THESIS

THE DEVELOPMENT AND VALIDATION OF AN X-BAND DUAL POLARIZATION DOPPLER WEATHER RADAR TEST NODE FOR A TROPICAL NETWORK

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In partial fulfillment of the requirements
For the degree of Master of Science
Colorado State University
Fort Collins, Colorado
Fall 2012

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ABSTRACT

THE DEVELOPMENT AND VALIDATION OF AN X-BAND DUAL POLARIZATION DOPPLER WEATHER RADAR TEST NODE FOR A TROPICAL NETWORK

An automated network of three X-band dual polarization Doppler weather radars is in process of being deployed and operational on the western coast of Puerto Rico. Colorado State University and the University of Puerto Rico at Mayaguez have collaborated to install the first polarimetric weather radar network in a tropical environment, known as TropiNet, to observe the lowest 2 km of the troposphere where the National Weather Service NEXRAD radar in Cayey, PR (TJUA) has obstructed views of the west coast, below 1.5 km due to terrain blockage and the Earth curvature problem. The CSU-X25P radar test node was developed, validated, and deployed to Mayaguez, PR in early 2011 to make first observations of this tropical region, and served as a pilot project to verify the infrastructure of the TropiNet network.

This research describes the CSU-X25P radar test node, presenting the radar system specifications and an overview of the data acquisition and signal processing sub-systems, and the antenna positioner and control sub-system. The development and validation process included integration, sub-system calibration and test, and a final evaluation by conducting end-to-end calibration of the radar system. Validation of the calculated data moments, include Doppler velocity, reflectivity, differential reflectivity, differential propagation phase, and specific differential phase. The validation was accomplished by comparative analysis of data from coordinated scans between CSU-X25P and the well-established CSU-CHILL S-band polarimetric Doppler weather radar, in Greeley, CO.

Upon validation, CSU-X25P was disassembled, packaged, and shipped to Puerto Rico to be fully deployed for operation in a tropical seaside environment. This research presents select observations of severe weather events, such as tropical storms and hurricanes, which attest to the robustness of the radar test node, and the TropiNet network infrastructure.
I want to thank Dr. Francesc Junyent who shared his invaluable RF and Radar systems expertise. The staff at the CSU-CHILL Radar Facility were instrumental in facilitating the space for the assembly, testing, calibrating, and shipping of this radar. Many thanks to Dave Brunkow, Patrick Kennedy, Jim George, and especially Bob Bowie.

To the UPRM faculty, staff, and students who collaborated in the successful deployment of CSU-X25P to Puerto Rico, I owe a debt of gratitude. In particular, Dr. Jose G. Colom, Dr. Sandra Cruz-Pol, Dr. Rafael A. Rodriguez Solis, Don Carlos Mercado, Jose Rosario, Benjamin De Jesus, Jose Cordero, and Jose Kotshie Ortiz to name a few.

I would like to recognize Dr. Omnia El-Hakim who first believed that I could become a professor, and recommended me for the NSF-funded AMP Bridge to Doctorate Fellowship. I want to thank all my colleagues in the Radar Lab for their continued support over the years. Thank you Henry Obermeyer for donating engineering resources and materials to fabricate the transceiver mounting plate. To the girls I had the honor of coaching in soccer while working on my Masters, thanks for bringing me balance.

Lastly, Mil Gracias to my lovely wife Araceli, and my cousin Fernando Gálvez who kept me on track to finish this chapter of my life. It is done.
DEDICATION

To my wife Araceli, and my children Gemma del Rocio and Diego Miguel who were my support system throughout this process. You are the reason I do this. For the example of my father Miguel who was a man with the resolve and work ethic of ten. And for my mother Dora who selflessly sacrificed herself to always ensure that I could follow my dreams.
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A.1 SCR-584 Pedestal Riser Dimensions.
CHAPTER 1

Introduction

Meteorological radars have become an integral part of our lives with continuous demand from users for up-to-the-minute forecasts and severe weather warnings. The National Weather Service (NWS) has been providing the general public with periodic updates for many years via radio and television broadcasts, and now over the internet and even on portable devices, such as smart phones. There is an increasing demand for current, reliable, and higher resolution weather forecasts from NWS, which requires a complementary short-ranged, low-scanning, higher resolution, networked radar approach like that of CASA IP-1 [2, 3].

In 2006, the Collaborative Adaptive Sensing of the Atmosphere (CASA) Engineering Research Center deployed the first generation network of four low-power, short-range, X-band, polarimetric Doppler weather radars in Oklahoma’s "tornado alley," near Norman, OK. CASA is a collaboration of several universities that includes Colorado State University (CSU), University of Massachusetts Amherst, University of Oklahoma, and University of Puerto Rico at Mayaguez (UPRM). It was soon apparent that such a networked strategy could improve observations in tropical environments, where a single high-powered, long-range NEXRAD radar had obstructed views of the lowest 2 km of the troposphere. A major advantage of going to an operating frequency in the X-band is the radar nodes are a much smaller and more cost-effective solution relative to the currently employed NEXRAD WSR-88D radars, operating at S-band. This prompted UPRM and CSU to begin investigating the feasibility of deploying a network of dual-polarized X-band Doppler weather radars in the tropical environment of Puerto Rico.

An automated network of three X-band polarimetric Doppler weather radars (TropiNET) is in process of being deployed on the western coast of Puerto Rico, as proposed in [4]. UPRM and CSU have collaborated to identify potential site locations and prepared the infrastructure for this network. Prior to the development of the TropiNET radar network, a single X-band
polarimetric Doppler weather radar, CSU-X25P was deployed in January 2011 to make observations in the tropical environment around Mayaguez, Puerto Rico as a pilot project. This radar is mounted on a tower atop the R&D Center building on the UPRM campus, in Figure 1.1, and is maintained and operated locally and remotely by staff from both UPRM and CSU. The CSU-X25P radar node is the first dual polarization Doppler weather radar to be deployed in Puerto Rico, and made first observations of the lower troposphere of the islands west coast, where the region has a lack of coverage at low altitudes by the WSR-88D radar (TJUA) in Cayey, PR, due to the Earth’s curvature and terrain blockage.

The CSU-X25P radar node description, development, validation, and select first observations by a polarimetric X-band Doppler weather Radar of the western coast of Puerto Rico will be presented herein. This Thesis will provide detail of the system presented here in the following chapters.

Chapter 2 will present a detailed system description of the CSU-X25P radar node. Details
of the transceiver analog hardware are given, including system specifications, RF section, and system level block diagrams. An overview of the data acquisition system gives background information on the analog-to-digital conversion of the intermediate frequency, to later be digitally down converted to provide the in-phase and quadrature-phase components. Another section will present the pedestal, motors, and antenna control. Finally, calibration results of the transmitter and receiver are given.

Chapter 3 discusses signal processing of the digital I&Q complex components used to calculate the data moments. It is from these moments that reflectivity, Doppler velocity, differential reflectivity, differential propagation phase, and specific differential phase are obtained. Rain rates and clutter suppression algorithms are also presented. Significant statistical requirements of the I&Q data, to ensure estimation errors are minimized, are discussed and validation with actual data are presented.

Chapter 4 presents weather events used to validate the system with the CSU-CHILL Radar in Greeley, CO. A brief discussion will explain how the data is presented and how the radar meteorologist interprets it. This chapter will also show select observations of tropical rain events during deployment of the CSU-X25P on the western coast of Puerto Rico.

Chapter 5 will be a summary and conclusion of this Thesis. A further discussion of lessons learned from this deployment, and how they can be applied to ensure the success of the TropiNet network will be presented.
CHAPTER 2

System Description

2.1 Overview

The CSU-X25P radar was built at the CSU-CHILL Radar facility in Greeley, CO where it was tested and validated in 2010 (Figure 2.1). The radar is composed of a transceiver, data acquisition unit (DAQ), signal processor, antenna, antenna positioner, and associated computers and disk space for control, data storage and processing. Figure 2.2 shows the major components of the CSU-X25P Radar system.

![Figure 2.1: CSU-X25P validation at CSU-CHILL Radar Facility](image)

The magnetron-based transmitter, operating in the 9.41 GHz ± 30 MHz frequency range, with a maximum 25 kW peak output power, is jointly housed, in an enclosure, with the receiver, eDAQ, and processor near the antenna, above the azimuth axis of rotation. In order to accommodate the higher dual PRF bursts of pulses needed for high Doppler velocity resolution and clutter filtering, the transmitter was tuned below its maximum peak power.
Figure 2.2: CSU-X25P Radar System Block Diagram
allowing for increased duty cycle. All radar signals and control commands between the transceiver and the signal processor are made via a Gigabit Ethernet rotary joint installed through the azimuth axis [5].

The center-fed, dual linearly polarized parabolic antenna provides a minimum gain of 44 dB with a beamwidth of 1.03 degrees (in azimuth and elevation) for both channels. The antenna comes equipped with an antenna-mounted radome to help protect it from the environment. Arbitrary scanning patterns are achieved by 360 degrees of rotation in azimuth, and up to 90 degrees in elevation by the antenna positioner. The radar specifications are listed in Table 2.1.

<table>
<thead>
<tr>
<th>Table 2.1: CSU-X25P Radar Specifications</th>
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<tbody>
<tr>
<td><strong>TRANSMITTER</strong></td>
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<tr>
<td>Type</td>
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<tr>
<td>Center Frequency</td>
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<tr>
<td>Peak Power output</td>
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<td>Average Power output</td>
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<td>Pulse Width</td>
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<tr>
<td>Polarization</td>
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<tr>
<td>Max. Duty Cycle</td>
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<tr>
<td><strong>ANTENNA AND PEDESTAL</strong></td>
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<td>Type (diameter)</td>
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<tr>
<td>3-dB Beam width</td>
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<tr>
<td>Gain</td>
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<tr>
<td>Maximum Scan Rate</td>
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<tr>
<td><strong>RECEIVER</strong></td>
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<tr>
<td>Type</td>
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<tr>
<td>Dynamic Range</td>
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<td>Noise Figure</td>
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<tr>
<td><strong>DATA ACQUISITION SYSTEM</strong></td>
</tr>
<tr>
<td>Sampling Rate</td>
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<tr>
<td>Dynamic Range</td>
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Figure 2.3: CSU-X25P Transceiver Block Diagram
2.2 Transceiver

At the heart of the radar is the transceiver which is composed of a high power transmitter, a duplexer, and a receiver as seen in the block diagram of Figure 2.3. It is here that a sub-microsecond pulse is generated by the Raytheon Mark 2 high power modulator board and transmitted by a magnetron with up to 25 kW (25 W average) output power, and an operating center frequency of 9.41 GHz. The transmitter output power level is reduced by 2 dB, from it’s default maximum of 74 dB at 1 µs pulse width, to a maximum pulse length of 660 ns which allows for an increased duty cycle. Higher pulse repetition frequencies (PRF) result in transmitter waveforms that set the ability of the system to unambiguously resolve higher Doppler velocities, suppress second trip echoes, and filter unwanted ground clutter. Doppler velocity for a uniform PRF pulsed radar is found as,

\[ v_{\text{max}} = \pm \frac{\lambda PRF}{4}, \]  

and in the case for a wavelength (\( \lambda \)) of 3 cm, at 1 kHz PRF, the maximum unambiguous Doppler velocity is \( \pm 7.5 \text{ m/s} \). The transmitter waveform is modified to allow blocks of pulses at two higher PRFs, like 2.4 kHz and 1.6 kHz, thereby extending the achievable unambiguous Doppler velocity. Most meteorologically significant phenomena occur with velocities above the \( \pm 20 \text{ m/s} \) interval, therefore operating in a dual PRF mode allows the unambiguous Doppler velocity to be resolved to the interval \( \pm 38 \text{ m/s} \). The maximum unambiguous velocity, \( v_{\text{max}} \), for this scheme can be found as,

\[ v_{\text{max}} = \pm \frac{\lambda}{4(\text{PRF}_2^{-1} - \text{PRF}_1^{-1})}. \]  

The duplexer, constructed of WR90 waveguide components, equally divides the power from the magnetron into two RF channels that connect to the Horizontal and Vertical ports of the center-fed parabolic antenna. The duplexer also provides protection to the magnetron against damaging reflections from each channel, with an isolator and circulator in the path to the antenna ports. The duplexer includes, per channel, a 40 dB coupler that provides
dedicated sample paths of the transmit pulse (via an additional 30 dB pad and 20 dB coupler) which is used to estimate the transmitted frequency, amplitude, and starting phase. The duplexer also directs the received echoes from the antenna to the receiver front end, while also providing protection to the Low Noise Amplifiers (LNAs) via limiters.

The dual-channel receiver front end is composed of connectorized components, followed by a mixed signal IF-PC board that together down-convert the RF signal to a final analog intermediate frequency (IF). The receiver has three analog down-conversion stages, the first IF at 510 MHz is accomplished by mixing the received echo with a disciplined local oscillator (PLL DRO) at 8.9 GHz. The first IF is then mixed with the IFPCB 400 MHz LO, to obtain the second IF at 110 MHz. The final analog IF is obtained by down-converting the second IF with the output from a programmable Direct Digital Synthesizer (DDS) to yield an IF in the 9.75-21.25 MHz band, to then be sampled and digitized by the data acquisition system A/D converters. The digital receiver in the data acquisition unit (eDAQ) will down-convert this digitized signal to baseband by setting the numerically controlled oscillator (NCO) to the appropriate frequency. At baseband we are only left with the weather echo return, which can then be further processed to remove ground clutter and estimate meteorological moments for radial velocity ($v_d$), reflectivity ($Z$), differential reflectivity ($Z_{dr}$), differential propagation phase ($\phi_{dp}$), and specific differential phase ($K_{DP}$). These products will be discussed in more detail in Chapter 3.

The DDS output frequency is set by a software-controlled Automatic Frequency Control (AFC) loop in the eDAQ, determining the analog receiver’s final down-conversion frequency by tracking the transmitted frequency. The naturally occurring random phase of the magnetron can also be estimated, and the DDS allows for the setting of any arbitrary phase difference between the two receive channels. This can be used to mitigate second-trip echoes. The analog receiver also includes a four level phase matched fast-switching Automatic Gain Control (AGC) which allows extension of the dynamic range, thus preventing saturation of the LNAs in strong clutter environments.
2.3 Data Acquisition System

The data acquisition system of the CSU-X25P is based on the proven FPGA-based board set from Dynamic Sensing Technologies (eDAQ), designed for networked weather radar nodes [6]. The dual-channel eDAQ serves to digitize the intermediate frequency, is a digital receiver, AFC and AGC controller, real time data manager, and is also responsible for generating the triggering scheme for the transmitter. At the eDAQ front end, the two impedance matched channels include programmable gain of the H and V signals prior to each being digitized by 14-bit ADCs, at a 100 MHz sampling frequency. The sampled data signal is processed in real-time and a digital I & Q (in-phase and quadrature) detection is implemented, along with band limiting and decimation to achieve the final bandwidth. As previously mentioned, a numerically controlled oscillator (NCO), with a better than 150 dB spurious-free dynamic range (SFDR), is digitally implemented in the processor core, within the FPGA, to down-convert the incoming IF to baseband. The NCO frequency is selected based on the magnetron frequency estimate obtained from samples of the transmit pulse, collected at the beginning of each data ray. The estimated magnetron frequency is used to set the DDS output in the range of 75 MHz to 125 MHz, at steps of 12.5 MHz, to keep the incoming analog IF near 10 MHz thus ensuring sampling above Nyquist and allowing the NCO to effectively digitally down-convert to baseband.

The digital AFC loop monitors the sampled transmit frequency to guarantee the received signal is in band and not affected by frequency shifts of the magnetron output due to thermal variations and/or instantaneous changes in duty cycle, resulting from dual PRF operation. The eDAQ board also performs the critical function of making this a coherent-on-receive radar by subtracting the estimated transmit pulse phase from the received echoes. Finally, User Datagram Protocol or UDP-formatted data packages containing the filtered, decimated, coherent, baseband I & Q samples are streamed in real time over a gigabit link to a post-processing computer on the local network. UDP is a protocol used to transport data across an IP based network, and skips handshaking, which makes it faster than most other protocols during data transmission.
2.4 Pedestal, Motors/Drives, and Control

In selecting the antenna positioner for the CSU-X25P, the key requirements were that it must achieve 360° azimuthal (AZ) rotation at a minimum of 30°/s and 90° tilt in the elevation (EL) axis. Aside from meeting these requirements, the SCR-584 pedestal was selected because of its robust design, ability to operate in most adverse environments, and its modest cost and availability. The Signal Corps Radio (SCR) #584 was an S-band radar (operating frequency ≈ 3000 MHz) developed by the Radiation Laboratory at MIT during World War II [7]. These pedestals are sought after by many researchers designing both mobile and stationary weather radar systems for their proven deployability in the harshest of environments.

To improve the rotational resolution of the SCR-584, the original motors were replaced by two high power rotary brushless servo motors, for high inertia (Parker-Hannifin model MPJ0922D41KPSV), as seen in Figures 2.4 and 2.5. The servo motors are driven by Advanced Motion Controls digital servo drives (model DPRANIR03A400, for data sheet see http://www.a-m-c.com/products/dpr.html), which can be controlled using AMC DriveWare or custom position control software, via standard RS-485/232 communication protocol. AMC DriveWare is used to initialize the drives with current-loop, acceleration, and digital I/O configurations for proper operation of the motors. A custom position control moment daemon, residing on the antenna control computer is used to specify start/stop angles, scan speeds, incremental steps, and select axis of motion to achieve volume coverage patterns (VCP) for weather feature identification.

SICK-Stegmann ARS-25 SSI encoders were used as replacements to the original SCR-584 synchros, to provide feedback of the angular position of both azimuth and elevation axes. The servo drives selected did not include incremental encoder inputs, and therefore a custom circuit board (Appendix A) was designed by a CHILL staff member to interpret the encoder input before interfacing to the antenna control computer via an RS-485 port. It is important to note that the servo amp power wires, for Motor A, B, and C, must be wrapped through toroid ferrites (see Advanced Motion Controls Ferrite Recommendations, under
'Downloads & Support,’ at http://www.a-m-c.com/products/dpr.html), near the source, in order to eliminate transient electromagnetic interference onto adjacent encoder, gigabit ethernet, and other control signals. Not doing so results in intermittent and/or incorrect I & Q data packet and antenna position transmissions. This was one of the key problems solved in the development stage of this radar.

Adapter plates (Figure 2.5), machined from aluminum stock, were necessary to mate both Parker motors to the original motor mounting holes, while the motor shaft was lengthened to allow proper engagement of the planetary gears. The incremental encoders also required custom machined adapter plates for mounting into the original synchro mounts, and proper engagement of position feedback gears in the elevation and azimuth axes. It was necessary to modify the slip-ring wiring (Figure 2.6) and the terminal blocks above and below the axis of rotation, to separate the elevation axis motor power wiring from the EL encoder wiring. The pedestal was acquired as a refurbished unit and based on two units received, both arrived
with unique wiring order through the slip-ring.

A gigabit ethernet rotary joint assembly (MOOG Inc.) was installed to enable real-time high speed transfer of the I&Q data from the eDAQ (mounted behind the antenna, above the azimuth axis) to the remotely located signal processing computer. The rotary joint passes through the SCR-584 pedestal’s slip-ring center shaft via the azimuthal axis into the elevation terminal block wiring chamber, seen in Figure 2.7. This center shaft was originally designed to pass waveguide from the pedestal base up to the antenna via waveguide rotary joints in the azimuth and elevation axes. A custom mounting plate and conduit assembly were fabricated, as seen in Figure 2.8. The Ethernet slip-ring must be mounted from underneath the pedestal, and thus required it be mounted onto a stand that allowed easy access during the assembly stages.

### 2.5 Calibration

Prior to collecting meteorological data with the CSU-X25P, it was necessary to obtain different measurements to characterize the transmitter, receiver gain and dynamic range, followed by a metallic sphere calibration. These tests were conducted at the CSU-CHILL Radar Facility.
2.5.1 Transmitter Calibration

The transmitter pulse peak power was measured for both horizontal (H) and vertical (V) ports using a power meter connected to the respective duplexer antenna port through a 40 dB coupler plus an additional 30 dB attenuator to prevent damage to the power sensor head. Because the transmit pulse power out of the magnetron is expected to be around 72 dBm, a high power dummy load was connected in series to the 40 dB coupler’s through port. The measured H-polarization signal was found to be \( P_{TH} = 68.1 \text{ dBm} \), whereas the corresponding value of the V-polarization signal was \( P_{TV} = 68.7 \text{ dBm} \). To validate the measured values at the antenna ports, the duplexer losses are added to these values to obtain the magnetron output power, which was found to be as expected for both channels. The coupler, duplexer, and all other RF paths were characterized prior to final assembly. The transmitter results are summarized in Table 2.2.

A sample of the transmit pulse is recorded by the eDAQ, via the Tx sample path (40 dB
coupler, 30dB attenuator, 20dB coupler, and high speed RF switch) into the LNA to verify its temporal and spectral characteristics. This sample is taken on a pulse-to-pulse basis, by controlling the high speed switch between the transmit and receive paths, then the signal proceeds to the downconversion stages. The transmit pulse was found to have a duration of \( \tau = 450\text{ns} \) (Figure 2.9), and a corresponding transmit frequency spectrum with a 3 dB bandwidth of \( B_{Tx} = 2.01\text{MHz} \), as seen in Figure 2.10.

### 2.5.2 Receiver Calibration

The gain and dynamic range of the CSU-X25P receiver were established using an external signal source connected at the receiver H and V antenna ports. A sinusoidal signal, set at the operating center frequency \( f = 9.41\text{GHz} \), is swept in power from the receiver’s noise floor to saturation, while the digital receiver’s measured output power is recorded. Figure 2.11 shows the results. The dynamic range is obtained by determining the difference in power levels corresponding to the 1 dB compression point and noise floor. The RF-to-IF gain is
established by subtracting the input and output power levels in the linear region.

The receiver bandwidth was established by injecting a sinusoidal signal of known frequency, $f = 9.41\,GHz$, into the receiver antenna ports. The receiver downconversion frequency was then swept, while the variation in power at the digital receiver output was recorded. This resulted in a 3 dB receiver bandwidth of $B_{3dB} = 1.01\,MHz$, in good agreement with expected theoretical values.

The previous measurements indicate a mismatch between the transmitted pulse bandwidth ($B_{Tx} = 2.01\,MHz$) and the receiver bandwidth ($B_{3dB} = 1.01\,MHz$), resulting in signal power loss due to the receiver finite bandwidth. To investigate this power loss, the following procedure was followed: the measured receiver transfer function was used to filter the transmitted pulse spectrum, and the signal power difference between the filter input and output yielded a signal loss $L_r = 3.7\,dB$. 
Figure 2.9: Transmit pulse time sample
Figure 2.10: Transmit pulse frequency spectrum
2.5.3 Metallic Sphere Calibration

A sphere calibration provides an end to end calibration of the radar system. This involves tethering, or free-flying, a metal sphere of known radius, to a 300 gram weather balloon, and use it as a target in the antenna far field. Figure 2.12 shows the battery-powered winch that was used to hold the balloon at a relatively fixed height of about 500 ft. above the ground. Figure 2.13 shows the balloon and metal sphere at approximately 1 km away from the CSU-X25P. Time series data was collected while scanning around the target and was used to obtain the backscattered signal power. The measured power was then compared to the expected theoretical power. The expected theoretical power at the digital receiver
output is obtained by the following:

\[ P_r = P_t + 2G_a + 10\log\left[\frac{\pi a^2 \lambda^2}{(4\pi)^3 r^4}\right] - L_r + G_{dBm2dBu} \] (2.3)

where \( P_t \) is the transmitted power in dBm, \( G_a \) is the antenna gain in dB, \( a \) is the radius of the metal sphere in m, \( \lambda \) is the radar wavelength, \( r \) is the radial target range from the radar, \( L_r \) is the radar receiver finite bandwidth loss, and \( G_{dBm2dBu} \) is the radar receiver analog plus digital gain. Table 2.3 summarizes the parameters used and the obtained results.

<table>
<thead>
<tr>
<th>( P_r ) [dBu]</th>
<th>( P_t ) [dBm]</th>
<th>( G_a ) [dB]</th>
<th>( a ) [in]</th>
<th>( \lambda ) [m]</th>
<th>( r ) [m]</th>
<th>( L_r ) [dB]</th>
<th>( G_{dBm2dBu} )</th>
</tr>
</thead>
<tbody>
<tr>
<td>78.1</td>
<td>68</td>
<td>44</td>
<td>5</td>
<td>0.03</td>
<td>1020</td>
<td>3.7</td>
<td>122</td>
</tr>
</tbody>
</table>

The time series data collected during the experiment was processed to first obtain the target returns as the radar scanned in azimuth. Once a return signal was identified, a gaussian fit (corresponding to the two-way antenna pattern azimuth) was used to determine its power and position in azimuth. This was repeated for the adjacent range gates to obtain a gaussian fit in range. The maximum values were selected from the gaussian fit in range at multiple elevations to obtain a final gaussian fit in elevation from which to obtain the measured metal sphere target backscatter.
Figure 2.12: Launching balloon and tethered metal sphere using electric winch
Figure 2.13: Balloon and metal sphere target down-range, 1 km from CSU-X25P
CHAPTER 3

Signal Processing

3.1 Overview

After the correct analog and digital downconversion stages of the IF signal are achieved, the
decimated and band limited received echo signal is phase corrected relative to the transmit
pulse phase. This is a standard coherent-on-receive technique used to obtain the wanted echo
signal information, such as phase of the scattered wave, the complex signal amplitude, and
Doppler shift relative to the radar. The received signal due to a single scatterer is a scaled
replica of the transmitted waveform, $U_{tr}(t)$ shifted in time by its delay in range ($\tau_s = 2r/c$)
and in frequency by the Doppler shift ($f_D = -2v/\lambda$). At the antenna port, the received
voltage (i.e. complex envelope of received signal) for a spherical particle can be expressed
as [8],

$$V_r(t) = \frac{\lambda G \sqrt{P_t}}{4\pi r^2}(S)e^{-j2k_0 r}U_{tr}(t - \tau_s). \quad (3.1)$$

where $S$ is the scattering matrix of the particle, $P_t$ is the transmitted power, $G$ is the antenna
gain, $r$ is the particle’s range, and the wave number $k_0 = 2\pi/\lambda = 2\pi f_0/c$.

In the case of scattering from precipitation, large numbers of hydrometeors will occupy a
resolution volume defined by the range gate at $r + \Delta r_{min}$, and the minimum 3dB beamwidths
of the antenna, in azimuth and elevation ($\phi_{3dB}$ and $\theta_{3dB}$). The hydrometeors can extend over
a large range with varying scattering amplitudes and moving at different velocities, where the
received voltage is the resultant of contributions from the scatterers. The received voltage
$V_r(t)$, at a particular range ($r = c\tau_s/2$), can be expressed as a discrete sum from the individual
contributions in the resolution volume as,

$$V_r(t) = \sum_k A_k(\tau_k; t)e^{-j2f_0 \tau_k}U_{tr}(t - \tau_k). \quad (3.2)$$
where $A_k$ is the scattering amplitude of the $k$th hydrometeor. For a given $t$ the receive voltage can be represented as the sum of these elemental phasors, and when considering $V_r(t + \Delta t)$, the particles are time-varying. This movement causes change in their location $r_k$ and phase $\theta_k = (4\pi/\lambda)r_k$, where $-\pi \leq \theta_k \leq \pi$, and by estimating the received power and phase shift relative to the transmit pulse the various moments necessary for meteorological analysis are obtained. These moments are calculated from the I&Q components of the complex envelope, to include the effective reflectivity ($Z$) and Doppler velocity ($v_d$ in m/s), differential reflectivity ($Z_{dr}$), differential propagation phase ($\phi_{DP}$), and specific differential phase ($K_{DP}$).

Due to its transmitter configuration, the CSU-X25P radar test node only operates with simultaneous transmission of horizontal and vertical polarization states, but in addition to copolar return power, each receiver antenna port experiences cross-coupling from the other polarization (i.e., cross-polar power). This mode of operation is know as the "hybrid mode," and experiences perturbations in the retrieved copolar radar parameters previously mentioned. Care was taken to mitigate radar parameter biases, by limiting and addressing system polarization errors, like those due to antenna properties and orientation of hydrometeors, as presented in [9], [10], and [11]. The calculation of data products throughout this paper are made under the general assumption of a hybrid mode of operation.

### 3.2 Digital I & Q Detection

Digital detection of the in-phase and quadrature (I&Q) components of the sampled receiver echoes is the most important step in radar signal processing in order to calculate the Doppler spectral moments. At the output of the analog receiver, the radar signal can be represented by amplitude and phase modulation as follows,

$$x(t) = a(t) \cos(2\pi f_0 t + \theta(t))$$

$$= a(t) \cos \theta(t) \cos 2\pi f_0 t - a(t) \sin \theta(t) \sin 2\pi f_0 t$$

$$= I(t) \cos 2\pi f_0 t - Q(t) \sin 2\pi f_0 t$$
where \( f_0 \) is the carrier frequency, \( \theta \) is the phase modulation, and \( a(t) \) is the amplitude modulation. The I&Q components are extracted from the complex signal \( s(t) \) after downconversion to baseband, and as seen below, are the orthogonal basis components of the receiver voltage \( V_r(t) \) as,

\[
s(t) = V_r(t) e^{j2\pi f_0 t} = a(t) e^{j\theta(t)} e^{j2\pi f_0 t} = [I(t) + jQ(t)] e^{j2\pi f_0 t}
\]

The CSU-X25P digital receiver converts the analog IF complex signal into a discrete sampled version which will finally be digitally downconverted to baseband with a numerically controlled oscillator (NCO). The discretized baseband complex signal is,

\[
V_r[n] = s[n] = I[n] + jQ[n] = a[n] e^{j\theta[n]}
\]

The received power is obtained from the complex signal as,

\[
P_r[n] = V_r[n] V_r^*[n] = |V_r[n]|^2 = I^2[n] + Q^2[n]
\]

and is typically represented as the mean power \( \overline{P_r}[k] \) of \( N \) randomly distributed samples with random scattering amplitudes, at a particular range in the \( k \)-th range gate. For this random process, the mean power is simply the variance of \( V_r(t) \).

The real and imaginary components, I and Q, carry information about the stochastic signal and have to satisfy the following properties as presented in [8]:

1. The I & Q components for the same time, are uncorrelated

\[
cov[I(t), Q(t)] = 0
\]
2. The means of I & Q are zero,

\[ E[I(t)] = E[Q(t)] = 0 \]  \hspace{1cm} (3.14)

3. The I & Q components have the same variance,

\[ E[I^2(t)] = E[Q^2(t)] = \sigma^2 \]  \hspace{1cm} (3.15)

4. The real and imaginary parts have the same autocorrelation function,

\[ E[I(t_1)I(t_2)] = E[Q(t_1)Q(t_2)] = \sigma^2 \rho_0(t_2 - t_1) \]  \hspace{1cm} (3.16)
5. The I & Q components at two different times, are correlated

\[ E[I(t_1)Q(t_2)] = -E[I(t_2)Q(t_1)] = \sigma^2 \alpha_0(t_2 - t_1) \] (3.17)

Time series data samples collected by the CSU-X25P radar in precipitation were analyzed to illustrate the properties of a stochastic process. Figure 3.1 shows the histograms of I and Q which are similar to Gaussian density functions, with zero mean. Also, the similarity in their distribution suggests they have the same variance. Figure 3.2 shows that the scattergram of I vs. Q is spread equally, without any preferential direction, indicating they are independent. Figure 3.3 shows the histogram of the power samples, \((I^2 + Q^2)\), having a near linearly
Figure 3.3: Log histogram of \((I^2 + Q^2)\). The linear decrease indicates \(I^2 + Q^2\) is exponentially distributed.

decreasing form indicating an exponential distribution, as expected for a probability density function of signal power given by,

\[
f_P(P) = \frac{1}{2\sigma^2} e^{-P/2\sigma^2}, \quad P > 0.
\]  

(3.18)

It is evident from the previous figures that the sampled In-phase and Quadrature phase components of a typical rain event satisfy the properties of a valid stochastic process. This validation will ensure accurate estimation of the Doppler data moments to be calculated.
### 3.3 Reflectivity (Z)

Reflectivity, $\eta$, is the general radar term describing the volumetric backscatter cross-section per unit volume, and is calculated from the received complex signal envelope, derived from the discrete I&Q components at the output of the digital receiver. The reflectivity from meteorological backscatterers can be written as, $\eta = (\pi^5/\lambda^4)|k_w|^2Z_e$. The equivalent reflectivity factor $Z_e$ can be estimated from the received power, $P$, using the weather radar equation [12], shown here in units commonly used by radar meteorologists,

$$P(mW) = \frac{\pi^{5}10^{-17}P_t(W)G^2G_r\tau(\mu s)\theta_{3dB}^2(deg)|k_w|^2Z_e(mm^6/m^3)}{6.75x2^{14}\ln2\ r^2(km)\lambda^2(cm)^2l_r}$$

(3.19)

Receive power at range $r$, $P(r)$ is measured in $mW$, transmit power $P_t$ in Watts, antenna and receiver gains, $G$ and $G_r$ are dimensionless, transmit pulse width $\tau$ in micro seconds ($\mu s$), antenna 3dB-beamwidth (H and V) $\theta_{3dB}$ in degrees, the dielectric factor of water $|k_w|^2 = |(\epsilon_r - 1)/(\epsilon_r + 2)|^2$ where $\epsilon_r$ is the complex relative permittivity of water, the range $r$ is in kilometers, wavelength $\lambda$ is in centimeters, and attenuation and receiver losses, $l$ and $l_r$ are dimensionless. The equivalent reflectivity factor $Z_e$ is simply referred to, by radar meteorologists, as reflectivity for H,V polarizations at a range $r$, $Z_{h,v}(r)$ measured as $mm^6m^{-3}$ in $dBZ$.

By rearranging the weather radar equation the reflectivity, at range $r$, for a given power measured at the H and V channel outputs of the digital receiver, can be expressed as,

$$Z_{h,v}(r) = \left(\frac{6.75x2^{14}\ln2}{\pi^{5}10^{-17}}\right)\left(\frac{l_r\ G^2}{\tau\ \theta_{3dB}^2\ |k_w|^2}\right)\left(\frac{\lambda^2(f_{DDS},f_{NCO})\ l^2}{P_t\ G_r(f_{DDS},f_{NCO})}\right)^2\ P_{h,v}(r)$$

(3.20)

where the quantities within the parentheses represent parts of the radar system constant $C = C_kC_sC_d$, $C_k$ is a numerical constant, $C_s$ is dependent on static system hardware parameters including the antenna and receiver gains, and $C_d$ is dependent on dynamic hardware parameters such as sampled transmit frequency and associated automatic frequency control parameters (i.e., $f_{DDS}, f_{NCO}$) set during the coherent-on-receive step. For example, the transmitted wavelength perturbations are due to the random variations in frequency output.
from the magnetron, and calculated as \( \lambda(f_{DDS}, f_{NCO}) = c/(f_{DDS} + f_{NCO} + f_{STALO}) \), where

the STALO frequency is fixed at 8900 MHz, and the DDS and NCO frequencies vary depending on the estimated frequency of the sampled transmit pulse. It is widely accepted in the radar meteorology community to represent \( Z \) using a logarithmic scale and written as,

\[
Z_{h,v}(r) = 10 \log[C \ P_{h,v}(r) \ r^2] \quad \text{[dBZ]} \quad (3.21)
\]

and will henceforth be referred to in dBZ.

### 3.4 Doppler Velocity (\( v_d \))

Meteorological events include precipitation with varying phases including rain, hail, snow, ice, and everything in between which can be detected by its equivalent reflectivity factor. But, in order to better understand the event from initiation to extinction, and to provide reasonable forecasting, the Doppler velocity \( v_d \) is an essential observation. The CSU-X25P radar node employs the coherent-on-receive technique, by means of detecting the transmit pulse phase of the magnetron, which can later be removed thus only leaving Doppler shifts introduced by the weather echoes.

In a pulsed radar, the received voltage \( V_r(t) \) is due to backscattering by particles in the resolution volume at a particular range-time \( t = \tau_s = 2r/c \). Transmit pulses are separated in time by a period \( PRT = T_s \), and thus for a particular range-time the voltage can be seen as a sequence of temporal samples at \( V_r(\tau_s) \), \( V_r(\tau_s + T_s) \),..., \( V_r(\tau_s + mT_s) \). Due to the random nature of the time-varying properties of the hydrometeors, sample-time fluctuations in the received voltages can be treated as a complex stochastic process. Assuming the average velocity \( \overline{v_d} \) and spectral width \( \sigma_v \) define a Gaussian-shaped Doppler spectrum, the sample-time autocorrelation function is given by [8],

\[
R[n] = S_0 \exp \left( -\frac{8\pi^2 \sigma_v^2 n^2 T_s^2}{\lambda^2} \right) \exp \left( -j \frac{4\pi \overline{v_d}nT_s}{\lambda} \right); \quad n = 1, 2, ..., N \quad (3.22)
\]

where \( S_0 = |R[0]| \) is the magnitude of the autocorrelation at lag 0, and the correlation
coefficient can be expressed as $\rho[n] = R[n]/R[0]$. From the above relationships, we can obtain the power spectral density (i.e., Doppler spectrum) from the autocorrelation function by,

$$ S(\Omega) = \sum_{n=-\infty}^{\infty} R(n)e^{-j\Omega n}, \quad (3.23) $$

where $S(\Omega)$ is the discrete time Fourier transform of the autocorrelation function, with Doppler frequency components in the interval (-\(\pi\), \(\pi\)). The Doppler velocity spectrum is thereby given as,

$$ S(v) = \frac{S_0}{\sigma_v\sqrt{2\pi}}exp\left(-\frac{(v - \bar{v})^2}{2\sigma_v^2}\right), \quad (3.24) $$

where the mean velocity is $\bar{v} = -\frac{\lambda}{4\pi T_s}\Omega$ and the standard deviation of velocity is $\sigma_v = \frac{\lambda}{4\pi T_s}\sigma_\Omega$.

The maximum radial velocity that can be observed without aliasing is $v_{max} = -\lambda f_{max}/2 = -\lambda/4T_s$, where $f_{max} = 1/2T_s$ and the maximum unambiguous range is defined as $r_{max} = cT_s/2$. Their product $v_{max}r_{max} = c\lambda/8$, fixed by the radar wavelength, shows that an increase in maximum range comes at the cost of a decrease in maximum unambiguous velocity, and vice versa. This is one of the fundamental relationships taken into account by radar engineers when designing a pulsed Doppler radar system.

In [13] and [14], the authors present a parametric estimation algorithm for X-band pulsed Doppler radars that provides improved performance in estimation of spectral moments in the presence of strong clutter signals, with a lower number of samples. With shorter dwell times, increased scan rates can be achieved without losing the ability to achieve higher Nyquist velocities up to 38 m/s, using dual-PRF waveforms. The authors also presented a table of optimal dual-PRF waveforms for various scan speeds (deg/s), PRF combinations ($PRF_1$ and $PRF_2$), 3/4 stagger ratio of pulse blocks (i.e., $N_1 = 40$, $N_2 = 54$), detectable Doppler velocities and Azimuthal resolution. In practice, the mean Doppler velocity for each block of pulses is computed and then combined for an extended unambiguous interval as, $\mp\lambda/4(T_{s1} - T_{s2})$. The waveforms thereby presented were used as a guideline to improve operation of the CSU-X25P radar node.
3.5 Differential Reflectivity ($Z_{dr}$)

The CSU-X25P radar node is a dual polarized radar system operating in a hybrid mode, with a transceiver that simultaneously transmits, then receives both Horizontally and Vertically oriented signal waveforms. There are many advantages polarimetric radars have, mainly their ability to obtain two different "looks" from the same resolution volume, in range. Because the transceiver has independent dual-linear channels, estimated powers for the H- and V-channels can be found as,

$$\hat{P}_{h,v} = \frac{1}{N} \sum_{n=1}^{N} |V_{h,v}[n]|^2,$$

for N samples. The average axis ratio of particles, such as precipitation in the resolution volume can be estimated based on the differential reflectivity $\hat{Z}_{dr}$, or estimated ratio of reflected power from the H- and V-polarizations as,

$$\hat{Z}_{dr} = 10 \log \left( \frac{\hat{P}_h}{\hat{P}_v} \right) \quad \text{[in dB]},$$

(3.26)

From the differential reflectivity estimate, radar meteorologists can measure rainfall and discriminate hydrometeor types as described in [8, 12]. As $Z_{dr}$ increases, in the absence of severe attenuation, it can be inferred that the horizontal radar cross section of raindrops is greater than the vertical, which is indicative of larger raindrops reaching hydrostatic equilibrium. Once a raindrop reaches terminal velocity, the larger its fluid volume, the more it spreads horizontally, thus yielding values of $Z_{dr} > 2$ dB. For very small raindrops, with lower reflectivity values, this ratio will be closer to unity (or 0 dB). It must be noted though, that hail will also have differential reflectivity values near 0 dB, but will also have large reflectivity values, typically above 50 dBZ.

The differential reflectivity measurement is affected by imbalance in boresight antenna gains and orthomode transducer channel differences in orthogonal signal propagation via the feedhorn. To account for this, a procedure presented in [15] to calibrate $Z_{dr}$ using properties of the rain medium was followed. They suggest collecting rain data with the antenna
pointed at an elevation of 90°, while rotating 360° in azimuth. When viewed at vertical incidence, raindrops are azimuthally symmetric with a $Z_{dr} = 0$ dB, and as the polarization basis rotates it is periodic with the azimuth angle. The measured $Z_{dr}$ is also periodic, but independent of the scattering, and for data in the far field (i.e., range $\geq 2D^2/\lambda$ or $\geq 361$ m), the mean differential reflectivity represents the gain imbalance in the two channels. The $Z_{dr}$ correction factor for the CSU-X25P was found to be 0.18 dB as seen in Figure 3.4. It’s important to note that when propagating through heavy rain the cumulative effects of differential attenuation will cause $Z_{dr}$ to become negative, and will be apparent in the loss of horizontal reflectivity ($Z_h$) beyond the heaviest rain. Attenuation correction methods using the differential propagation phase ($\phi_{DP}$) will be discussed later.

Figure 3.4: CSU-X25P $Z_{dr}$ Calibration
3.6 Differential Propagation Phase ($\phi_{DP}$) and Specific Differential Phase ($K_{DP}$)

The differential propagation phase measurement is obtained by taking the difference in phase between the horizontal and vertical polarizations, from simultaneous data profiles in the same range gate. The individual channel phase is obtained after phase-correction (i.e., removal of initial transmit pulse phase) to ensure it is due only to the received echoes. If we consider the simultaneous (i.e., lag 0) orthogonal received signal samples $V_h[n]$ and $V_v[n]$, in a hybrid transmission mode, the covariance is

$$\hat{R}_{h,v}[0] = \frac{1}{N} \sum_{n=1}^{N} V_v[n] V_h^*[n]. \quad (3.27)$$

From this the estimated correlation coefficient $\rho_{h,v}$ can be calculated as,

$$\hat{\rho}_{h,v} = \frac{\hat{R}_{h,v}[0]}{(\hat{P}_h \hat{P}_v)^{1/2}} \quad (3.28)$$

and subsequently its magnitude $|\hat{\rho}_{h,v}|$ and differential phase $\phi_{DP}$ as,

$$|\hat{\rho}_{h,v}| = \frac{|\hat{R}_{h,v}[0]|}{(\hat{P}_h \hat{P}_v)^{1/2}} \quad \phi_{DP} = arg(\hat{R}_{v,h}[0]). \quad (3.29)$$

The differential propagation phase is always positive in rain and is a monotonically increasing function of range. Unlike reflectivity $Z$ which is adversely affected by attenuation as the electric field propagates through rain, $\phi_{DP}$ does not experience extinction. An important property of differential phase is its independence from system gain calibration.

The magnitude of the correlation coefficient can be used to filter out unwanted reflectivity when displaying the weather echoes in a Planned Position Indicator (PPI) plot, where a typical value used is $|\rho_{h,v}| \geq 0.98$. A high correlation coefficient will also ensure a higher accuracy in the estimate of differential propagation phase.

Specific differential phase $K_{DP}$ can be estimated as the range derivative of $\phi_{DP}$, and is
valid between ranges \( r_1, r_2 \). The one-way finite difference estimate of \( K_{DP} \) is written as,

\[
\hat{K}_{DP} = \frac{\phi_{DP}(r_2) - \phi_{DP}(r_1)}{2(r_2 - r_1)} \quad \text{in } ^\circ \text{km}^{-1} \tag{3.30}
\]

while care should be taken in selecting the range interval to avoid sudden fluctuations in differential phase due to noise or other factors. Methods to improve the accuracy of the \( K_{DP} \) estimate by smoothing of, or applying a linear / non-linear regression fit to, the range profile of \( \phi_{DP} \) can be found in [8].

A rain rate estimate based on \( K_{DP} \) can be written in the form presented by (Sachidananda and Zrnic, 1986) as,

\[
R(K_{DP}) = aK_{DP}^b \quad \text{in } \text{mm hr}^{-1} \tag{3.31}
\]

where the values for coefficient \( a \) and exponent \( b \) are frequency dependent and found empirically. In [16] the authors found that for \( R < 30 \text{ mm hr}^{-1} \), with \( \lambda = 3.2 \text{ cm} \), assuming an oblate equilibrium drop shape aspect ratio, the lognormal distribution coefficient \( a = 12.1 \), with an exponent of \( b = 0.85 \). The rain rate estimate accuracy is also dependent on the range interval selected for \( K_{DP} \), and should be taken from the lowest elevation angle possible.

### 3.7 Clutter Suppression

For any pulsed radar, the echoes will most likely contain strong returns from unwanted ground targets such as buildings, trees, or mountains. These returns, typically in the 40 to 60 dBZ range are backscatter from the main beam at boresight or from a sidelobe of the antenna, and are referred to as ground clutter. These non-meteorological targets bias the reflectivity and velocity estimates, and must be mitigated in order to extract the observed meteorological phenomena. The clutter contaminated weather echo can be modeled with a spectral density given by,

\[
S_m(\mu) = \frac{S_c}{\sqrt{2\pi \sigma_c^2}} \exp\left(\frac{v_k^2}{2\sigma_v^2}\right) + \frac{S_0}{\sqrt{2\pi \sigma_v^2}} \exp\left(-\frac{(v_k - v_m)^2}{2\sigma_v^2}\right) + N \tag{3.32}
\]
where \( \mathbf{\mu} = [S_c \sigma_c S_0 v_m \sigma_v N]^T \), subscript "c" refers to clutter parameters and \( N \) refers to the noise power density.

The CSU-X25P uses a real-time spectral fit clutter filter as implemented in [13, 14], where the zero-velocity clutter spectral coefficients are notched with a spectral clipper, using a Gaussian model for the clutter spectral density. A Gaussian weather spectral density is recursively fit to the remaining points and the notched spectral coefficients are interpolated with the model. The filtered signal will experience up to a 50 dB clutter suppression ratio, but unfortunately may lose weather echoes with low radial velocity.
CHAPTER 4

System Validation with the CSU-CHILL Radar and Deployment on the Western Coast of Puerto Rico

4.1 Description of Validation with CSU-CHILL Radar

The CSU-X25P radar test node was assembled and tested on site at the CSU-CHILL Radar Facility in Greeley, CO, so it was natural that it would be validated with this firmly established research radar. During the validation period, the CSU-CHILL radar was a single frequency S-Band, high power, dual polarized, Doppler weather radar, and is well recognized among researchers in this community. It has since gone through major upgrades where it was recently modified to be a dual-frequency (S- and X-band) dual-polarization Doppler weather radar.

For the validation, CSU-X25P and CSU-CHILL simultaneously collected data during convective storms in the fall of 2010. Care was taken to coordinate RHI and PPI scans of the two radars during these events, to maximize the amount of data available for analysis. The CSU-X25P radar was secured to a concrete pad at ground level, and due to blockage from structures on the property, it only had $180^\circ$ of unobstructed views to the North, between $270^\circ \leq \text{AZ} \leq 90^\circ$. Minor sectors available to the south were used for comparison in setting a benchmark.

A typical coordinated scanning effort occurred while the CHILL radar was collecting precipitation data for other research projects, typically focused over rain gauges in the Denver metro area, to the south. The X25P radar took advantage of the ongoing operations by requesting to be included in the scan strategies planned for the particular experiment, whenever available. It was during these operations that the author gained firsthand experience on effectively operating a dual polarization, Doppler weather radar. This knowledge proved to be invaluable during operational training provided to UPRM personnel.
4.2 Coordinated Observations

On August 9, 2010, a convective storm was observed near Greeley, CO and coordinated scans were performed between CSU-X25P and CSU-CHILL radars. The coordination was done manually, while operators were both in the CHILL radar user van, sharing scan strategies. For reference in Figure 4.1, the one degree beamwidths of the transmit signals, for X25P and CHILL, are plotted versus range for elevation scan angles of 2° and 0.8°, respectively. The difference in elevation scan angles is due the location of the X25P near the ground, which has increased ground clutter in the lowest angles near the radar. The CHILL reflector is mounted on a pedestal, about 10 m above the ground, and experiences less interference at the lowest scan angles. In Figure 4.2(a) & (b), the coordinated PPI scans show good agreement in the first 22 km of the X25P, but due to higher reflectivites beyond that range, as seen in the CHILL plot, severe attenuation is observed for X-band. Note that although the elevation angles and scan start times are not perfectly synchronized there is good agreement in increasing differential phase, $\Phi_{DP}$, as a function of radial range, as seen in (c) & (d). This is further highlighted in (e) & (f) by the $\Phi_{DP}$ range profiles, at an azimuth of 51.5°. It is important to note that as the transmit pulse propagates through oriented hydrometeors, the differential propagation phase, between the H and V receive echo phases, is inversely proportional to wavelength. Therefore, the X-band accumulation between the ranges of 10 km to 20 km is about three times that seen by the CHILL S-band radar. The noisy $\Phi_{DP}$ seen in both figures is a result of inaccurate estimation of differential phase in regions with lower cross-correlation coefficient (i.e., $\rho_{h,v} \leq 0.7$). This is more prevalent for a 3 cm wavelength where signal extinction occurs in horizontal polarization, before affecting the vertically oriented return, in heavy rain with large oblate raindrops.

A comparison was made between attenuation correction algorithms for reflectivity and differential reflectivity as presented by [17], and an algorithm for estimation of specific differential attenuation presented in [18]. The X25P implements the former in real-time, while a C code implementing the latter algorithm, known as DROPS, was used to obtain plots of the corrected reflectivity (CdBZ) and corrected differential reflectivity (CZdr), shown here.
Figure 4.1: CSU-X25P vs. CHILL range profiles of 1° beamwidth. X25P centered at EL=2°, with starting height of 2 m. CHILL centered at EL=0.8° with starting height of about 10 m above ground level (AGL).
Figure 4.2: X25P ($EL = 2^\circ$) and CHILL ($EL = 0.8^\circ$) PPI scans on 2010-08-09
in Figures 4.3(a)&(b). In both figures the measured and DROPS corrected range profiles were plotted together, as the best attenuation correction was obtained from the DROPS algorithm. The DROPS attenuation correction method was shown to provide up to 31.5 dB improvement for heavier rain, as seen at about 20 km range. The advantage of the Liu et al. algorithm is that it is implemented in real time, where DROPS is done in post processing. A direct result of attenuation correction, is a noticeable improvement (up to 3dB) in the negative bias in Zdr, becoming positive or near zero, due to the increase in horizontal reflectivity, $Z_h$ after attenuation correction.

In Figure 4.4(a), the intensity plot of corrected Zdr vs. $Z_h$ displays agreement with the expected relationship of drops with increasing oblate shape in heavier rainfall. The more oblate hydrometeors will have radar cross-sections with larger horizontal reflectivity, than for the vertical orientation. Here it is evident that in heavier rain ($Z_h \geq 20dB$), an almost linearly increasing differential reflectivity is obtained. For lighter rain ($Z_h \leq 20dB$), differential reflectivity is near zero, indicating more spherical hydrometeors with more similar $Z_h$ and $Z_v$ reflectivities. The intensity plot for specific differential phase, $K_{DP}$, versus corrected reflectivity in Figure 4.4(b), shows good agreement with the theoretical, and validated, $K_{DP} - Z_h$ relationship [19, 20].

### 4.3 Selected Observations

In Figure 4.5, the RHI scan at an azimuth of $217^\circ$ shows good agreement between CSU-X25P and CSU-CHILL reflectivities and Doppler velocity. Due to ground clutter spectral filtering, the CSU-X25P inadvertently removes some meteorological data at the zero-velocity transition. There is good agreement in both radars, with respect to the location of the bright band at an altitude of 3 km, where the cross-correlation coefficient, $\rho_{hv}$, shows a band of lower values due to the decorrelation effect of the melting layer. However, the lower $\rho_{hv}$ values of X25P at farther ranges are attributed to limited signal-to-noise ratio.

Figure 4.6 shows an RHI scan through a tropical rain event in June 2011, during the X25P deployment in Mayaguez, Puerto Rico, with a visible melting layer at about 4 km
Figure 4.3: X25P Measured vs. DROPS-Corrected $Z_h$ and Zdr Range Profiles
Figure 4.4: CSU-X25P Corrected Reflectivity vs. (a) CZdr and (b) $K_{DP}$.

Corrected Reflectivity vs. Corrected ZDR (\# of occurrences)

CSU-X25P Corrected Reflectivity vs. Corrected ZDR (\# of occurrences)

CSU-X25P-20100809-221024 UTC

(a) Corrected Z vs Zdr

Corrected Reflectivity vs. Specific Differential Phase KDP (\# of occurrences)

CSU-X25P-20100809-221024 UTC

(b) Corrected Z vs $K_{DP}$
Figure 4.5: CSU-X25P and CHILL RHI scans at AZ=217° on 2010-08-09 (UTC). Images created using CSU VCHILL.
above ground level (AGL). In (a) & (b) uncorrected reflectivity and differential reflectivity show signs of attenuation, highlighted by the negative bias in $Z_{DR}$. This is confirmed in (c) with a phase accumulation in $\Phi_{DP}$ of about $50^\circ$, through the region down range from from the heavier precipitation (at 15 km away from the radar and 4 km AGL). Considering the copolar cross-correlation coefficient in (d), we can see the decreasing $\rho_{hv}$ due to the de-correlating effect of the melting layer, and also negative effect due to low SNR with increasing range. Using the reflectivity and differential reflectivity correction algorithms presented by Liu et al., we can obtain a more accurate estimate of these meteorological parameters, as seen in (e) & (f). Notice the improvement of the negative bias in $Z_{DR}$. An interesting observation regarding the location of the melting layer on the western coast of Puerto Rico is its height above ground level, with respect to that seen in Greeley, Colorado. The bright band was typically observed to be about 4.5 km AGL in this Tropical region, where a 2-3 km AGL height is typical in Colorado during the same season. Further investigation of the melting layer, and specific attenuation is needed to better understand this tropical environment and its effect on radar operations at X-band frequencies.

4.4 Deployment of CSU-X25P Radar to the Western Coast of Puerto Rico

Upon the agreement and validation with the CSU-CHILL radar, the CSU-X25P node was now ready for its first deployment. Through a collaborative effort between CSU and UPRM, the design of a three radar network, known as TropiNet [4], was underway with potential site locations and infrastructure being identified. The site selected to host Puerto Rico’s first polarimetric X-band Doppler weather radar test node was the top of the R & D Building on the UPRM campus, geographically located at lat 18.2112°, lon −67.1368°, as seen in Figure 4.7.

The UPRM personnel procured and installed the twenty foot tower (Figure 4.8) and ensured the availability of adequate infrastructure such as power and network connectivity. Once the tower and infrastructure were in place, in the fall of 2010, the radar system was
Figure 4.6: Tropical rain event observed by X25P. RHI scan on June 8, 2011 06:25:37 (UTC), AZ= 288°. Images created using CSU VCHILL.
disassembled at the CSU-CHILL Radar facility and prepared for shipment to Mayaguez, Puerto Rico, as seen in Figures 4.9. The shipping of the radar system was coordinated such that its departure from Colorado avoided the harshest of winter conditions, as well as its arrival coinciding with Puerto Rico’s dry season.

In mid January of 2011, the radar system was delivered to the R & D building on the UPRM campus. The author played a key role in disassembly and the assembly, which is an important aspect of the radar, that it should hold its system integrity in disassembly and assembly. Figure 4.10 captures the installation of the antenna control server, motors,
pedestal, and 8 ft. parabolic antenna. Once the antenna and pedestal were secured to the tower, the azimuth and elevation motors, servo drives, and their associated cabling were installed. With this task completed, line power, control signaling, and Gigabit Ethernet cabling were also connected to the antenna control server enclosure and to the pedestal base riser.

With the radar transceiver, antenna, pedestal, motors, and antenna control server installed on the tower, the signal processing computer and mass storage device were housed in an environmentally controlled room inside the R & D building. A fiber optic link was made, between this computer and a gigabit Ethernet switch in the antenna control enclosure. An external gateway to the internet was established on the processing computer’s Eth0 port. A solar calibration was done to account for any errors in the pedestal position control.

The collaboration between CSU and UPRM also served as a training opportunity that would enable the staff and students to locally operate and maintain the radar node. The
author supervised and trained several UPRM graduate students about installation of a X-band dual polarization Doppler radar system, as seen in Figure 4.11. Additional support from CSU, via email and cellphone, was made possible by remote internet access to the radar node servers to assess system status, to restart processes, or identify potential root cause hardware failures. A key feature of this radar node is that it can be remotely operated, and as such, most of the collected data to be presented later in this chapter was obtained by the author, operating the radar from Colorado.

To assist in remote operations, a web camera was installed to provide a real-time visualization of the antenna. Having this view allows for safe remote operations, ensuring the tower platform is void of personnel and any obstructions. A radar control GUI is used to operate the radar (and can also be operated using scripts), while the GTK-based CSU-CHILL Real-time Display was adapted to display the CSU-X25P moment data, overlaid on a map of Puerto Rico. A web-based weather radar data visualization tool was developed by UPRM to simultaneously display moment data from NEXRAD (TJUA), CSU-X25P, and the Off-the-Grid radar nodes. Figure 4.12 shows a typical screen from the remote radar control server, with the Radar Control GUI, CSU-CHILL Real-time displays, a live feed from the video camera, and the web-based visualization tool.

It’s been the author’s experience that during CSU-X25P operations, it is beneficial to
Figure 4.10: Raising and mounting the pedestal, antenna, equipment onto building
have access to NEXRAD or Terminal Doppler Weather Radar (TDWR) data. The advantage of having the longer (radial) range reflectivity and Doppler velocity PPIs (NEXRAD = 230 km; TDWR = 90 km) from these radars provides the operator ample time to plan future scans in the lowest altitudes of incoming weather, using the faster scanning, shorter range, but higher resolution X-band radar. For example, based on the reflectivity from the TJUA radar in Figure 4.12b, and assuming easterly winds, a sector volumetric scan pattern can be planned between $90^\circ < AZ < 215^\circ$.

4.5 First Observations in a Tropical Environment

One of the first interesting weather events captured occurred on April 26, 2011 when low velocity circulation ($\pm 8 m/s$) was observed just off the northwest coast, near Rincon, PR, as presented in [21]. In Figure 4.13, a velocity couplet indicating counterclockwise circulation was detected at approximately 25 km from the radar node, centered around azimuth $336^\circ$, shown here at an elevation of $3^\circ$. Waterspouts, having similar signatures, are commonly seen along the West coast of Puerto Rico and have been known to make landfall, damaging property and causing injuries, in the immediate vicinity. The NWS Weather Forecast Office website (http://www.crh.noaa.gov/apx/science/waterspouts/waterspouts.php) defines two waterspout types: tornadic and fair weather. Tornadic waterspouts develop downward from a thunderstorm like tornadoes, and can form over water, or move from land onto water.
Figure 4.12: CSU-X25P Radar Control and Visualization Tools

(a) Control GUI, Real-time Display Control, and Live Video

(b) Real-time Display, UPRM web tool displaying NEXRAD data (background)
Fair weather waterspouts form in light wind conditions and develop on the warm water surface and work their way up toward the much colder flat base of developing cumulus clouds. Figure 4.14 shows a waterspout that formed off the coast of Mayaguez, in September of 2005. For the event of April 26th, ground validation of a water spout was not available probably because the circulation was observed several kilometers offshore, where the altitude at this range was approximately 1.3 km above sea level. The significance of capturing this event demonstrates the need for a network of shorter-range, higher resolution, and faster scanning of the lowest altitudes along the West coast of Puerto Rico, like that proposed by the TropiNET radar network.

4.5.1 Tropical Storm Emily

During its first year of deployment, the CSU-X25P radar was operated both locally and remotely from Colorado, during tropical storms and hurricanes. Depending on the time of year, the time difference between Puerto Rico and Colorado is +2 or +3 hours and thus operational duties are shared to maximize data collection of entire events. In early August of 2011, Tropical Storm Emily formed in the Caribbean Sea and was tracked over 200 km south of Puerto Rico. But as seen from the night time infrared satellite and NEXRAD reflectivity WunderMap images (www.wunderground.com) in Figure 4.15, Emily reached out well north of Puerto Rico bringing several north-westbound rain bands across the island, due to the counterclockwise rotation of the storm. The direction of the rain bands is further highlighted by the relative storm tracks present in the TJUA reflectivity image. The UPRM team and the author shared operational duties in the data collection of Tropical Storm Emily.

Most rain events observed by CSU-X25P, while deployed in Puerto Rico, have easterly to southeasterly winds, either dissipating over land or continuing westward toward the Dominican Republic. But as seen in the selected images from CSU-X25P during TS Emily (Figure 4.16), this was not the case. Due to its vortex centered in the Caribbean Sea and its counterclockwise rotation, the rain bands observed had southerly or southeasterly winds with velocities in the ±25m/s range. Figures 4.17(a-d) further show RHI scans where the relative
Figure 4.13: Circulation off coast of Rincon, PR on 26 April 2011 18:42:45 UTC. Images created using CSU VCHILL and Google Earth
storm movement of the entire vertical column clearly indicates tropical storm strength inbound velocities from $AZ = 118^\circ$, and outbound velocities through the $AZ = 303^\circ$ direction. A more typical RHI scan through a thunderstorm would indicate inbound velocities in the lower altitudes below 2 km, with perhaps outbound movement of the cloud tops. Due to the size and relative location of the tropical storm, one rain band crosses the entire island and covers most of the PPI scan range.

4.5.2 Hurricane Irene

Twenty days later, from August 20-29, Hurricane Irene was first tracked as a tropical storm, with unusually high sustained winds, east of the Lesser Antilles, then gradually peaking as a Category 3 hurricane with 120 mph winds, while over the Bahamas. Irene was the first and largest hurricane of the 2011 Atlantic hurricane season, eventually making landfall over New York, then moving over northern New England as seen in the storm track of Figure 4.18. Hurricane Irene caused hundreds of millions of dollars in damage, especially in Puerto Rico.
Figure 4.15: Tropical Storm Emily on 3 August 2011: (a) Infrared satellite storm track, and (b) TJUA reflectivity with relative storm tracking enabled (images created using Weather Underground’s WunderMap web application)
Figure 4.16: CSU-X25P PPI images (15:21 UTC) of Tropical Storm Emily. Note storm approach from SE to NW due to CCW rotation of Emily from south. Images created using CSU VCHILL and Google Earth.
Figure 4.17: Tropical Storm Emily, RHI scans through Inbound and Outbound rain band on 2011-08-03 at 15:57 UTC (a & c) and 15:28 UTC (b & d). For reference, cloud tops at 8-10 km; melting layer at 4.5-5 km. Images created using CSU VCHILL and Google Earth.
and the many east coast cities affected by the strong winds and heavy rain.

![Hurricane Irene storm track](image)

**Figure 4.18: Hurricane Irene storm track (using Weather Underground)**

Early on August 22, Tropical Storm Irene made landfall on the southeastern coast of Puerto Rico where it continued to track northwest across the island, all while strengthening to hurricane force winds as it returned to water. At approximately 09:00 UTC, Irene was classified as a category 1 hurricane. Figure 4.19, shows satellite, TJUA, TDWR, and CSU-X25P radar reflectivity and velocity images depicting the evolution of Irene. In the MODIS satellite image, the organization of the storm is evident as it tracks towards Puerto Rico on 20110821-1745 UTC. About 8 hours later, TJUA radar captured reflectivities in the 35-45 dBZ range as Irene was on its final approach toward the island. In this and subsequent figures, corrected reflectivity and Doppler velocity PPI scans from CSU-X25P, at Elevations of 4°, are presented at hourly intervals starting at 20110822-07:09 UTC through 20110823-
02:08 UTC. This particular PPI scan angle was selected based on minimal ground clutter that might otherwise block large sectors in the southeastern quadrant.

At approximately 20110822-03:30 UTC, the TJUA radar went offline for the next nine hours, due to power losses experienced by over 800,000 households to the eastern half of the island. With the NEXRAD radar down, the TDWR radar located in San Juan was then used to track Irene during CSU-X25P operations. Around the time of the power outages, remote operations from within the island were no longer possible, and thus the author assumed remote operations from his home in Colorado. This was possible due to proper planning of the infrastructure, ensuring ample backup power and internet access to the R&D building, on the UPRM campus.

During this evolution observed by CSU-X25P, we can see that during the first four hours, from 07:09-10:08 UTC, light rain bands came from the northeast and gradually shifted to northwesterly winds. An assumption of the general storm direction can be made based on the zero-Doppler transition region (white color), where the relative velocity here is orthogonal to the radial direction of the radar, resulting in near zero velocities. Looking at the orthogonal direction to the zero-Doppler transition, the general direction of the rain bands can be determined based on the inbound to outbound velocities from the observed PPIs.

In Figure 4.20, the transition in the direction of light rain bands shifts from northwesterly to southwesterly, from 11:05-14:10 UTC. During this time, Irene’s eye was tracking westward, almost parallel the northern coast of the island. Measured Doppler velocities were in the 18 m/s (40 mph) range, but the eye was still over 80 km away from the radar. As the eye travelled past the west coast, the measured velocities saw an increase to the 24 m/s (54 mph) range, as seen in Figure 4.21. A notable shift from southwesterly to near southerly, in the relative wind velocities of the rain band direction can be seen over the next eight hours from 15:04-22:11 UTC, in Figures 4.21 and 4.22. In Figure 4.23, we continue to see the rain band direction shift from southerly to slightly southeasterly as the eye of Hurricane Irene traveled north of the Dominican Republic, while moisture was continually pulled toward its vortex.

The CSU-X25P radar node proved to be robust and operated under high sustained wind
velocities during several heavy precipitation events, to include Tropical Storm Emily and Hurricane Irene. It was also demonstrated that the hardware and software designs met the challenge of operating in the tropical environment on the western coast of Puerto Rico. It was the author’s experience that the CSU-UPRM collaboration was successful with respect to preparedness for future deployment and operations of the TropiNet radar network.
Figure 4.19: MODIS, TJUA, & CSU-X25P CdBZ & $v_D$ PPI images of Hurricane Irene. Note shift from northeasterly to northwesterly direction in rain bands. MODIS image source: NASA/GSFC, MODIS Rapid Response. TJUA image source: Weather Underground. CSU-X25P images created with CSU VCHILL.
Figure 4.20: CSU-X25P CdBZ & \( v_D \) PPI images of Hurricane Irene (created with CSU VCHILL). Note shift from northwesterly to southwesterly direction in rain bands.
Figure 4.21: TJUA (dBZ) & CSU-X25P (CdBZ & v_D) images of Hurricane Irene. Note eye located NW of island, the southwest-erly direction in rain bands and the significant increase in velocities to 18-25 m/s. TJUA image source: Weather Underground. CSU-X25P images created with CSU VCHILL.
Figure 4.22: CSU-X25P CdBZ & $v_p$ PPI images of Hurricane Irene (created with CSU VCHILL). Note the southwesterly to southerly shift in direction, and sustained velocities in the 18-22 m/s range.
Figure 4.23: CSU-X25P CdBZ & $v_D$ PPI images of Hurricane Irene (created with CSU VCHILL). Note the southerly to southeasterly shift in direction, and sustained velocities in the 15-18 m/s range.
CHAPTER 5

Summary and Conclusions

This Thesis has described the development and validation of an X-band polarimetric Doppler weather radar, that was the first dual polarization radar to be deployed in Puerto Rico, to observe the lower troposphere in a tropical environment. The overall radar system was described, providing detailed descriptions of the major functional blocks to include: the transceiver, data acquisition system, antenna, and antenna positioner. Key modifications made to upgrade a SCR-584 pedestal with modern motors, position control, and a gigabit Ethernet rotary joint were presented and documented. Calibration procedures were separately performed on the transmitter, analog receiver, and digital receiver, while a metallic sphere calibration provided an end to end radar system calibration.

An overview of the signal processing techniques used to calculate the Doppler spectral moments from an incoherent magnetron transmitter was presented. The calculated radar parameters include the Doppler velocity, effective reflectivity, differential reflectivity, differential propagation phase, and specific differential phase. In addition, a description of the real-time clutter filtering procedure used to mitigate unwanted return echoes from ground targets was included.

Results from the system validation of the CSU-X25P with the CSU-CHILL S-band radar provided good agreement in the calculated radar parameters. The deployment of CSU-X25P, the initial test node for the TropiNet radar network, to the western coast of Puerto Rico is documented. This deployment served as a platform to strengthen the collaboration between CSU and UPRM, and equally important, to test the robustness of the TropiNet infrastructure.

Having successfully operated during Hurricane Irene and Tropical Storm Emily in 2011, the CSU-X25P has proven it can function, as designed, under adverse environmental conditions in the tropics. These and other cases were presented with example PPI and RHI
scan images that testify to the capabilities of this X-band polarimetric Doppler weather radar.
REFERENCES


APPENDIX A

Antenna Positioner Assemblies
Revision History

1.0 - Board rev 1. Includes microwave switch assembly components, DAC SHDN pin tied to +5V instead of GND
1.1 - Change P25-11 to use component M25R21X102C instead of Q2P32

Build Variants

UMAC chassis compatible

UMAC - No UMAC compatibility
Do not populate:

P1

UMAC - Draw power from UMAC
Do not populate:
P1, P10, D2B, U27, G45

Parallel I/O

NOPAR - No Parallel I/O
Do not populate:
P18-21, C30-42, J13, J14

PAR10 - With Parallel I/O
Do not populate:
Figure A.1: SCR-584 Pedestal Riser Dimensions
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### SEPT. 14, 2009

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### EL LIMIT SW

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### EL LOWER LIMIT SWITCH - NC

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### EL UPPER LIMIT SWITCH - NC

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81
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<td>+24V DC In</td>
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<tr>
<td>TB-1-2</td>
<td>+24V DC In</td>
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<tr>
<td>TB-1-3</td>
<td>+24V DC In</td>
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<tr>
<td>TB-1-4</td>
<td>+24V DC Return</td>
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<tr>
<td>TB-1-5</td>
<td>+24V DC Return</td>
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<td>TB-1-6</td>
<td>+24V DC Return</td>
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<tr>
<td>TB-2-3</td>
<td>N/C</td>
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<td>N/C</td>
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<td>L1, Servo Amp Power</td>
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<td>TB-3-3</td>
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<td>AZ Analog In +</td>
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<td>TB-4-5</td>
<td>AZ Drive, Analog Out +</td>
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<td>AZ Drive, Analog Out -</td>
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<td>TB-4-7</td>
<td>AZ Motor Power C</td>
<td>16 AWG wire</td>
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<td>TB-4-8</td>
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<td>TB-4-11</td>
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<td>TB-5-1</td>
<td>Opto 1+</td>
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<td>TB-5-3</td>
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<td>TB-5-5</td>
<td>Opto 3+</td>
<td>Run/Safe &amp; AZ Stow</td>
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<td>Opto 4-</td>
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<td>via limit switches &amp; run/safe</td>
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<td>TB-6-3</td>
<td>EL Brake +</td>
<td>EL Motor Brake + (BR 1)</td>
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<td>TB-6-4</td>
<td>EL Brake -</td>
<td>EL Motor Brake + (BR 2)</td>
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<td>EL Drive, Analog Out +</td>
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