variance	Cumulative %
1	5.7	71.6	71.6	5.7	71.6	71.6	4.3	53.9	53.9
2	1.6	19.4	91.0	1.6	19.4	91.0	3.0	37.1	91.0
3	.494	6.174	97.157						
4	.140	1.755	98.912						
5	.054	.669	99.581						
6	.020	.246	99.828						
7	.013	.158	99.986						
8	.001	.014	100.000						

Extraction Method: Principal Component Analysis.

We can also use the scree plot, which presents eigenvalue of each component in the initial solution, to determine the optimal number of components (Figure 13). Generally, the components on the steep slope are extracted, because the components on the shallow slope contribute little to the solution (SPSS, 2004). As shown in Figure 13, the last big drop occurs between the second and the third components. Thus, using the first two components is a better option (SPSS, 2004).

Scree Plot

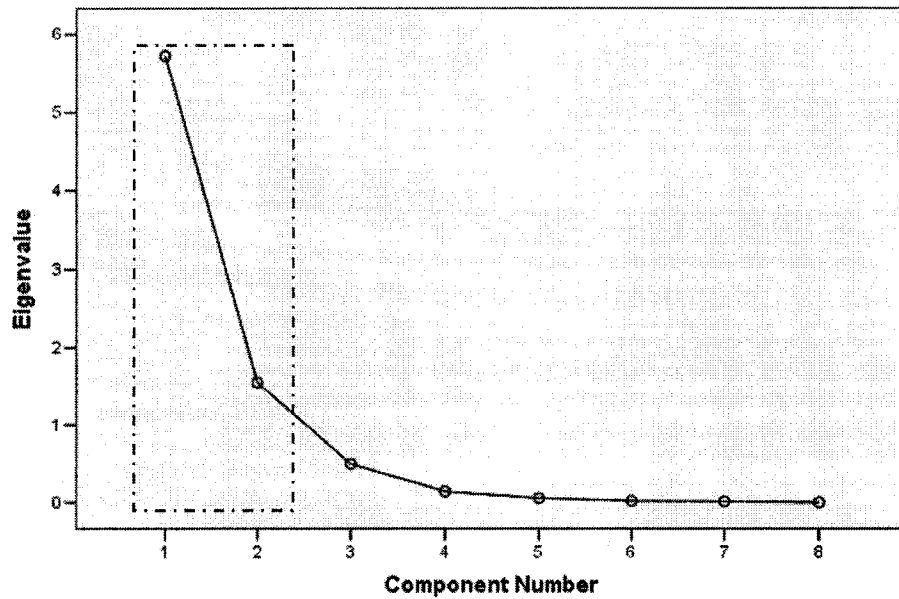


Figure 13. Scree plot of the initial eigenvalue of each component in sandstone watersheds.

In the rotated component matrix (Table 39), the first component is most highly correlated with *Forest*, *Developed*, and *Water*. However, *Forest* is the best representative, because it is less correlated with the second component. The second component is most highly correlated with *Population* and *Precipitation* followed by *Area and Agriculture*. Nevertheless, *Population* is a better representative for the second component because it is also much less correlated with the first component (Table 39) (component correlation coefficient of *Population* with component 2 is 0.964 while the coefficient with the first component is only -0.086).

Table 39. The Rotated Component Matrix of the log transforms of the 8 variables in sandstone watersheds.

Rotated Component Matrix(a)		
	Component	
	1	2
Water	.910	-.341
Developed	.921	-.354
Forest	.947	.098
Agriculture	.847	-.515
Precipitation	-.235	.887
Population	-.086	.964
Slope	-.417	.730
Area	.888	-.449
Extraction Method: Principal Component Analysis. Rotation Method: Varimax with Kaiser Normalization.		
a Rotation converged in 3 iterations.		

The two regression factor scores (Table 41) used for the PCR procedures were calculated by multiplying component score coefficients (Table 40) by the corresponding standardized values (Zscores).

Table 40. Component Score Coefficient Matrix of the log transforms of the 8 variables in sandstone watersheds.

Component Score Coefficient Matrix		
	Component	
	1	2
Water	.228	.037
Developed	.229	.033
Forest	.337	.258
Agriculture	.168	-.062
Precipitation	.118	.378
Population	.185	.448
Slope	.023	.261
Area	.196	-.021
Extraction Method: Principal Component Analysis. Rotation Method: Varimax with Kaiser Normalization. Component Scores.		

Table 41. Regression Factor Scores for the first 2 principal components of the log transforms of the 8 variables in sandstone watersheds.

Station No.	REGR factor score 1	REGR factor score 2
USGS 01545600	-1.497	0.737
USGS 03049625	1.260	1.724
USGS 03237280	-1.994	1.001
USGS 06452000	0.004	-0.861
USGS 06478500	0.103	-1.031
USGS 07026000	0.394	0.886
USGS 07231500	0.806	0.052
USGS 07234000	-0.326	-1.229
USGS 07237500	-0.057	-1.009
USGS 07245000	1.109	0.267
USGS 06438000	0.198	-0.538

4.2.2. Principal component regression:

There is no regression between the first two selected principal components with orthophosphate, phosphorus, Mg, chloride, and sulfate. The PCRs of total dissolved solids, dissolved oxygen, pH, total nitrogen, calcium, potassium, sodium, and bicarbonate are shown in Table 42. As shown in Table 42, the *REGR factor score 2* is significant positively in most of those regressions (except for the PCR of pH) while *REGR factor score 1* is significant positively in the regression of sodium only.

In general, the squared correlations of the principal component regressions expressed in Table 42 are higher than those in limestone watersheds (Table 36), ranging from 0.39 to 0.83.

Table 42. Principal component regressions in sandstone watersheds (p < 0.05)

Sandstone	R ²	Adj. R ²
TDS = 10 ^{2.334 + 0.213 x (REGR factor score 2)} (46)	.408	.343
DO = 10 ^{0.873 + 0.735 x (REGR factor score 2)} (47)	.861	.841
pH = 10 ^{0.888 - 0.029 x (REGR factor score 2)} (48)	.673	.637
Total nitrogen = 10 ^{0.002 + 0.551 x (REGR factor score 2)} (49)	.834	.810
Ca = 10 ^{1.426 + 0.242 x (REGR factor score 2)} (50)	.406	.340
Na = 10 ^{1.422 + 0.262 x (REGR factor score 1)} (51)	.539	.488
K = 10 ^{0.448 + 0.402 x (REGR factor score 2)} (52)	.640	.600
Bicarbonate = 10 ^{1.776 + 0.291 x (REGR factor score 2)} (53)	.495	.432

*(REGR factor score 1 = 0.228 * ZWater + 0.229 * ZDeveloped+ 0.337 * ZForest + 0.168 * ZAgriculture + 0.118 * ZPrecipitation + 0.185 * ZPopulation + - 0.023 * ZSlope + 0.196 * ZArea;*

*REGR factor score 2 = 0.037 * ZWater + 0.033 * ZDeveloped + 0.258 * ZForest + -0.062 * ZAgriculture + 0.378 * ZPrecipitation + 0.448 * ZPopulation + 0.261 * ZSlope + -0.021 * ZArea)*

(ZWater = (log transformed value of Water (i) – average log transformed value of all log transformed Water values)/std. deviation of all log transformed Water values).

The reason that total dissolved solids, potassium, and bicarbonate are positively related to component 2 because they are proportional related to *Population* (equations 23

through 26, Table 28), which has the highest component correlation with component 2 (0.964) while *Population*'s correlation with component 1 is only -0.086 (Table 39). Similarly, the reason that dissolved solids, pH, total nitrogen, and phosphorus are significant related to component 2 because they are related to *Precipitation* (equations 32 - 35, Table 29), which has the second highest component correlation with component 2 (0.887) while *Precipitation*'s correlation with component 1 is only -0.235 (Table 39). Likewise, the reason that sodium is positively related to component 1 because it is positively related to *Water* (equation 14, Table 24), which has much higher component correlation (0.910) with component 1 compared to (-0.341) of component 2 (Table 39).

Note that while almost all principal component regressions in Table 42 have higher squared correlations than those of the corresponding multiple linear regressions, some of them (e.g. regressions of dissolved oxygen, pH, and bicarbonate) have lower squared correlations when compared to those of equations (23) and (26) (Table 28), and equation (33) (Table 29) . In addition, although principal component regression gives higher correlations, it is not necessary that it give better predictions because most of the regressions in Table 42 are significant with component 2, which is accounted for only 37% of the total variation of the 8 independent variables.

4.3. Crystalline watersheds:

4.3.1. Principal component analysis:

The extractions of the principal components of the log values of the 8 environmental variables in crystalline bedrock are also higher than those of limestone watersheds, ranging from 0.71 to 0.98 (Table 43).

Table 43. Extraction communities of the principal components of the log transforms of the 8 environmental variables in crystalline watersheds.

Communalities		
	Initial	Extraction
Water	1.000	.921
Developed	1.000	.873
Forest	1.000	.906
Agriculture	1.000	.848
Precipitation	1.000	.864
Population	1.000	.705
Slope	1.000	.977
Area	1.000	.952
Extraction Method: Principal Component Analysis.		

Unlike the results for limestone and sandstone watersheds, crystalline watersheds have three components containing eigenvalues greater than 1. Also, the variation explained by the extracted components is more evenly spread over the rotated components (Table 44). Together, the first three extracted components explain nearly 88% of the variability in the log transforms of the original 8 variables.

Table 44. Total variance explained for principal components in crystalline watersheds.

Total Variance Explained									
Component	Initial Eigenvalues			Extraction Sums of Squared Loadings			Rotation Sums of Squared Loadings		
	Total	% of Variance	Cumulative %	Total	% of Variance	Cumulative %	Total	% of Variance	Cumulative %
1	4.2	53.0	53.0	4.2	53.0	53.0	3.8	47.6	47.6
2	1.6	20.4	73.5	1.6	20.4	73.5	1.7	21.6	69.3
3	1.2	14.6	88.1	1.2	14.6	88.1	1.5	18.8	88.1
4	.554	6.924	95.004						
5	.284	3.556	98.560						
6	.072	.904	99.463						
7	.042	.522	99.985						
8	.001	.015	100.000						
Extraction Method: Principal Component Analysis.									

As shown in the rotated component matrix (Table 45), the first component is most highly correlated with *Forest*, *Agriculture*, and *Developed* land covers (0.933, 0.920, and 0.854, respectively, Table 45). However, *Agriculture* and *Forest* might be the best representatives for component 1 because they are less correlated with the other components (Table 45). The second component is most highly correlated with *Precipitation* and *Population* (0.913 and 0.810, Table 45). *Precipitation* is a very good representative for the second component because it is also much less correlated with the first component and the third one. The third component is highly correlated with *Slope* (-0.983, Table 45), and *Slope* is also much less correlated with the other two components (Table 45). Therefore, *Slope* is the best representative for component 3.

Table 45. The Rotated Component Matrix of the log transforms of the 8 variables in crystalline watersheds.

Rotated Component Matrix(a)			
	Component		
	1	2	3
Water	.663	.294	.629
Developed	.854	.378	-.033
Forest	.933	-.071	.173
Agriculture	.920	.036	.032
Precipitation	-.026	.913	-.173
Population	.138	.810	.174
Slope	-.074	.071	-.983
Area	.949	-.010	.229
Extraction Method: Principal Component Analysis. Rotation Method: Varimax with Kaiser Normalization.			
a Rotation converged in 4 iterations.			

The highest component score coefficient in component 1 is *Agriculture* (0.285) while in component 2 it is *Precipitation* (0.554), and in component 3, it is *Slope* (-0.745) (Table 46).

Table 46. Component Score Coefficient Matrix of the log transforms of the 8 variables in crystalline watersheds.

Component Score Coefficient Matrix			
	Component		
	1	2	3
Water	.067	.133	.371
Developed	.244	.147	-.174
Forest	.273	-.127	-.042
Agriculture	.285	-.065	-.146
Precipitation	-.063	.553	-.104
Population	-.066	.484	.131
Slope	.153	.023	-.745
Area	.262	-.090	.000

Extraction Method: Principal Component Analysis.
Rotation Method: Varimax with Kaiser Normalization.
Component Scores.

The regression factor scores (Table 47) for the PCR analysis were calculated by multiplying the component score coefficients (Table 46) by the corresponding standardized values (Zscores).

Table 47. Regression Factor Scores for the first 3 principal components of the log transforms of the 8 variables in crystalline watersheds.

Station No.	REGR factor score 1	REGR factor score 2	REGR factor score 3
USGS 01059400	1.152	1.184	0.493
USGS 01066000	0.621	1.299	0.421
USGS 04014500	-0.713	-0.027	1.864
USGS 05124480	-0.766	-0.205	1.952
USGS 06623800	-0.374	0.109	-0.361
USGS 07311200	-1.279	1.027	0.985
USGS 08431700	-0.542	-1.344	-0.614
USGS 09352900	-0.049	0.870	-1.075
USGS 09430600	-0.587	-0.655	-1.322
USGS 09508300	-1.472	-0.651	-0.502
USGS 11264500	0.048	1.017	-0.589
USGS 12447390	-0.931	0.115	-1.219
USGS 13169500	0.090	-2.156	0.743
USGS 13331500	0.196	0.688	-0.967
USGS 14048000	1.859	-1.249	-0.368
USGS 14103000	1.855	-0.585	0.440
USGS 14113000	0.892	0.563	0.118

4.3.2. Principal component regression:

In crystalline bedrocks, there are only 2 principal component regressions are significant, that is, the PCR of dissolved oxygen and PCR of pH (Table 48). *REGR factor score 1* and *REGR factor score 3* are not significant in any of these 2 regressions while *REGR factor score 2* is positively significant in the relation with dissolved oxygen and negatively significant in the relation with pH. The reason for that is because of the significant relations between *Precipitation* and dissolved oxygen, and between *Precipitation* and pH as shown in equations (36) and (37), Table 29 while *Precipitation* has the highest component correlation with component 2 but is much less correlated with component 1 and component 3 (Table 45).

Table 48. Significant principal component regressions in crystalline watersheds ($p < 0.05$).

Crystalline	R^2	Adj. R^2
DO = 10 [1.525 + 0.246 x (REGR factor score 2)] (54)	.296	.249
pH = 10 [0.876 - 0.02 x (REGR factor score 2)] (55)	.450	.413

In general, the squared correlations of the principal component regressions expressed in Table 48 are higher than those in limestone watersheds (Table 36) but lower than those in sandstone watersheds (Table 42). In addition, the squared correlations in Table 48 are lower than those of the corresponding multiple linear regressions in equations (36) and (37) (Table 29).

5. Summary of the principal component regression analysis

Principal component regression (PCR) is based on principal component analysis and the multiple regression technique. In a principal component analysis, all variables are used to compute components. This technique reduces the number of variables and therefore, gives a simple descriptive if the principal components are easily interpreted. In addition, principal components are nearly not collinearity. Unlike the multiple regression technique, PCR composites variables rather than eliminates variables in the regression processes.

In this study, generally speaking, PCR equations have lower correlation (except for the regressions of total dissolved solids, pH, and bicarbonate in sandstone watershed, Table 49) and approximately equal Standard Error of the Estimate when compared to those of multiple regression technique using stepwise method of model selection (Table 49). Besides that, using multiple regressions can detect more relationship than PCR (Table 49). In addition, most principal component regressions in Tables 37, 43, and 49 are significant with component 2 which is accounted for only 20 to 37% of the total variation of the 8 independent variables. Those clues indicate that using PCR equations for ion yield estimations might not be as good as using the corresponding multiple regression equations. This inference will be tested again in the next step of **Stream water ion yield prediction, model validation, and spatial distribution of the predicted ion yields.**

So why do most PCR equations have lower correlations than those of multiple regression technique using stepwise method of model selection? This can be explained that PCR takes into account all independent variables although some of them are not

actually effective predictors (e.g. *Slope* and *Area*) while multiple regression selects only those predictors most affecting the ion yield responses into an equation (e.g. *Precipitation* and *Population*).

Table 49. Comparison of Multiple Regression and Principal Component Regression using Correlation and Standard Error of the Estimate for each bedrock type (more details are presented in Appendix 4, 5, 6).				
Limestone				
	Multiple Regression Using Log Scale		Principal Component Regression Using Log Scale	
	R²	Std. Error of the Estimate	R²	Std. Error of the Estimate
TDS	No sig. reg.		No sig. reg.	
DO	.419	0.181	.284	0.201
pH	.633	0.013	.256	0.018
Total Nitrogen	No sig. reg.		No sig. reg.	
Orthophosphate	.456	0.329	No sig. reg.	
Phosphorus	.302	0.313	.217	0.332
Ca	No sig. reg.		No sig. reg.	
Mg	.862	0.130	No sig. reg.	
Na	.783	0.210	No sig. reg.	
K	.582	0.234	No sig. reg.	
Chloride	.608	0.305	No sig. reg.	
Sulfate	.426	0.332	No sig. reg.	
Bicarbonate	.364	0.249	No sig. reg.	
Sandstone				
	Multiple Regression using log scale		Principal Component Regression using log scale	
	R²	Std. Error of the Estimate	R²	Std. Error of the Estimate
TDS	.624	0.228	.408	0.270
DO	.843	0.342	.861	0.322
pH	.939	0.009	.673	0.021
Total Nitrogen	.791	0.300	.834	0.268
Orthophosphate	No sig. reg.		No sig. reg.	
Phosphorus	.516	0.430	No sig. reg.	
Ca	No sig. reg.		.406	.309
Mg	.376	0.363	No sig. reg.	
Na	.453	0.278	.539	0.255
K	.586	0.341	.640	0.317
Chloride	No sig. reg.		No sig. reg.	
Sulfate	No sig. reg.		No sig. reg.	
Bicarbonate	.717	0.262	.495	0.327
Crystalline				
	Multiple Regression using log scale		Principal Component Regression using log scale	
	R²	Std. Error of the Estimate	R²	Std. Error of the Estimate
TDS	No sig. reg.			
DO	.482	0.337	.31	0.39
pH	.418	0.023	.63	0.020
Total Nitrogen	.334	0.412	No sig. reg.	
Orthophosphate	No sig. reg.		No sig. reg.	
Phosphorus	No sig. reg.		No sig. reg.	
Ca	No sig. reg.		No sig. reg.	
Mg	No sig. reg.		No sig. reg.	
Na	No sig. reg.		No sig. reg.	
K	No sig. reg.		No sig. reg.	
Chloride	No sig. reg.		No sig. reg.	
Sulfate	No sig. reg.		No sig. reg.	
Bicarbonate	No sig. reg.		No sig. reg.	

6. Stream water ion yield prediction, model validation, and spatial distribution of the predicted ion yields.

There are only 4 stations available for the testing data in sandstone bedrock (see Appendices 1, 2, 3 and the websites listed in “References for additional (testing) sandstone stations” at the end of Literature Cited section). This data set is too small for an actual validation. However, this step is necessary for a completed conceptual methodology. Based on the data available for the 4 stations, and based on high correlations and a compatible comparison between multiple regression and PCR approaches, the following equations in sandstone watersheds were selected for predictions (Table 50).

Table 50. Selected equations for predictions and validations of stream water ion yields in sandstone watersheds.			
Multiple Regression		R²	Adj. R²
$DO = 10^{(-4.271)} \times (\text{Precipitation})^{(3.631)}$	(32)	0.843	0.821
$pH = 10^{(1.137)} \times (\text{Precipitation})^{(-0.176)}$	(33)	0.939	0.932
$\text{Total nitrogen} = 10^{(-3.790)} \times (\text{Precipitation})^{(2.678)}$	(34)	0.791	0.761
$K = 10^{(-0.416)} \times (\text{Population})^{(0.672)}$	(25) *	0.586	0.540
Principal Component Regression		R²	Adj. R²
$DO = 10^{[0.873 + 0.735 \times (\text{REGR factor score } 2)]}$	(47)	0.861	0.841
$pH = 10^{[0.888 - 0.029 \times (\text{REGR factor score } 2)]}$	(48)	0.673	0.637
$\text{Total nitrogen} = 10^{[0.002 + 0.551 \times (\text{REGR factor score } 2)]}$	(49)	0.834	0.810
$K = 10^{[0.448 + 0.402 \times (\text{REGR factor score } 2)]}$	(52) *	0.640	0.600

(*: Although the regressions of potassium yield does not have both multiple regression and PCR square coefficients ≥ 0.640 , these equations are taken into account for prediction just for a comparison between the 2 regression techniques).

The predicted values of ion yields are shown in Table 51 and the corresponding Average Percentage Errors are shown in Table 52.

Table 51. Predicted Ion Yields Using Multiple Regression, Principal Component Regression, and GIS techniques in sandstone watersheds.				
	DO	pH	Total Nitrogen	K
Multiple Regression	<i>kg/ha/yr</i>	<i>units</i>	<i>kg/ha/yr</i>	<i>kg/ha/yr</i>
USGS 08053500	21.8	7.3	2.2	6.2
USGS 06810000	17.2	7.4	1.9	4.2
USGS 06485500	5.7	7.8	0.8	4.0
USGS 05413500	17.5	7.4	1.9	4.7
PCR				
USGS 08053500	13.9	7.5	1.6	3.9
USGS 06810000	9.1	7.7	1.2	3.1
USGS 06485500	4.6	7.9	0.7	2.1
USGS 05413500	12.6	7.6	1.5	3.7
Actual Values				
USGS 08053500	6.2	8.1	NoData	2.7
USGS 06810000	20.9	7.9	NoData	7.2
USGS 06485500	19.6	8.3	5.7	12.0
USGS 05413500	NoData	8.3	NoData	8.1
PCR Using GIS				
USGS 08053500	15.1	7.6	1.6	3.9
USGS 06810000	13.2	7.6	1.5	3.7
USGS 06485500	4.8	8.0	0.7	2.0
USGS 05413500	19.2	7.5	2.0	4.6
Multiple Regression Using GIS				
USGS 08053500	21.9	7.3	2.2	5.3
USGS 06810000	17.3	7.4	1.9	4.1
USGS 06485500	5.9	7.8	0.8	3.5
USGS 05413500	17.5	7.4	1.9	4.7

Table 52. Average Percent Error (APE, in fraction) for Predicted Ion Yields Using Multiple Regression, Principal Component Regression, and GIS techniques in sandstone watersheds.				
	DO	pH	Total Nitrogen	K
Multiple Regression	<i>(APE, in fraction)</i>	<i>(APE, in fraction)</i>	<i>(APE, in fraction)</i>	<i>(APE, in fraction)</i>
USGS 08053500	-2.52	0.10	NoData	-1.30
USGS 06810000	0.18	0.06	NoData	0.42
USGS 06485500	0.71	0.06	0.86	0.67
USGS 05413500	NoData	0.11	NoData	0.42
General Average Percent Error (GAPE)*	1.13	0.08	0.86	0.70
PCR				
USGS 08053500	-1.24	0.07	NoData	-0.44
USGS 06810000	0.56	0.03	NoData	0.57
USGS 06485500	0.77	0.05	0.88	0.83
USGS 05413500	NoData	0.08	NoData	0.54
General Average Percent Error (GAPE)*	0.86	0.06	0.88	0.60
PCR Using GIS				
USGS 08053500	-1.4	0.1	NoData	-0.4
USGS 06810000	0.4	0.0	NoData	0.5
USGS 06485500	0.8	0.0	0.9	0.8
USGS 05413500	NoData	0.1	NoData	0.4
General Average Percent Error (GAPE)*	0.85	0.06	0.88	0.55
Multiple Regression Using GIS				
USGS 08053500	-2.53	0.10	NoData	-0.96
USGS 06810000	0.17	0.06	NoData	0.43
USGS 06485500	0.70	0.06	0.86	0.71
USGS 05413500	NoData	0.11	NoData	0.42
General Average Percent Error (GAPE)*	1.13	0.08	0.86	0.63

General Average Percent Error (GAPE) is calculated by absolute values of APE from each watershed because we are interested in understanding how far the predicted results are from the actual values using different methods.*

1- For multiple regression and PCR not using GIS layers: dissolved oxygen, pH, total nitrogen, potassium, and bicarbonate are predicted.

a. For multiple regression, pH has the smallest Average Percent Error (APE, 0.08)

followed by potassium (0.70), total nitrogen (0.86), and dissolved oxygen (1.13) (Table 52)

b. For PCR, pH has the smallest APE (0.06) followed by potassium (0.60), dissolved oxygen (0.86), and total nitrogen (0.88) (Table 52).

2- For multiple regression and PCR using GIS layers: dissolved oxygen, pH, total nitrogen, and potassium are predicted.

a. For multiple regression, pH has the smallest APE (0.08) followed by potassium (0.63), total nitrogen (0.86), and dissolved oxygen (1.13) (Table 52).

b. For PCR, pH has the smallest APE (0.06), followed by potassium (0.55), dissolved oxygen (0.85), and total nitrogen (0.88) (Table 52).

From (1) and (2), it is possible to conclude that the ion yield predictions made by multiple regression and PCR are not very much different although using PCR has lower APE for pH and potassium. Practically speaking, multiple regression is much preferred because it requires a smaller number of independent variable data. Besides that, the use of GIS models gives almost the same average predicted results as calculated by the corresponding regressions without using GIS. While the non-GIS models just use average values for each ion yield of each watershed (lumped/aggregated), GIS models take into account spatial distribution of the input data (disaggregated) which may be far from those average values taken to build the regression models. Therefore, using the GIS approach may introduce more standard deviation (Table 53) and larger range of the results than non-GIS method.

In addition, the GIS models of PCR give larger standard deviation and larger range of the predictions when compared to GIS models of multiple regressions, except for potassium (Table 53, and Figure 14 – Figure 16). For example, as shown in Figure 14, the range of dissolved oxygen predicted by PCR using GIS is from 0 – 61.3 (kg/ha/yr) while the range of dissolved oxygen predicted by multiple regression using GIS is only from 3.3 – 11.0 (kg/ha/yr). The reason for this might be due to the fact that PCR takes into account all variables for prediction so that the range of the input data is also larger than some selected variables used in the multiple regressions, and consequently, PCR using GIS introduces more uncertainty.

Nevertheless, using multiple regressions with GIS-based models is a better choice for ion yield prediction in this study because this approach not only gives the same average predicted results as using the regression models without GIS, but also gives a visualization depicting the spatial effects. Therefore, we can see and delineate different areas of different levels of ion yields (Figure 14 – Figure 16). As shown, the shape and pattern of the map of predicted ion yield reflects the shape and the pattern of the corresponding predictor. For example, the map of potassium predicted (Table 14) using multiple linear regression and GIS is similar to the map of population density because potassium depends on population density, which was mapped by counties. However, the map of potassium predicted using PCR and GIS (Table 14) is much smoother compared to that of using multiple linear regression due to effects of other variables with higher resolution that were significant in the model.

Station No.	DO		pH		Total Nitrogen		K	
	Multiple Re.	PCR	Multiple Re.	PCR	Multiple Re.	PCR	Multiple Re.	PCR
USGS 08053500	3.10	11.8	0.05	0.2	0.23	0.9	4.59	1.5
USGS 06810000	1.83	6.8	0.04	0.2	0.15	0.6	1.86	1.0
USGS 06485500	1.74	4.3	0.11	0.3	0.18	0.4	2.61	0.9
USGS 05413500	0.00	6.2	0.00	0.1	0.00	0.5	0.00	0.9

Generally speaking, as indicated in this study, the regression models for predicting pH are most valid because they have the smallest APE (0.06 – 0.08) (Table 52), and regression models for predicting dissolved oxygen are least valid because of highest APE (1.13, Table 52). However, those models are still very useful for ion yield predictions of those watersheds where no water quality data are available.

For multiple regression perspective, most of the significant regression equations in this study have higher percentage variance of each constituent explained by the related predictors, compared to those of Peters 1984. For example, in Peters 1984 most regression squared correlations were just from 0.4 to 0.7 while in this study, most regression squared correlations are from 0.5 to 0.93 (Table 17, Table 42, and Table 48).

Figure 14. Predicted values of DO, pH, Total Nitrogen, and K using multiple regression and PCR equations for station No. USGS 06485500 - Big Sioux R At Akron, IA.



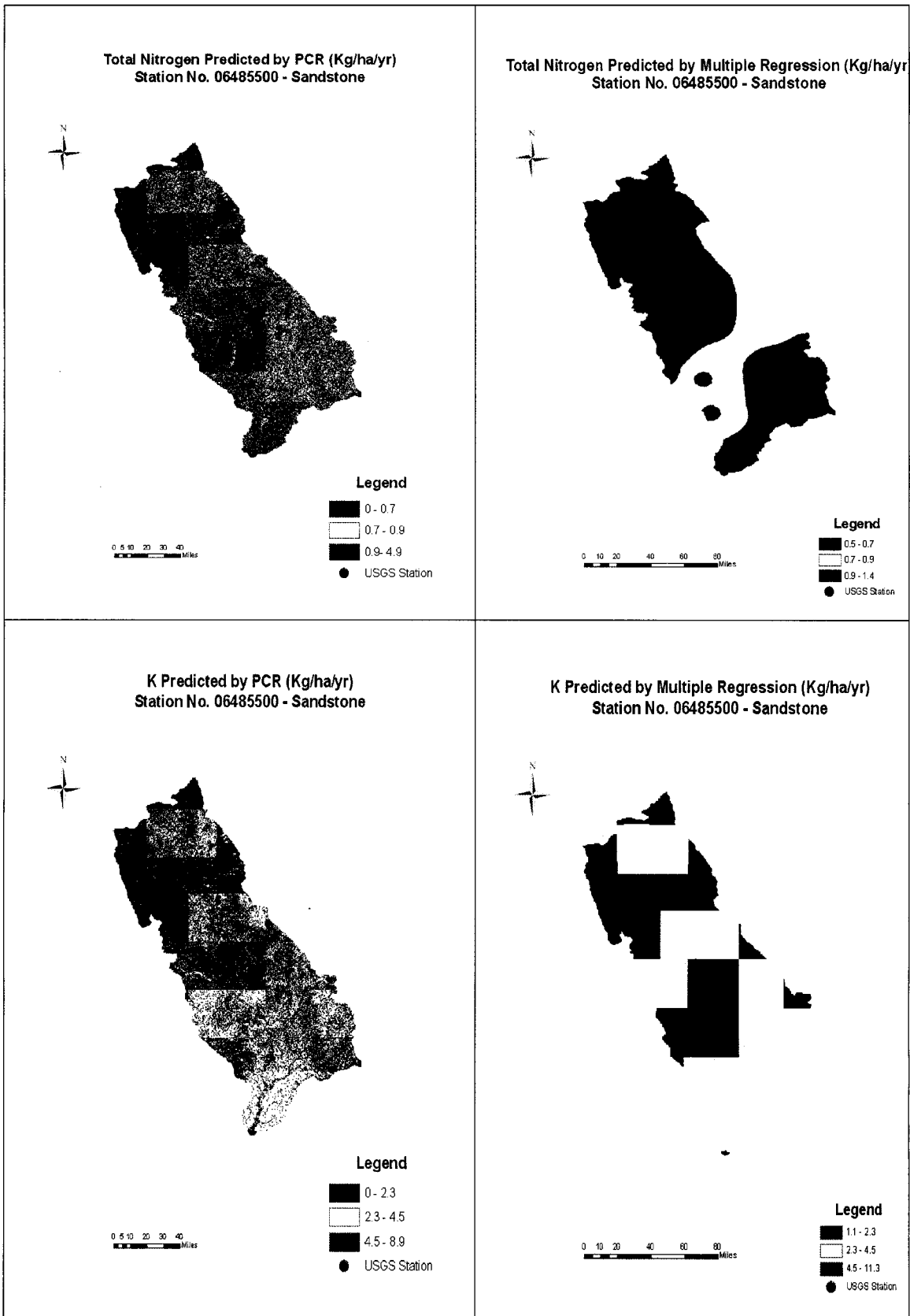
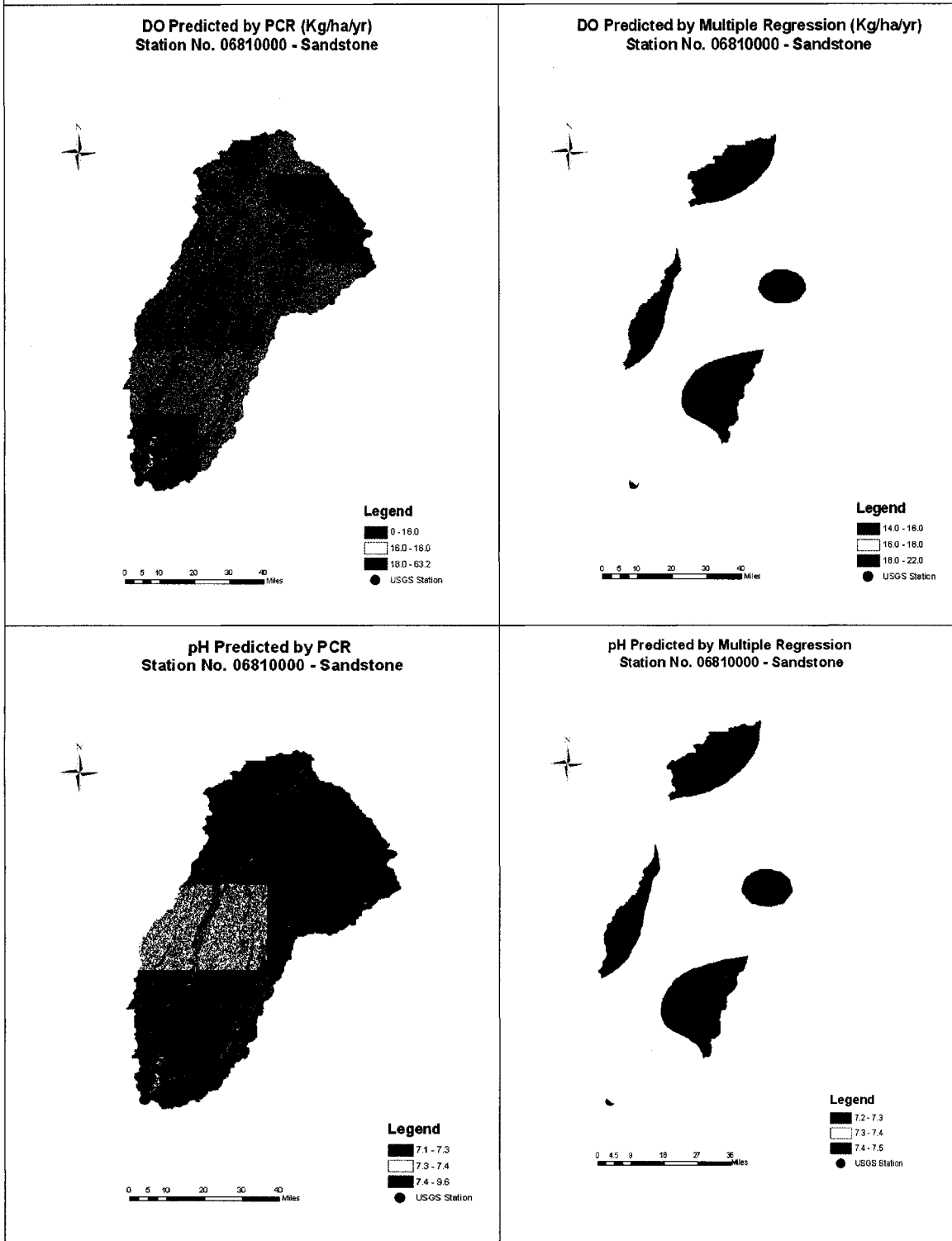


Figure 15. Predicted values of DO, pH, Total Nitrogen, and K using multiple regression and PCR equations for station No. USGS 06810000 - Nishnabotna River above Hamburg, IA.



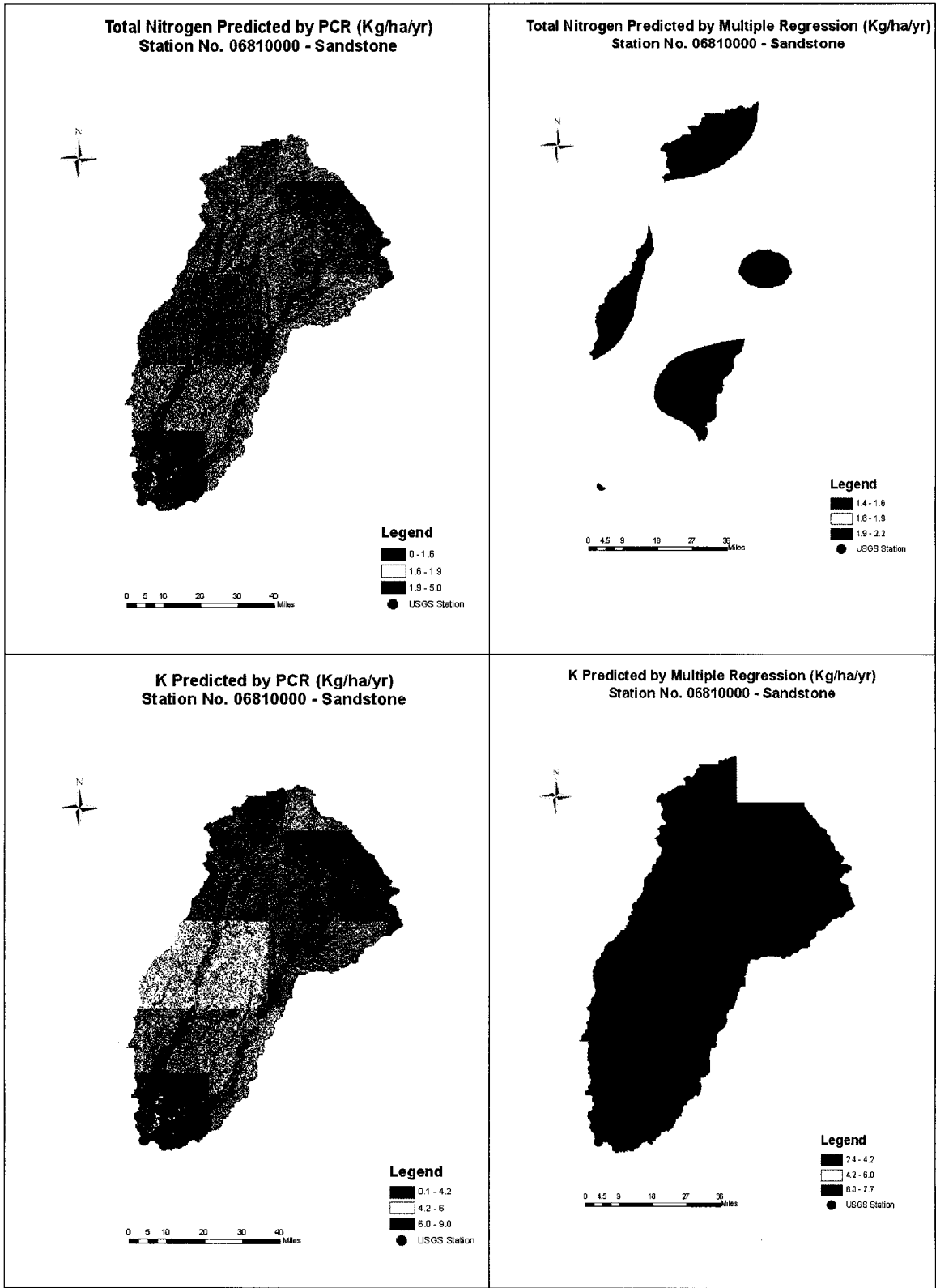
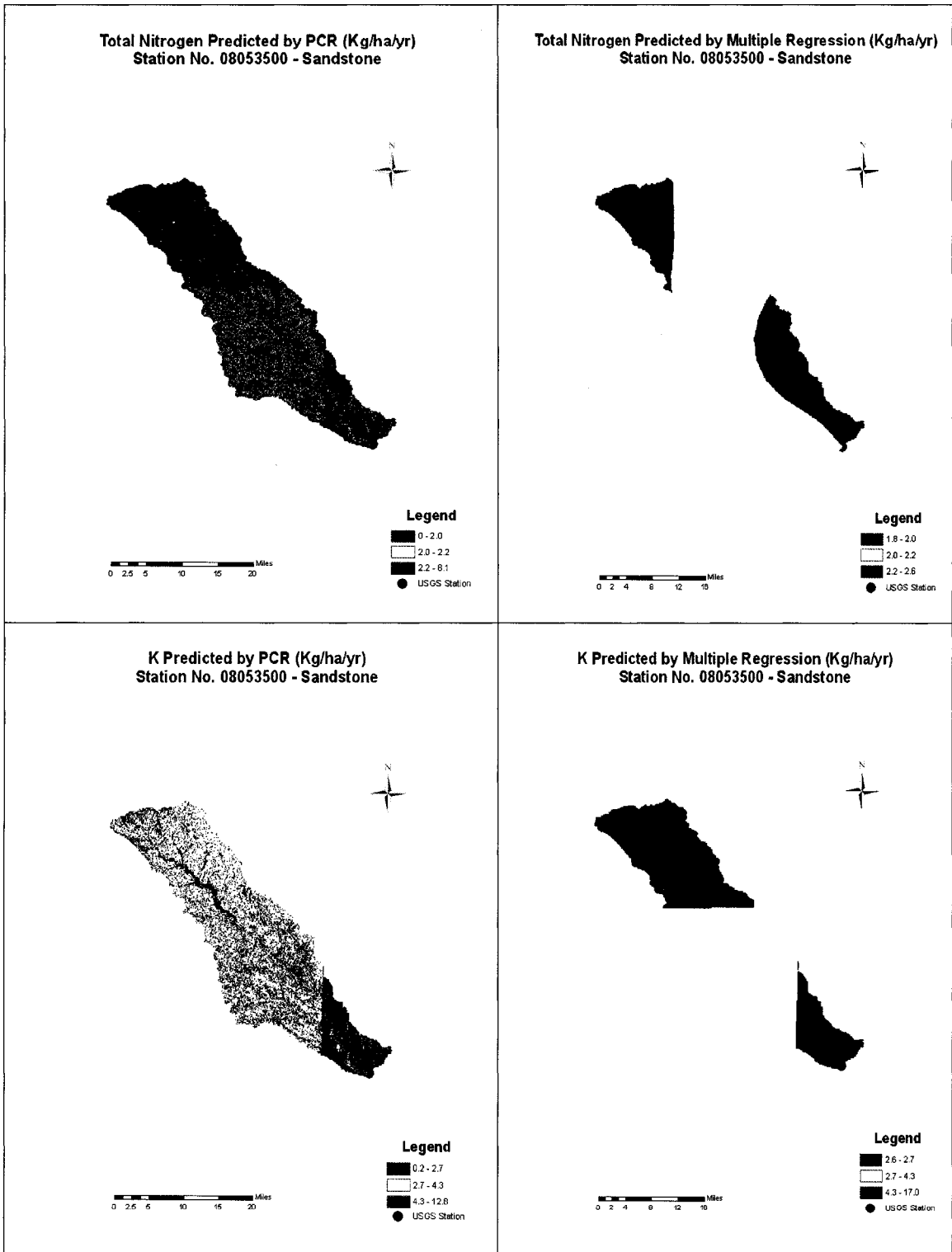


Figure 16. Predicted values of DO, pH, Total Nitrogen, and K using multiple regression and PCR equations for station No. USGS 08053500 - Denton Ck nr Justin, TX.





V. CONCLUSIONS

In this study, 13 stream water constituents – total dissolved solids, dissolved oxygen, pH, total nitrogen, orthophosphate, phosphorus, calcium, magnesium, sodium, potassium, sulfate, chloride, and bicarbonate from 47 watersheds in the United States were analyzed in relation to the environmental factors, including bedrock types (limestone, sandstone, and crystalline), land cover type, average annual precipitation, population density, watershed slope, and watershed area.

The results from this study indicate that there is a strong exponential relationship among environmental factors and selected stream water ion yields for selected watersheds. The most important factors affecting annual stream water ion yields are average annual precipitation and bedrock type. Stream water ion yields of limestone bedrock are much different from those of the other two, including total dissolved solids, orthophosphate, calcium, magnesium, and bicarbonate. Sandstone and crystalline ion yields are not very much different, except for chloride and sulfate. For example, some average annual ion yields of limestone watersheds are 2 to 5 times larger than those of sandstone or crystalline watersheds, including total dissolved solids (723 kg/ha/yr), bicarbonate (441 kg/ha/yr), magnesium (36 kg/ha/yr), chloride (61 kg/ha/yr), sulfate (123 kg/ha/yr), and calcium (144 kg/ha/yr). However, calcium is not related to any environmental factor of all three bedrocks except for the principal component regression of calcium in sandstone watershed ($R^2 = 0.406$).

The yields of dissolved oxygen, potassium, sodium, and total nitrogen were not different among the three bedrock types nor was the median pH values. Ion yields in

limestone and sandstone bedrocks are more closely related to the environmental factors than those in crystalline bedrock. Most of the ion yields in crystalline bedrock have no relation to the environmental factors, except for dissolved oxygen, pH, and total nitrogen.

For the stepwise multiple regression of this study shows that out of 22 significant equations of all watersheds in the three bedrock types, average annual precipitation appears in 11 equations, including regressions of dissolved oxygen (sandstone- $R^2 = 0.84$ and crystalline- $R^2 = 0.48$), pH (in all three bedrocks- R^2 ranges from 0.42 to 0.63 and 0.94), total nitrogen (sandstone- $R^2 = 0.79$ and crystalline- $R^2 = 0.33$), phosphorus (sandstone- $R^2 = 0.52$), magnesium (limestone- $R^2 = 0.37$), sodium (limestone- $R^2 = 0.12$), sulfate (limestone- $R^2 = 0.18$), and bicarbonate (limestone- $R^2 = 0.36$).

In the significant regressions, average annual precipitation is either a unique predictor (in 9 equations) or among the most significant predictors. These significant regressions and the corresponding squared regression coefficient values indicate that average annual precipitation in company with other predictors is the most important factor affecting stream water ion yields, accounting for 12 to 94 percent of the variance of ion yields of sodium, sulfate, bicarbonate, magnesium, phosphorus, total nitrogen, dissolved oxygen and pH. The validation of this study indicates that it is applicable to use average annual precipitation to predict pH in all three bedrock types.

Followed by bedrock type and average annual precipitation, population density is the next most important factor in most ion yield prediction. Out of the 22 significant equations of all watersheds in the three bedrock types, population density is found to be in 7 equations, including regressions of total dissolved solids (sandstone- $R^2 = 0.37$),

sodium (limestone- $R^2 = 0.54$), magnesium (sandstone- $R^2 = 0.31$), potassium (limestone- $R^2 = 0.26$, sandstone- $R^2 = 0.54$), chloride (limestone- $R^2 = 0.61$), and bicarbonate (sandstone- $R^2 = 0.41$). Like average annual precipitation, in these regressions, population density is either a unique predictor or among the most significant predictors having highest beta values. The significant squared regression coefficient values indicate that population density, in company with other factors, accounted for 26 to 61 percent of the variance of ion yields of potassium, magnesium, total dissolved solids, bicarbonate, sodium, and chloride.

Nevertheless, population is not present in the models of crystalline watersheds. In limestone watersheds, population affects sodium and chloride most, while in sandstone watersheds, it affects total dissolved, magnesium, potassium, and bicarbonate most.

In principal component regression analysis, precipitation and population are also among the highest score coefficient weighted factors, especially for sandstone watersheds. This explains why the predicted yield maps of dissolved oxygen, pH, total nitrogen, and potassium in sandstone watersheds show a similar shape and pattern of population maps.

Land cover type also affects ion yields in limestone and sandstone watersheds. The effects of the four land cover types (*Water*, *Developed*, *Forest*, and *Agriculture*) on yield prediction are not very different, although *Developed* and *Water* cover types tend to be more effective. For example, *Developed* appears in the regressions of orthophosphate and phosphorus in limestone watersheds ($R^2 = 0.28$), *Water* appears in the regressions of orthophosphate in limestone watershed ($R^2 = 0.19$), total dissolved solids ($R^2 = 0.25$),

sodium ($R^2 = 0.45$), and bicarbonate ($R^2 = 0.31$) in sandstone watersheds while *Agriculture* is found in only a regression of magnesium ($R^2 = 0.15$), and *Forest* is found in only a regression of sulfate ($R^2 = 0.24$) in limestone watersheds. These results indicate that effects of land cover types were significant with other factors and accounted for 15 to 45 percent of the variance of ion yields in limestone and sandstone watersheds, including orthophosphate, phosphorus, magnesium, and sulfate (limestone watersheds), total dissolved solids, sodium, and bicarbonate (sandstone watersheds).

Watershed slope was an ineffective predictor, significant only in limestone watersheds in company with other factors, accounted for 12 to 42 percent of variance of ion yields of sodium, potassium, magnesium, and dissolved oxygen. Watershed area was not significant in any ion yield regression analysis.

Multiple linear regression (MLR) provided simpler models but larger R square, smaller standard of error of estimate, and could detect more significant equations than principal component regression (PCR).

The use of GIS-based regression models to predict ion yields introduces more variation and uncertainty because of the larger range of input values, compared to those selected to build the regression models, although the mean predicted values of non-GIS and GIS models are almost the same. This GIS-based approach, however, is very useful for visualization and comparison of different patterns of predicted ion yield.

VI. RECOMENDATIONS

The digital environmental data obtained by the National Elevation Dataset, National Land Cover Dataset, National Hydrography Dataset, Census data from the National Atlas, and Average Annual Precipitation from Data Gateway have an accuracy level which is most suitable for large areas like state wide or national wide regions, especially for population and land cover datasets. Thus, applying these datasets to obtain data for the small watersheds, such as some of this study, is not very applicable. This explains why land cover types are not effective for the ion yield prediction models in this study.

The frequency of data of this study is based mainly on average annual basis (e.g. precipitation). In fact, constituent yields vary very much in different seasons and even in different months of the year. Thus, shorter time periods of the data is recommended. In addition, the effects of soil types are not considered (due to difficulty of obtaining data), although soil is very important for water quality. Both training and testing sample sizes are small. Point source pollution is not separated from non point sources because of the difficulty of collecting point sources due to the fact that most of the project watersheds are large. The use of mean slope values is not very suitable for large watersheds. These limitations conceal and even can distort the relationship between stream ion yields and the environmental factors. Furthermore, the effects of bedrock types on ion yields are indirectly reflected and are not very explicitly because other factors were not controlled. The GIS-based prediction results were based on a small testing sample size and on a cell-

by-cell basis without taking accumulative effects into account. Thus, those results would be used for relative spatial distribution comparison of ion yields only.

The prediction of stream ion yields using 4 selected sandstone watersheds results some high Average Percent Error (APE) values (e.g. APE of total nitrogen). These results might be partly due to some parts of the testing watersheds do not contain sandstones. Nevertheless, prediction is necessary and useful for environmental planning, remediation, and protection. Prediction can be improved by adding soil type into the regression processes, because water quality is very sensitive to soil properties in the drainage system. In addition, the effects of slope and watershed area should be reconsidered, although they are not very effective for ion yield estimation as shown in this study. The use of larger data sets, more accurate, and shorter frequency can also improve the regression models.

GIS-based implementation models are very robust and useful for a spatial delineation of predicted stream water ion yields. However, to reduce the uncertainty introduced by this method, one of the most effective ways is to collect as various ranges of the input data as possible. This would help to assure that the range of the input data in each GIS layer is in the same scale of or not much different from the data ranges used to build regression models.

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<http://seamless.usgs.gov/website/seamless/viewer.php>

<http://edcftp.cr.usgs.gov/pub/data/landcover/states/>

<http://www.mrlc.gov/>

<http://gisdata.usgs.net/website/MRLC/viewer.php>

<http://www.mrlc.gov/scripts/mapserv.exe?map=d%3A%5Cinetpub%5Cwwwroot%5Cipc%5Cmrlc2k%5Czones%5Czones.map>

Websites referenced for the accuracy of land cover:

<http://landcover.usgs.gov/accuracy/>

Websites referenced for Population Density of the US in 1990:

<http://www.nationalatlas.gov>

Websites referenced for Precipitation and Hydrologic Units:

<http://datagateway.nrcs.usda.gov/GatewayHome.html>

Websites referenced for Hydrologic Units and Flowlines:

<http://nhd.usgs.gov/>

Websites referenced for Water Quality Data:

<http://water.usgs.gov/waterwatch/>

<http://waterdata.usgs.gov/nwis>

Websites referenced for additional (testing) sandstone stations:

<http://digicoll.library.wisc.edu/cgi-bin/EcoNatRes/EcoNatRes-idx?type=div&did=ECONATRES.0027.0280.0010&isize=M>

<http://72.14.203.104/search?q=cache:CEfXiJCFhYsJ:www.seta.iastate.edu/publicservice/s/water/plan/chapter3/+sandstone+of+of+cambrian-ordovician+bedrock&hl=en&gl=us&ct=clnk&cd=9>

http://72.14.203.104/search?q=cache:iWHCtloK_VoJ:courses.unt.edu/hwilliams/lab_7.htm+sandstone+bedrock+and+water&hl=en&gl=us&ct=clnk&cd=9

<http://www.igsb.uiowa.edu/gsbpubs/pdf/OFM-1997-1.pdf>

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http://plantsci.sdstate.edu/woodardh/Geology/Eastern/General_Geology/general_geology.htm

<http://www.northern.edu/natsource/EARTH/Aquife1.htm>

<http://www.beg.utexas.edu/UTopia/images/pagesizemaps/physiography.pdf#search=%22bedrock%20map%20of%20texas%22>

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http://www.cretaceousfossils.com/formations/trinity_group/paluxy_sandstone.htm

APPENDICES

All appendices are recoded in a CD attached in the inner side of the back cover of the dissertation. They can be open, printed, and copied. Appendices 1, 2, 3, and 7 are in Excel Worksheets. Appendices 4, 5, 6, and 8 are in Microsoft Word Documents.

Appendix 1: Station Information.

Appendix 2a. Water Quality Data and Corresponding Number of Records Compiled from <http://waterdata.usgs.gov/nwis> for the Project Watersheds in Limestone Bedrock.

Appendix 2b. Water Quality Data and Corresponding Number of Records Compiled from <http://waterdata.usgs.gov/nwis> for the Project Watersheds in Sandstone Bedrock.

Appendix 2c. Water Quality Data and Corresponding Number of Records Compiled from <http://waterdata.usgs.gov/nwis> for the Project Watersheds in Crystalline Bedrock.

Appendix 3a_1. Land Cover, Precipitation, Population, Slope, and Area Data in the Project Watersheds in Limestone Bedrock.

Appendix 3a_2. Water Quality Data in the Project Watersheds in Limestone Bedrock (kg/ha/yr).

Appendix 3a_3. Combined Land Cover Classes in the Project Watersheds in Limestone Bedrock (% & SQMI).

Appendix 3a_4. Explanation of Land Cover codes and Combined Land Cover Classes.

Appendix 3b_1. Land Cover, Precipitation, Population, Slope, and Area Data in the Project Watersheds in Sandstone Bedrock.

Appendix 3b_2. Water Quality Data in the Project Watersheds in Sandstone Bedrock (kg/ha/yr).

Appendix 3b_3. Combined Land Cover Classes in the Project Watersheds in Sandstone Bedrock (% & SQMI).

Appendix 3c_1. Land Cover, Precipitation, Population, Slope, and Area Data in the Project Watersheds in Crystalline Bedrock.

Appendix 3c_2. Water Quality Data in the Project Watersheds in Crystalline Bedrock (kg/ha/yr).

Appendix 3c_3. Combined Land Cover Classes in the Project Watersheds in Crystalline Bedrock (% & SQMI).

Appendix 4_1. Correlation matrix and Multiple Linear Regression between the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Enter” Method of Model Selection to see a Multicollinearity Diagnosis for Watersheds in Limestone Bedrock.

Appendix 4_2. Multiple Linear Regression between the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Stepwise” Method of Model Selection for Watersheds in Limestone Bedrock ($\alpha = 0.05$).

Appendix 4_3. Factor Analysis and Principal Component Regression among the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Stepwise” Method of Model Selection for Watersheds in Limestone Bedrock ($\alpha = 0.05$).

Appendix 5_1. Correlation matrix and Multiple Linear Regression between the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Enter” Method of Model Selection to see a Multicollinearity Diagnosis for Watersheds in Sandstone Bedrock.

Appendix 5_2. Multiple Linear Regression between the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Stepwise” Method of Model Selection for Watersheds in Sandstone Bedrock ($\alpha = 0.05$).

Appendix 5_3. Factor Analysis and Principal Component Regression among the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Stepwise” Method of Model Selection for Watersheds in Sandstone Bedrock ($\alpha = 0.05$).

Appendix 6_1. Correlation matrix and Multiple Linear Regression between the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Enter” Method of Model Selection to see a Multicollinearity Diagnosis for Watersheds in Crystalline Bedrock.

Appendix 6_2. Multiple Linear Regression between the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Stepwise” Method of Model Selection for Watersheds in Crystalline Bedrock ($\alpha = 0.05$).

Appendix 6_3. Factor Analysis and Principal Component Regression among the Predefined Environmental Factors and Stream Water Ion Yields (kg/ha/yr) Using Log-Transformed Values and “Stepwise” Method of Model Selection for Watersheds in Crystalline Bedrock ($\alpha = 0.05$).

Appendix 7. Comparison between average annual ion yields (kg/ha/yr) from Peters's research in 1984 and the current research.

Appendix 8. Watersheds Delineated in the Three Bedrock Types.